- 1 Explaining the large variability in empirical relationships between magnetic pore
- 2 fabrics and pore space properties
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22 Summary

23 The magnetic anisotropy exhibited by ferrofluid-impregnated samples serves as a proxy for their 24 pore fabrics, and is therefore known as magnetic pore fabric. Empirically, the orientation of the 25 maximum susceptibility indicates the average pore elongation direction, and predicts the preferred 26 flow direction. Further, correlations exist between the degree and shape of magnetic anisotropy and 27 the pores' axial ratio and shape, and between the degrees of magnetic and permeability 28 anisotropies. Despite its potential, the method has been rarely used, likely because the large 29 variability in reported empirical relationships compromises interpretation. Recent work identified an 30 additional contribution of distribution anisotropy, related to the arrangement of the pores, and a 31 strong dependence of anisotropy parameters on the ferrofluid type and concentration, partly 32 explaining the variability. Here, an additional effect is shown; the effective susceptibility of the 33 ferrofluid depends on the measurement frequency, so that the resulting anisotropy depends on measurement conditions. Using synthetic samples with known void geometry and ferrofluids with 34 known susceptibility (4.04 SI and 1.38 SI for EMG705 and EMG909, respectively), magnetic 35 36 measurements at frequencies from 500 Hz to 512 kHz are compared to numerical predictions. 37 Measurements show a strong frequency-dependence, especially for EMG705, leading to large discrepancies between measured and calculated anisotropy degrees. We also observe artefacts 38 39 related to the interaction of ferrofluid with its seal, and the aggregation of particles over time. The 40 results presented here provide the basis for a robust and quantitative interpretation of magnetic 41 pore fabrics in future studies, and allow for re-interpretation of previous results provided that the 42 ferrofluid properties and measurement conditions are known. We recommend that experimental 43 settings are selected to ensure a high intrinsic susceptibility of the fluid, and that the effective 44 susceptibility of the fluid at measurement conditions is reported in future studies.

45 Keywords

- 46 Magnetic fabrics and anisotropy
- 47 Permeability and porosity
- 48 Magnetic properties

49 1. Introduction

50 Magnetic pore fabrics (MPF) have been proposed as a fast and efficient way to characterize the 51 anisotropy of pore space in rocks (Pfleiderer and Halls, 1990), and to predict permeability anisotropy 52 and preferred flow directions (Pfleiderer and Halls, 1994, Hailwood et al., 1999). They are defined as 53 the anisotropy of magnetic susceptibility (AMS) of ferrofluid-impregnated samples, and may reflect 54 depositional or tectonic fabrics (Pfleiderer and Kissel, 1994, Hailwood and Ding, 2000, Parés et al., 55 2016). As pore fabrics control fluid flow in porous media, their accurate description is important in 56 many areas of geophysics and geology, including convective flow models, aquifer and reservoir 57 characterization, geothermal energy and CO₂ storage applications (Ayan et al., 1994, Huang et al., 58 2017, Ijeje et al., 2019, Panja et al., 2021, Sinan et al., 2020, Wang et al., 2014, Wang et al., 2019, 59 Willems et al., 2017, Storesletten, 1998). Traditional pore characterization methods such as X-ray 60 tomography face trade-offs between sample size and resolution, and generate large amounts of 61 data that need to be processed (Cnudde and Boone, 2013, Landis and Keane, 2010). For applications 62 that require characterization of the average pore fabric, MPFs provide a promising alternative in that 63 they describe the average pore fabric as a single second-order tensor, measured on a representative 64 sample volume, and potentially capturing pores down to 10 nm, without being affected by mineral 65 and grain boundary properties unlike seismic anisotropy (Robion et al., 2014, Almqvist et al., 2011, 66 Pfleiderer and Halls, 1990, Benson et al., 2003).

67 Correlations between average pore axial ratio and MPFs have been proposed and investigated since 68 the earliest MPF studies, using both natural and synthetic samples (Pfleiderer and Halls, 1990, 69 Pfleiderer and Halls, 1993, Hrouda et al., 2000, Jones et al., 2006, Jezek and Hrouda, 2007, Nabawy 70 et al., 2009). Additionally, MPFs were compared to other measures of pore space anisotropy, e.g. 71 anisotropy of elastic properties or electrical conductivity (Louis et al., 2005, Robion et al., 2014, 72 Benson et al., 2003, Nabawy et al., 2009). Although reported empirical relationships for fabric 73 orientation are similar for all studies (maximum susceptibility indicating the average pore elongation 74 direction and maximum permeability), there is a large variability in reported relationships between

75 MPF anisotropy degree and pore aspect ratio or degree of permeability anisotropy (Fig. 1) 76 (Pfleiderer and Halls, 1990, Pfleiderer and Halls, 1993, Pfleiderer and Halls, 1994, Louis et al., 2005, 77 Jones et al., 2006, Nabawy et al., 2009). Therefore, quantitative and robust interpretation of MPF 78 data is not yet possible, and while the method is promising, it has been used rarely. For the method 79 to become more widely applied, understanding the variability between reported empirical 80 relationships is crucial, and the goal of this paper. The basis for interpreting the empirical 81 relationships reported in rocks is to understand the fundamentals, and this is achieved here on 82 synthetic samples with simple and known pore geometries.

83 One explanation for the large variability in empirical relationships is that different types of 84 ferrofluids at different concentrations have been used when these relationships were established. In 85 the meantime, it has become evident that for a given pore axial ratio, the MPF anisotropy degree 86 increases nonlinearly with increasing fluid susceptibility (Biedermann, 2019, Jones et al., 2006). The 87 same applies to correlations with permeability anisotropy, which are further complicated by the fact 88 that only few MPF studies report full permeability tensors based on six independent measurements 89 (Pfleiderer and Halls, 1994, Hailwood et al., 1999), whereas measurements along only two or three 90 directions parallel to the macroscopic fabric are more common (Benson et al., 2003, Louis et al., 91 2005, Nabawy et al., 2009). If the number of measurements is lower than that needed to define the 92 full tensor, the calculated anisotropy underestimates the true anisotropy, unless the measurement 93 directions coincide with the principal axes of the tensor.



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95 Fig. 1: (a) Simplified empirical relationships between magnetic pore fabric (MPF) and average pore

alignment, or permeability anisotropy; (b,c) Literature data from which empirical relationships were
 derived show large scatter.

99 Secondly, the impregnation process and associated changes in the pore space properties may result 100 in differences between studies. Two standard methods are used for impregnation, (1) evacuating the 101 pore space under vacuum conditions and then supplying ferrofluid (Parés et al., 2016, Pfleiderer and 102 Halls, 1990, Benson et al., 2003, Robion et al., 2014, Hrouda et al., 2000), or (2) injecting the 103 ferrofluid under pressure, which leads to different fabrics depending on the injection pressure 104 (Esteban et al., 2006). It is not clear, however, whether this is related to smaller pores being 105 impregnated at higher pressure, or the destruction of pore walls during impregnation. Additional 106 impregnation methods are being tested (Pugnetti et al., 2021).

107 A third reason for the variability is that the MPF data has been largely compared to the average pore 108 axial ratio, shape and orientation, i.e., assuming that MPFs are controlled by shape anisotropy 109 (Pfleiderer and Halls, 1990, 1993, Hrouda et al., 2000, Jones et al., 2006, Jezek and Hrouda, 2007). 110 Shape anisotropy results from self-demagnetization, a process that occurs when a strongly magnetic 111 body with a high intrinsic susceptibility k_{int} (e.g., an ore body, magnetite grain, or ferrofluid-filled 112 pore) is surrounded by weakly magnetic material (e.g., rock) (Clark and Emerson, 1999). Selfdemagnetization reduces the observed susceptibility k_{obs} to $k_{obs} = (I + k_{int}N)^{-1}k_{int}$, where I is the 113 114 unit matrix, and N the self-demagnetization tensor, which depends on the shape of the strongly 115 magnetic body (e.g. Clark, 2014). It can be easily calculated for ellipsoids (Osborn, 1945, Stoner, 116 1945), and approximated for other simple body shapes (Sato and Ishii, 1989, Joseph, 1966, Joseph, 117 1967). However, self-demagnetization tensors may change throughout a body of complex shape 118 (Joseph, 1976, Joseph and Schlömann, 1965). In addition to the shape preferred orientation of single 119 pores, also their arrangement controls the measured MPF. Rocks contain numerous pores in a 120 complex and irregular three-dimensional network, and distribution anisotropy, arising from 121 magnetostatic interaction of the ferrofluid in different pores, also contributes to the measured 122 anisotropy (Biedermann, 2019, Biedermann, 2020). Distribution anisotropy has been extensively 123 investigated for magnetite grains in rocks (Grégoire et al., 1998, Grégoire et al., 1995, Hargraves et 124 al., 1991, Cañón-Tapia, 1996, Cañón-Tapia, 2001, Stephenson, 1994), and is described in a similar

way for MPFs (Biedermann, 2019, Biedermann, 2020). Thus, the MPF depends not only on the pores'
shape preferred orientation as proposed initially, but also on the distribution of the pores
throughout the rock. The mathematical treatment of distribution anisotropy in MPF studies relies on
the assumption that the fluid susceptibility is homogeneous throughout the pore space, and that
impregnated pores possess similar magnetic properties to solid grains of the same susceptibility.
Recent work is testing these models by comparing measured MPFs to predictions based on pore
characterization using X-ray microtomography (Zhou *et al.*, 2021).

132 Finally, measurement conditions, specifically frequency, may affect MPF results. Ferrofluids are 133 colloidal suspensions of magnetic nanoparticles in non-magnetic water- or oil-based carrier fluid. 134 The nanoparticles are coated with surfactant to avoid agglomeration, and their size of ~10 nm 135 ensures they are kept in suspension by Brownian motion (Odenbach, 2004, Joseph and Mathew, 136 2014, Torres-Diaz and Rinaldi, 2014, Rosensweig, 1987, Rosensweig, 1988, Papaefthymiou, 2009). 137 Magnetite particles in this size range behave superparamagnetically at room temperature, and their 138 susceptibility is frequency-dependent (Söffge and Schmidbauer, 1981, Muscas et al., 2013, Néel, 139 1949, Bean and Livingston, 1959, Brown, 1959, Stephenson, 1971, Dormann, 1981, Jones and 140 Srivastava, 1989, Coffey and Kalmykov, 2012). This characteristic is exploited in environmental 141 magnetism, where frequency-dependence of susceptibility is used to infer grain size distributions 142 (Dearing et al., 1996, Eyre, 1997, Worm, 1998, Worm and Jackson, 1999, Hrouda, 2011). Out-of-143 phase susceptibility is a second property related to frequency dependence, and also used for 144 magnetic granulometry (Hrouda et al., 2013). Other possible sources of frequency-dependence and 145 out-of-phase susceptibility are eddy currents or low-field hysteresis, observed in pyrrhotite and Ti-146 magnetite (Jackson, 2003-2004, Kosterov et al., 2018, Hrouda et al., 2013, Jackson et al., 1998). 147 Physical motion of particles in response to the magnetic field may play an additional role (Brown, 148 1959, Brown, 1963, Dormann, 1981). Brownian motion is constrained by the pore walls, and may be 149 restricted in certain pores due to their size. If this affects frequency dependence, it may help to 150 distinguish between fabrics of different pore size fractions. Frequency-dependent properties and

151 out-of-phase susceptibility are thus expected for the ferrofluid used in MPF studies. Of particular 152 interest here is whether the frequency-dependence of susceptibility also affects the anisotropy. One 153 indication that this may be the case is a large variability in effective anisotropy constants of 154 magnetite nanoparticles depending on whether the measurements were obtained in DC or AC fields 155 (Goya et al., 2003). Unfortunately, neither the intrinsic susceptibility of the fluid, nor the 156 measurement frequency have been reported in most MPF studies. Even though the frequency can 157 sometimes be estimated from the instrument used, the lack of information on fluid susceptibility 158 makes it impossible to compare results and empirical relationships between studies. Thus frequency-159 dependence and its potential effect on anisotropy and MPF interpretations remain to be 160 investigated.

161 This study characterizes MPFs and their frequency dependence in synthetic samples with a range of 162 pore sizes, aspect ratios, and arrangements. Measurements obtained at a range of frequencies are 163 compared to numerical models taking into account shape and distribution anisotropy. Models are 164 based on the initial susceptibilities given in the fluids' technical specifications. Differences between 165 expected and effective susceptibilities and related discrepancies between models and 166 measurements are discussed. The term 'expected susceptibility' is used here to describe the 167 susceptibility calculated from the initial susceptibility and shape of the fluid-filled void. 'Effective 168 susceptibility' is used to describe the actually measured susceptibility. Both expected and effective 169 susceptibilities refer to observables and are affected by self-demagnetization, i.e., they depend on 170 the shape of the void. They should be equal if the intrinsic susceptibility of the fluid at measurement conditions equals the initial susceptibility reported in the fluid's technical specifications. A major 171 172 finding of this work is the strong decrease of effective ferrofluid susceptibility with frequency, in 173 particular for water-based ferrofluid EMG705, with important consequences for the interpretation 174 of MPFs. The experiments shown here also identify difficulties and unwanted effects that may 175 complicate the interpretation of MPFs in rocks.

176 2. Material and Methods

177 2.1 Samples

Two sets of samples have been prepared for this study. The first group (label prefix ZK) contains one 178 179 cylindrical pore with a ratio diameter:height equal to 1:4, and four different sizes, defined by 180 cylinder diameters of 0.6 – 5 mm. In the second group (label D.T., where the number after D indicates the diameter, and the number after T the cylinder height), each sample contains a set of 9 181 182 cylindrical pores, with different samples having diameter:height ratios of 1:2, 1:4 and 1:8, and 183 diameters of 0.5 mm, 1 mm, and 2 mm (Fig. 2). The ZK sample group was used to investigate the 184 effects of ferrofluid type and concentration, as well as testing different types of sealing. The 185 anisotropy parameters of the four different sizes should in theory be equal for the same ferrofluid 186 and concentration, so that these samples allow to investigate size-dependent effects. Conversely, 187 the main purpose of the D.T. samples is to investigate the interplay of shape and distribution 188 anisotropies for different configurations of filled pores. Therefore, MPFs on the D.T. sample series 189 were measured using a single ferrofluid and a single concentration.

190 The ZK samples were prepared from a 1-inch diameter polycarbonate cylinder, using an HSS/CNC 191 drill at the Institute of Geological Sciences, University of Bern. The samples were prepared such that 192 the diameter:height ratio and the expected MPF is the same for all samples, although the void 193 volume and therefore mean susceptibility are different. The volume-effect can be removed by 194 normalizing all magnetic data by the ferrofluid volume rather than the sample volume. Initially, eight 195 sets of samples comprising four sizes each were drilled. These were filled with water- and oil-based 196 ferrofluids, EMG705 and EMG909, respectively, at 1:10, 1:20, 1:25 and 1:50 volume concentrations of ferrofluid to carrier liquid. Attempts to dilute the ferrofluids at a ratio 1:100, as used in Parés et al. 197 198 (2016), failed due to aggregation of the particles, and their precipitation before the fluid could be 199 filled into the samples. The initial susceptibilities of EMG705 and EMG909 are reported as 4.04 (SI) 200 and 1.38 (SI) (EMG 705 Specifications and Physical Properties;

201 https://ferrofluid.ferrotec.com/products/ferrofluid-emg/water/emg-705/ and EMG909

- 202 Specifications and Physical Properties <u>https://ferrofluid.ferrotec.com/products/ferrofluid-</u>
- 203 emg/oil/emg-909/). The measured susceptibilities of the carrier fluids are -1.0(±0.1)*10⁻⁵ and -
- 204 1.6(±0.1)*10⁻⁵ (SI) for water and oil respectively, orders of magnitude lower than those of the
- 205 ferrofluid, and thus negligible. The diluted ferrofluids have nominal intrinsic susceptibilities ranging

206 from 0.03 to 0.4 (SI).

(a) ZK samples: single pore, constant ratio major/minor axis (4:1), four sizes - effect of ferrofluid type and concentration



(b) D.T. samples: multiple pores, three aspect ratios and sizes - effect of distribution anisotropy and impregnation efficiency

D1 (diameter 1 mm) D05 (diameter 0.5 mm) D2 (diameter 2 mm) Filling and measurement sequence _____ '_₽₽[₽] Height/diameter 8:1 Height/diameter 4:1 Height/diameter 2:1 · 00 ^{ال} 0 יח D2T4 D1T2 D05T1 а d **1**00 000000000 <mark>॔╻╹╻╹╻</mark>╹ D2T8 D1T4 D05T2 b <u>hould by a state of the state </u> <u>──</u>──────── ╷╹╻╹╻╹ D1T8 D05T4 D2T16 с

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Fig. 2: Pore dimensions and assemblies of filled pores for the studied samples. The scientific questions
addressed with each sample group were different: (a) ZK samples served the purpose of investigating
the influence of ferrofluid properties and sample preparation; (b) the D.T. samples allowed to

211 characterize the effects of pore shape and arrangement.

213 Prior to MPF measurements, each void was sealed with different materials, including tape or various 214 kinds of glue. Sealing with tape was unsuccessful, because it was hardly possible to prevent the 215 formation of air bubbles, and because the fluid migrated into the space between the cylinder surface 216 and the tape, likely due to capillary forces, over timespans of hours. Glued seals showed mixing 217 artefacts, i.e., a small portion of the ferrofluid would diffuse into the glue while the glue was drying 218 (Fig. 3). This was particularly problematic for oil-based ferrofluid and the largest voids, which 219 required most glue and therefore long drying times. Additionally, oil-based ferrofluid would react 220 with the glue and destroy its sealing capacities over timeframes of a few days.

A second set of ZK samples was then drilled, and these were filled with special care to prevent air bubbles or diffusion of ferrofluid outside the void. To achieve this, the voids were sealed with hot glue that dries faster than normal glue thus minimizing interaction with ferrofluid, and a combination of hot glue with a plastic plate containing two smaller holes to allow exchange of air during filling and sealing, while at the same time reducing the amount of glue and drying time.

226 Despite all precautions taken, trapped air could not be avoided completely, and in all samples, air

227 bubbles appeared to develop over time.

228 The D.T. samples were prepared from polycarbonate, using CNC milling machines at the Physics 229 Institute, University of Bern. A total of nine cubic samples were made, with three aspect ratios 230 (diameter:height ratios of 1:2, 1:4 and 1:8), and three sizes (2 mm, 1 mm and 0.5 mm diameter). 3x3 231 voids were drilled in a single face of the cube, at 8 mm distance from each other. To reduce the 232 number of samples that needed to be prepared, the voids of each sample were filled sequentially, 233 measuring the MPF before filling the next void(s). This procedure allowed to obtain six datasets from 234 each of the nine samples. Based on the experience with the ZK samples, water-based EMG705 235 ferrofluid diluted with distilled water at 1:10 was used to fill the voids, and hot glue for sealing. 236 Water-based fluid is less prone to particle aggregation and sedimentation, and interacts less with 237 glue than oil-based fluid, and this stability over time was important for the chosen sequence of filling

- and measuring the different sets of voids one after the other. A lower-case letter at the end of the
- sample name indicates the pattern of filled voids.



(a) Sample preparation and preparation-related artefacts: Interaction of ferrofluid and seal

(b) Sample preparation and preparation-related artefacts: Nanoparticle aggregation and sedimentation



(c) Changes in sealing capabilities three weeks after sample preparation



Water-based fluid: Good sealing

(d) Changes in fluid configuration over time

Oil-based fluid: Poor sealing, ferrofluid leaks



- 241 Fig. 3: Sample preparation and preparation-related artefacts: Conceptual sketches, pictures, and
- 242 influence on measured anisotropy for (a) migration of ferrofluid along the sample-seal interface or
- 243 mixing of ferrofluid with seal and formation of air bubbles; (b) particle aggregation and
- 244 sedimentation over time. (c) Changes in sealing capabilities three weeks after sample preparation,
- resulting in ferrofluid leakage for oil-based EMG909; and (d) changes in fluid configuration over time,
- 246 *affecting interpreted pore shapes.*

247 All voids have been drilled from a single side of the sample, to simplify the manufacturing process, 248 resulting in asymmetric positioning of the void(s) within the cylinder or cube. A possible effect of the 249 sample asymmetry on the measured anisotropy was tested by repeat measurements with slightly 250 different sample positions. For the MFK1-FA, susceptibility measurements were independent of 251 sample position, indicating that the field inside the coil of the MFK1-FA is homogeneous on the scale 252 of the sample size and position variation. The large noise level of the SM150H/L instruments for 253 repeat measurements with a given sample positioning outweighs any potential variation resulting 254 from changes in sample positioning. Hence, the sample asymmetry does not affect the measured 255 susceptibilities.

256 2.2 Expected magnetic properties

The expected magnetic properties for the configurations shown in Fig. 2 were calculated based on the known initial susceptibilities for the respective ferrofluid and its concentration, and the known pore shape, as well as the pore arrangement in the case of the D.T. samples.

260 Each of the ZK samples contains a single ferrofluid-filled cylindrical void with equal diameter/height 261 ratio, so that their anisotropies are defined solely by shape anisotropy, and the demagnetization 262 tensor is the same for each of them. Due to the sample geometry, the maximum susceptibility is 263 expected along the z axis (cylinder axis), and there is a minimum susceptibility plane normal to that 264 axis. The self-demagnetization factors along the three sample axes are $N_x = N_y > N_z$, and using the 265 equation for cylinders given by Sato and Ishii (1989), $N_x = N_y = 0.450$ and $N_z = 0.0997$. Had an ellipsoidal approximation been used (Osborn, 1945), the self-demagnetization factors would have 266 267 been $N_x = N_y = 0.462$, $N_z = 0.0754$. Expectations of the observed directional susceptibilities depend 268 on the self-demagnetization tensor and fluid susceptibility, and the same is true for the expected 269 anisotropy parameters (Table 1, Fig. 4a). The susceptibility anisotropy is described by the directional 270 susceptibilities k_x , k_y , and k_z , and their ratios. Additionally, the anisotropy degree P =271 $\max(k_x, k_y, k_z)/\min(k_x, k_y, k_z)$ and anisotropy shape $U = (2^* \text{median}(k_x, k_y, k_z) - \max(k_x, k_y, k_z))$ 272 $\min(k_x,k_y,k_z)/(\max(k_x,k_y,k_z) - \min(k_x,k_y,k_z))$ were used, analogously to P and U calculated from the

eigenvalues of the susceptibility tensor (Jelinek, 1981). Note that we are not using the standard
notation in these equations, because *P* and *U* are defined based on the eigenvalues, and with only
three directional measurements, it is in general not possible to define the full tensor nor its
eigenvalues. Nevertheless, given the symmetry of the samples, *k_x*, *k_y* and *k_z* are measured parallel to
the expected principal susceptibility directions, and thus represent the eigenvalues.

All D.T. samples apart from the (a) series possess both shape and distribution anisotropy, and their

expected directional susceptibilities were computed using the FinIrrSDA code (Biedermann, 2020).

280 Because the spacing between the voids is constant for all samples, independent of void size, the

281 distribution anisotropy contribution leads to different total anisotropies even when the shape

anisotropies are the same (Table 1). For the (a) series, the anisotropy is equivalent to that expected

for the ZK samples in that $k_z > k_x = k_y$. Interactions lead to a slight increase of k_x compared to k_y in the

(b), (c) and (d) series, also affecting the shape of the anisotropy. For the configurations of these

285 samples, the *P*-values are mainly defined by the aspect ratios of each void, while the distribution

anisotropy has a smaller effect on the *P*-value, but largely affects the anisotropy shape *U* (Fig. 4b).

287 2.3 Magnetic measurements

The magnetic properties of the ZK samples had been measured prior to preparing the D.T. samples, and the results obtained for the ZK sample series were used the select suitable preparation and measurement sequences for the D.T. sample series. Therefore, the experiments performed on each series differ from each other.

- Table 1: Expected magnetic properties for the ZK (a) and D.T. (b) sample series. Directional
- susceptibilities (k_x , k_y , k_z) normalized by ferrofluid volume, and anisotropy indicated by ratios of

294 directional susceptibilities, anisotropy degree, and anisotropy shape. Initial susceptibilities: 4.04 (SI)

295 for water-based EMG705, and 1.38 (SI) for oil-based EMG909.

a) Expect	ed suscentibility (n	ormalized by fe	errofluid volume) fo	n 7K samnle							
u) Expect	cu susceptionity (ii	Ferrofluid	Void	Magnetic	nore fabric	naramete	rs				
		concentration	diameter:height	kx	kv	kz	v/x	z/x	z/v	Р	U
	water-based ferro	1:10	1:4	0.315	0.315	0.354	1.000	1.124	1,124	1.124	-1.000
		1:20	1:4	0.177	0.177	0.189	1.000	1.066	1.066	1.066	-1.000
		1:25	1:4	0.145	0.145	0.153	1.000	1.054	1.054	1.054	-1.000
		1:50	1:4	0.076	0.076	0.079	1.000	1.028	1.028	1.028	-1.000
	oil-based ferroflu	1:10	1:4	0.119	0.119	0.124	1.000	1.043	1.043	1.043	-1.000
		1:20	1:4	0.064	0.064	0.065	1.000	1.023	1.023	1.023	-1.000
		1:25	1:4	0.052	0.052	0.053	1.000	1.019	1.019	1.019	-1.000
		1:50	1:4	0.027	0.027	0.027	1.000	1.009	1.009	1.009	-1.000
b) Expect	ed susceptibility (n	ormalized by fe	errofluid volume) fo	or D.T. samp	les (water-	based flui	d)				
		Ferrofluid	Void	Magnetic	pore fabric	paramete	rs				
		concentration	diameter:height	kx	ky	kz	y/x	z/x	z/y	Р	U
	a series (all)	1:10	1:2	0.319	0.319	0.344	1.000	1.078	1.078	1.078	-1.000
		1:10	1:4	0.315	0.315	0.354	1.000	1.124	1.124	1.124	-1.000
		1:10	1:8	0.313	0.313	0.360	1.000	1.152	1.152	1.152	-1.000
	b series, D2	1:10	1:2	0.319	0.319	0.344	1.000	1.078	1.078	1.078	-0.994
		1:10	1:4	0.315	0.315	0.354	1.000	1.124	1.124	1.124	-0.993
		1:10	1:8	0.313	0.313	0.360	0.999	1.151	1.152	1.151	-0.988
	b series, D1	1:10	1:2	0.319	0.319	0.344	1.000	1.078	1.078	1.078	-0.999
		1:10	1:4	0.315	0.315	0.354	1.000	1.124	1.124	1.124	-0.999
		1:10	1:8	0.313	0.313	0.360	1.000	1.152	1.152	1.152	-0.998
	b series D05	1:10	1:2	0.319	0.319	0.344	1.000	1.078	1.078	1.078	-1.000
		1:10	1:4	0.315	0.315	0.354	1.000	1.124	1.124	1.124	-1.000
		1:10	1:8	0.313	0.313	0.360	1.000	1.152	1.152	1.152	-1.000
	c series, D2	1:10	1:2	0.320	0.319	0.344	0.997	1.076	1.078	1.076	-0.932
		1:10	1:4	0.316	0.315	0.354	0.995	1.118	1.124	1.118	-0.915
		1:10	1:8	0.315	0.312	0.359	0.990	1.139	1.151	1.139	-0.862
	c series, D1	1:10	1:2	0.319	0.319	0.344	1.000	1.078	1.078	1.078	-0.992
		1:10	1:4	0.315	0.315	0.354	0.999	1.123	1.124	1.123	-0.989
		1:10	1:8	0.313	0.313	0.360	0.999	1.150	1.152	1.150	-0.983
	c series D05	1:10	1:2	0.319	0.319	0.344	1.000	1.078	1.078	1.078	-0.999
		1:10	1:4	0.315	0.315	0.354	1.000	1.124	1.124	1.124	-0.999
		1:10	1:8	0.313	0.313	0.360	1.000	1.152	1.152	1.152	-0.998
	d series, D2	1:10	1:2	0.320	0.319	0.344	0.998	1.075	1.078	1.075	-0.941
		1:10	1:4	0.316	0.315	0.353	0.996	1.118	1.123	1.118	-0.927
		1:10	1:8	0.315	0.312	0.359	0.991	1.139	1.149	1.139	-0.880
	d series, D1	1:10	1:2	0.319	0.319	0.344	1.000	1.078	1.078	1.078	-0.993
		1:10	1:4	0.315	0.315	0.354	0.999	1.123	1.124	1.123	-0.991
		1:10	1:8	0.313	0.313	0.360	0.999	1.150	1.151	1.150	-0.985
	d series D05	1:10	1:2	0.319	0.319	0.344	1.000	1.078	1.078	1.078	-0.999
		1:10	1:4	0.315	0.315	0.354	1.000	1.124	1.124	1.124	-0.999
		1:10	1:8	0.313	0.313	0.360	1.000	1.152	1.152	1.152	-0.998
	e series, D2	1:10	1:2	0.320	0.320	0.344	1.000	1.076	1.076	1.076	-1.000
		1:10	1:4	0.316	0.316	0.353	1.000	1.118	1.118	1.118	-1.000
		1:10	1:8	0.314	0.314	0.358	1.000	1.139	1.139	1.139	-1.000
	e series, D1	1:10	1:2	0.319	0.319	0.344	1.000	1.078	1.078	1.078	-1.000
		1:10	1:4	0.315	0.315	0.354	1.000	1.123	1.123	1.123	-1.000
		1:10	1:8	0.313	0.313	0.360	1.000	1.150	1.150	1.150	-1.000
	1 8 8 8										
	e series D05	1:10	1:2	0.319	0.319	0.344	1.000	1.078	1.078	1.078	-1.000
		1:10	1:4	0.315	0.315	0.354	1.000	1.124	1.124	1.124	-1.000
		1:10	1:8	0.313	0.313	0.360	1.000	1.152	1.152	1.152	-1.000
	6	4.40	4.2					. ar-			
	t series, D2	1:10	1:2	0.320	0.320	0.343	1.000	1.075	1.075	1.075	-1.000
		1:10	1:4	0.316	0.316	0.353	1.000	1.116	1.116	1.116	-1.000
		1:10	1:9	0.314	0.314	0.357	1.000	1.135	1.135	1.135	-1.000
	6	1.10	4.2								
	t series, D1	1:10	1:2	0.319	0.319	0.344	1.000	1.078	1.078	1.078	-1.000
		1:10	1:4	0.315	0.315	0.354	1.000	1.123	1.123	1.123	-1.000
		1:10	1:9	0.313	0.313	0.360	1.000	1.150	1.150	1.150	-1.000
			1.0								
	t series D05	1:10	1:2	0.319	0.319	0.344	1.000	1.078	1.078	1.078	-1.000
		1:10	1:4	0.315	0.315	0.354	1.000	1.124	1.124	1.124	-1.000
		1:10	1:8	0.313	0.313	0.360	1.000	1.152	1.152	1.152	-1.000



297

Fig. 4: Model results for (a) ZK sample series, with constant aspect ratio of the void and variable fluid
 concentration, and (b) D.T. sample series, using water-based fluid at 1:10 concentration, three aspect
 ratios (AR, diameter:height) of each void, and a series of filled void distributions.

301

302 2.3.1 ZK sample series

303 AC susceptibility was measured in several fields, frequencies, and along two or three axes of the

- sample coordinate system (k_x , k_y and k_z , or k_x and k_z ; where the z-axis is along the cylinder axis, and
- 305 the *x* and *y*-axes oriented arbitrarily in the plane perpendicular to that axis). In combination with
- 306 constraints from the known sample geometry (rotational symmetry around z), two directions are
- 307 sufficient to determine the magnetic anisotropy tensor, and three directions allow to estimate data
- 308 quality. Additional estimates of data quality were obtained from repeat measurements.

309 Two instruments were initially used for directional susceptibility measurements, (1) the MFK1-FA 310 susceptibility bridge with three frequencies (976 Hz, 3904 Hz and 15616 Hz), and field ranges 2-706 311 A/m at 976 Hz, 2-356 A/m at 3904 Hz and 2-218 A/m at 15616 Hz; and (2) the SM150H/L 312 susceptometers with nominal frequency range 63 Hz – 512 kHz. Of these, a reduced range between 313 500 Hz – 512 kHz provided usable results, whereas lower frequencies were subject to large noise 314 levels. Measurements on the SM150H/L were conducted in 80 A/m, the highest field available at all 315 frequencies. For anisotropy determination on the MFK1-FA, a field of 200 A/m was used and 3-5 316 repeat measurements were taken for each direction, while field-dependence measurements contain 317 one measurement per field, and the measurement uncertainty is calculated from the measurements 318 at the previous and subsequent fields. On the SM150H/L susceptometers, a total of 20 repeat 319 measurements were necessary due to the larger instrumental noise level. 320 The magnetization of a sample exposed to an AC field can vary in-phase with the field, or be subject 321 to a phase shift. This phase shift may arise from viscous relaxation in small particles, electrical eddy 322 currents in conductive materials or weak-field hysteresis (Jackson, 2003-2004, Hrouda et al., 2017). 323 While the SM150H/L system measures the component of susceptibility in-phase with the field, the 324 MFK1-FA also provides information on the phase shift. However, note that unlike on the KLY5 325 kappabridge, the zero phase is not calibrated on the MFK1 kappabridges. Here, the phase measured 326 for the samples was corrected with the phase measured on the calibration sample with known zero 327 phase, according to the method outlined in Hrouda et al. (2015).

The diamagnetic susceptibility of the holder and polycarbonate sample cylinder were subtracted from all measurements as background. After background-correction, the measured susceptibility was normalized by (1) sample volume and (2) void volume. The directional susceptibilities and the anisotropy of susceptibility, as described by the ratios of directional susceptibilities as well as *P* and *U* values were then compared to the expected values.

333 2.3.2 D.T. sample series

- Based on the results obtained from the ZK samples, a subset of the above measurements was
- 335 selected for the D.T. sample series. All measurements were performed on the MFK1-FA, and
- directional susceptibilities (k_x , k_y , k_z) were measured at a field of 200 A/m, and frequencies of 976 Hz,
- 337 3904 Hz and 15616 Hz. Three to five repeat measurements were used to estimate directional
- 338 susceptibilities and the measurement noise. Prior to filling the voids successively with ferrofluid, the
- and the same conditions, and later subtracted as background. The
- 340 results are reported as directional susceptibilities normalized by sample or void volume,
- susceptibility ratios, or the anisotropy parameters *P* and *U* as for the ZK samples.
- 342 3. Results
- **343** 3.1 ZK samples
- 3.1.1 First set of ZK samples: Frequency and field dependence, instrumental noise
 The susceptibility of samples filled with water- and oil-based ferrofluid is frequency-dependent, and
- 346 the decrease in susceptibility with increasing frequency is stronger in the water-based ferrofluid
- 347 compared to the oil-based ferrofluid (Fig. 5a; Table A, Supplementary Material). For the
- 348 measurements obtained on the SM150H/L instruments, the susceptibility at 512 kHz is 35% 43% of
- that at 500 Hz for ZK1 and ZK2 filled with water-based fluid. Conversely, the susceptibility of the oil-
- based fluid decreases less, to 74-77 % of the value at 500 Hz. The corresponding measurements of
- 351 ZK3 and ZK4 are not interpretable due to the large instrumental noise level. The ratio of
- 352 susceptibility at 16 kHz to that at 1 kHz (k_{16}/k_1) as measured on the SM150H/L varies between 64%
- and 69% (water-based ferrofluid), and between 92% and 95% (oil-based fluid) for ZK1 and ZK2. The
- lower noise level of the MKF1 susceptibility bridge allows to analyse all samples, and the ratios k_{16}/k_1
- are 65%-66% for the majority of measurements for the water-based ferrofluid, and 85% 102% for
- 356 the oil-based ferrofluid. Compared to the initial susceptibilities from the fluids' technical
- 357 specifications, the median effective susceptibilities at ~1kHz and ~16kHz are ~31% (k_1/k_i) and ~21%
- (k_{16}/k_i) for the water-based fluid EMG705, while they are approximately 130% and 120% for the oil-

based fluid EMG909. These ratios were calculated by comparing the measured directional susceptibilities to the expected directional susceptibilities reported in Table 1. Note that the ratios k_1/k_i and k_{16}/k_i are subject to large variability and also time-dependence; e.g., the effective susceptibility of oil-based fluid at 1:50 is approximately 3 times the expected value, likely related to the increased magnetic interactions when particles are aggregated.

The measurements obtained on both instruments agree within measurement uncertainty (Fig. 5b).

The lower susceptibility of ZK1 filled with water-based fluid as measured on the MFK1-FA compared

to the SM150H/L is related to loss of fluid from the cavity due to capillary motion of the fluid under

the seal (cf Fig. 3). The noise level of the SM150H/L is several orders of magnitude larger than that of

the MFK1-FA, and reaches >100% of the measured susceptibility. In comparison, variability for

369 repeat measurements on the MFK1-FA is maximum 1-5% of the measured susceptibility (Fig. 5c).

370 Therefore, although the SM150H/L is preferable in that it covers a larger frequency range, the

371 quality of the data is not sufficient to analyse the anisotropy of the samples studied here, with

expected *P*-values of 1.05 (water-based ferrofluid at 1:25) or 1.02 (oil-based fluid at 1:25).

373 Phase shifts are observed for both the oil-based and water-based ferrofluids (Fig. 5d). For both types

of fluid, the phase is generally larger at ~4 kHz compared to 1 or 16 kHz. Samples filled with water-

based fluid show higher phase shifts (up to ~16°) and more variability than those filled with oil-based

376 fluid, where the maximum phase is ~8°.

Field-dependence measurements show that susceptibility and to a lesser degree phase are
independent of applied field (Fig. 6). Some jumps are observed in the field-dependent data, but
these are caused by changes in the dynamic range of the instrument and not sample properties.





Fig. 5: Frequency dependence of susceptibility and phase shift for the ZK samples: Frequencydependence of susceptibility normalized by sample (a) and void volume (b). Measurement
frequencies for the MFK1-FA are ~1 kHz, ~4 kHz and ~16 kHz, and for the SM150H/L, measurements
were performed at 0.5 kHz, 1 kHz, 4 kHz, 16 kHz, and 512 kHz. (c) Measurement uncertainty given as
percentage of standard deviation and mean of repeated measurements. Dashed lines indicate a
'standard' 1% noise level desirable for anisotropy measurements, and 5% noise, reflecting the
expected anisotropy of the water-based fluid filled ZK samples. (d) Phase shift of the susceptibility.





Fig. 6: Field dependence of susceptibility (a) and phase shift (b) for the ZK1 samples at frequencies ~1,
~4 and ~16 kHz. Note that the jumps in susceptibility and associated larger errors are related to
changes in the dynamic range of the instrument during the measurement (indicated by arrows).

404	samples. At the same time, additional aspects were investigated, e.g., the influence of ferrofluid
403	Based on the initial results, a reduced number of measurements was conducted on the second set of
401	3.1.2 Second set of ZK samples: Anisotropy and its dependence on concentration and frequency.
400	avoid artefacts as much as possible during preparation.
399	some samples. Anisotropy was therefore studied on the second set of samples, taking special care to
398	the fluid trapped between the sample and seal appears to dominate the observed anisotropy in
397	mixed with glue, and to a lesser degree trapped air inside the void. Although volumetrically small,
396	shown in Fig. 3, particularly fluid trapped at the interface between sample cylinder and seal, or fluid
395	is not consistently observed in this initial dataset. The most likely explanation are the artefacts
394	Note that although the maximum susceptibility is expected along the sample <i>z</i> -axis, with $k_x = k_y$, this

405 concentration, and time-dependence. Although preparation artefacts could not be prevented
406 entirely, they were less prominent than in the first set of samples, allowing the analysis of the
407 anisotropy of the ferrofluid-filled voids. Anisotropy depends both on the ferrofluid type,
408 concentration and the measurement frequency. Additionally, measured anisotropy changes over
409 time, reflecting the formation of air bubbles or particle aggregation.

410 Samples filled with water-based ferrofluid generally show the expected behaviour of $k_z > k_x$, $k_x \sim k_y$ 411 (Fig 7a,b; Table A, Supplementary Material). Only the samples with the smallest voids, ZK4, display a 412 behaviour contrary to expectation, in that their maximum susceptibility is often not parallel to the 413 long axis of the cylinder. A possible explanation is that these small voids are the hardest to fill, and 414 any fluid outside the void, or air bubbles trapped inside the void may outweigh the anisotropy of the 415 fluid in the void itself. ZK4 results will thus not be interpreted further. For ZK1, ZK2 and to a lesser 416 extent ZK3, a trend of stronger anisotropy degree (approximated here by k_z/k_x) for higher 417 concentration of ferrofluid is observed; however, the k_z/k_x ratios are lower than predicted in Table 1. 418 Most samples show similar properties for the initial and repeat measurements after 25 days, but 419 sometimes the degree of anisotropy decreased over time. This is interpreted here as the partial loss 420 of fluid from the void, and its migration to the interface between plastic cylinder and glue. The 421 anisotropy shapes are mostly prolate, as expected from the sample geometry (Fig. 7c). 422 MPFs of samples filled with oil-based fluid show large variability, and different behaviours for 1:50 423 concentration compared to 1:10 or 1:20. At concentrations of 1:10 or 1:20, the measured k_z/k_x ratios 424 cluster loosely around the expected values, except for the repeat measurements after 25 days, 425 which have significantly larger k_z/k_x ratios. Conversely, oil-based fluid concentrations of 1:50 lead to

426 k_z/k_x mostly < 1, opposite of the expected behaviour, and the ratios decreased further after 25 days
427 (Fig. 7a). This observation can be explained by the aggregation and precipitation of particles that is
428 strongest at the 1:50 concentration, leading ultimately to an oblate body of precipitated particles at
429 the bottom of the prolate void. Particle aggregation and precipitation is also accompanied by a

- 430 change of the anisotropy shape, from mostly prolate (as expected) at 1:10 or 1:20 concentration, to
- 431 oblate at a concentration of 1:50 (Fig. 7c). The MPFs of the samples filled with 1:50 oil-based
- 432 ferrofluid are clearly dominated by artefacts resulting from particle precipitation and are not related
- 433 to the sample geometry. Hence, they are not discussed further.



⁴³⁴

<sup>Fig. 7: Magnetic anisotropy of ferrofluid-filled voids in the second set of ZK samples, measured at 976
Hz (F1), 3904 Hz (F2), and 15616 Hz (F3). (a) Ratio of directional susceptibilities k_z to k_x, compared to
the expected value from the model shown in Fig. 4a. (b) Ratio of directional susceptibilities k_y to k_x.
Due to sample symmetry, susceptibilities should be equal along these two directions, indicated by the
dashed line. (c) Anisotropy shape U, which should be -1 (dashed line) according to sample symmetry.
Measurements at F1 were repeated after the sample had been stored for 25 days, and symbol sizes
reflect the size of the void in the ZK1, ... ZK4 samples.</sup>



442

Fig. 8: Comparison of the measured ratios k_z/k_x , k_z/k_y and $2*k_z/(k_x+k_y)$ with the modelled P-value in the frequency range ~1 kHz to ~16 kHz, for different ferrofluid concentrations, 1:10 (a), 1:20 (b), and 1:50 (c). Due to the symmetry in the x-y-plane, all these measured parameters are approximations of the modelled anisotropy degree, they should be equal and their variability indicates deviation from ideal behaviour. Large variability is observed in oil-based samples due to artefacts such as particle aggregation and poor sealing.

449

450 Higher measurement frequencies almost always result in weaker anisotropy degrees for samples

451 filled with water-based EMG705 (Figs 7a and 8). This observation can be directly related to the

452 decrease in mean susceptibility with increasing frequency (cf Fig. 5). The frequency-dependence of

- 453 the MPF anisotropy degree is less clear for samples filled with oil-based EMG909. A possible
- 454 explanation for this is that the effective susceptibility shows overall a smaller frequency-dependence
- 455 for the oil-based fluid. A second possibility is that the time-dependent artefacts (particle
- 456 aggregation, chemical reaction between oil-based ferrofluid and glue destroying the seal)
- 457 dominantly control the magnetic results of these samples.

458 3.2 D.T. samples

459 The anisotropy results of the D.T. samples show the combined effects of shape and distribution 460 anisotropies for different aspect ratios of the voids, and as successively more voids are filled with 461 EMG705 water-based ferrofluid (Fig. 9; Table B, Supplementary Material). The measured anisotropy 462 shapes are mainly prolate, similar to expectation, except for the samples with the smallest voids (0.5 463 mm diameter), and measurements at 15616 Hz (F3) in the a, b, and c series. The reason the D05 464 samples display MPF shapes different from the expected sample geometry or results of the D1 and 465 D2 samples containing larger voids, are preparation artefacts (cf Fig 3). Similar to the ZK4 samples, 466 the voids in the D05 samples are not filled completely with ferrofluid, making the spatial variation of 467 susceptibility and magnetization more complex than the sample geometry. The deviation in shape 468 for the a, b, and c series measured at F3 are likely related to these measurements having the largest 469 noise level. The susceptibility and its anisotropy are expected to be lowest at F3, and the smaller the 470 number of filled voids, the smaller the expected susceptibility. Measured shapes for corresponding 471 samples in the D1 and D2 series are almost identical. For a given measurement frequency and 472 configuration of filled voids, the degree of anisotropy increases with the aspect ratio of the 473 individual voids. The measured degree of anisotropy is almost always smaller than that expected, 474 and generally decreases nonlinearly with increasing measurement frequency (Fig. 10).



Fig. 9: Modelled and measured anisotropy degree P and shape U for the D.T. samples, filled with
water-based EMG705 fluid at concentration 1:10. Measurements were performed at 3 frequencies,
976 Hz (F1), 3904 Hz (F2), and 15616 Hz (F3). Preparation artefacts are observable in the D05 sample
series.





481 Fig. 10: Frequency dependence of the anisotropy degree P for the D.T. samples filled with EMG705
482 water-based ferrofluid at 1:10 concentration. Measurements were performed at the standard

483 frequencies of the MFK1-FA kappabridge, 976 Hz, 3904 Hz, and 15616 Hz. Preparation artefacts are

484 visible in the D05 series. The results for the D1 and D2 samples indicate a decrease of anisotropy

485 *degree with increasing measurement frequency.*

486 4. Discussion

4.1 Frequency-dependence leads to discrepancies between expected and measured 487 488 anisotropy A major finding of this study is the strong frequency-dependence of the magnetic susceptibility of 489 490 the ferrofluid and its anisotropy, in particular for the water-based EMG705. This frequency-491 dependence leads to large discrepancies between the initial susceptibilities reported in the fluid's 492 technical specifications, valid for measurements in weak DC fields, and the effective properties at 493 standard measurement conditions for MPF studies. Self-demagnetization and hence shape 494 anisotropy increase nonlinearly with the intrinsic fluid susceptibility. Observed deviations between 495 expected and measured anisotropy degrees in samples filled with EMG705 are clearly related to the 496 effective fluid susceptibility being lower than the specified initial susceptibility. Analogously, the 497 measured anisotropy degree of samples filled with EMG909 should be slightly higher than that 498 modelled, due to the higher effective susceptibility. However, this deviation is low compared to the 499 variability in the data, so that no unambiguous conclusion can be drawn. Measurements of the 500 EMG909-filled samples after 25 days show stronger anisotropies, possibly associated with an 501 increase in fluid susceptibility resulting from particle aggregation. In general, the measured MPF of a 502 given sample will depend on the measurement frequency in addition to ferrofluid type and 503 concentration, and all of these parameters (or the effective fluid susceptibility at measurement 504 conditions) need to be known before MPFs can be interpreted quantitatively. The intrinsic 505 susceptibility at measurement conditions can be determined by measuring directional 506 susceptibilities of fluid in a void of known shape and dimensions under the same conditions, and then calculating $k_{int} = (I + k_{obs}N)^{-1}k_{obs}$ (e.g. Clark, 2014). Note that k_{obs} approaches N^{-1} for 507 508 large k_{int} , so that high intrinsic susceptibilities cannot be measured reliably. In this case it may be 509 helpful to measure diluted fluid and then calculate k_{int} of the undiluted fluid from k_{int} of the 510 diluted fluid and the dilution ratio. In any case, we recommend to do several repeat measurements 511 of k_{obs} , as any uncertainty will be amplified when calculating k_{int} , especially when N is large along 512 the measurement direction.

513 Only one MPF study so far specified the measurement frequency used, which was 976 Hz (Parés et 514 al., 2016). In some additional studies, the frequencies can be estimated from the instrument 515 capabilities, and range from 750 Hz to 920 Hz (Pfleiderer and Halls, 1990, Pfleiderer and Halls, 1993, 516 Pfleiderer and Halls, 1994, Hrouda et al., 2000, Benson et al., 2003, Jones et al., 2006, Esteban et al., 517 2006, Robion et al., 2014, Humbert et al., 2012). Where neither the instrument nor measurement 518 frequency are specified (Pfleiderer and Kissel, 1994, Louis et al., 2005, Almqvist et al., 2011), or 519 where instruments with several operating frequencies were used (Nabawy et al., 2009), anisotropy 520 degrees are not interpretable, and derived empirical relationships not comparable to other studies. 521 Measurement frequencies just below 1 kHz appear most common in MPF studies, and it would be 522 desirable to define a universal relationship between the initial susceptibility of a ferrofluid and its 523 effective susceptibility at ~ 1 kHz. This would allow correction of previously published results and 524 empirical relationships for differences between the initial and effective fluid susceptibilities, and 525 facilitate modelling in future studies. However, the two ferrofluids used here, EMG705 and EMG909, 526 show largely different characteristics: for EMG705, the effective susceptibility at ~1 kHz is ~20-35% 527 of its initial susceptibility, and for EMG909 the ratio of effective to initial susceptibility is ~125-150%. 528 A consequence of this is that although the technical specifications (ferrotec.com) indicate a larger 529 susceptibility for EMG705 than EMG909, the effective susceptibilities at 976 Hz, 3904 Hz and 15616 530 Hz are higher for EMG909. Hence, the frequency-dependence and ratio of effective to initial 531 susceptibilities need to be measured for each ferrofluid used in MPF studies.

Some previous MPF studies used the same fluids that are investigated here, EMG705 (Pfleiderer and Halls, 1990, Pfleiderer and Halls, 1993, Pfleiderer and Halls, 1994, Pfleiderer and Kissel, 1994), and
EMG909 (Robion *et al.*, 2014, Parés *et al.*, 2016). Other fluids used include EMG905 (Hrouda *et al.*,
2000, Benson *et al.*, 2003, Jones *et al.*, 2006, Almqvist *et al.*, 2011), EMG507 (Robion *et al.*, 2014),
and EMG509 (Humbert *et al.*, 2012), and the type of ferrofluid was not always specified (Hailwood *et al.*, 1999, Louis *et al.*, 2005, Esteban *et al.*, 2006, Nabawy *et al.*, 2009). It is possible that all water-

538 based fluids show a similar frequency-dependence as EMG705, and all oil-based fluids a behaviour 539 similar to EMG909. If this is the case, it would explain why the empirical correlations between the 540 MPF anisotropy degree and pore aspect ratios are steeper in the study of Jones et al. (2006) 541 compared to (Pfleiderer and Halls (1990, 1993); it could be an effect of the higher effective fluid 542 susceptibilities in the former study compared to the later. The correlations with respect to 543 permeability are not comparable because essential information on the fluid properties and 544 measurement frequencies are missing. More work will be needed to determine the effective 545 properties of all ferrofluids used in MPF studies, and to systematically compare empirical 546 relationships.

547 4.2 Origin of the frequency-dependence

548 The measurements show frequency-dependent susceptibility and a phase shift, but no significant 549 field-dependence (cf. Figs 5 and 6). Both frequency-dependence and phase shift are larger for 550 samples filled with EMG705 water-based fluid compared to the EMG909 oil-based fluid. Three 551 mechanisms have been described to cause frequency-dependence and out-of-phase susceptibility: 552 (1) viscous relaxation of superparamagnetic particles, (2) eddy currents in conductive materials, and 553 (3) weak-field hysteresis (Jackson, 2003-2004, Hrouda et al., 2013, Jackson et al., 1998, Néel, 1949, 554 Brown, 1959, Dormann, 1981). The absence of any field-dependence (cf Fig. 6) makes it possible to 555 exclude weak-field hysteresis as a phenomenon occurring in the samples investigated here (Hrouda 556 et al., 2013). The electrical conductivities of both ferrofluids at concentration 1:25 were measured at 557 the Petrophysics Laboratory at the University of Bern; EMG705 has an electrical conductivity of 2 mS 558 while that of EMG909 is not measurable, so that eddy currents likely do not contribute to dissipation 559 in these samples. Therefore, the observed frequency-dependence and out-of-phase susceptibilities 560 are a result of viscous relaxation, either by Néel relaxation or Brownian motion.

Both Néel and Brownian relaxation times vary with particle volume. The Néel relaxation time τ_N is computed as $\tau_N = \tau_0 \exp(KV/kT)$, where τ_0 is a time constant, K the anisotropy constant, V the

particle volume, and kT the thermal energy, being the product of Boltzman's constant and temperature (Néel, 1949). The time constant τ_0 is not truly constant as it depends on particle size, coercivity and temperature; and varies between $0.4*10^{-9}$ s to $3.3*10^{-9}$ s for magnetite particles with sizes decreasing from 15 nm to 5 nm. A value of $1*10^{-9}$ s agrees with experimental evidence (Worm, 1998). The volume of a particle with 10 nm diameter is $5.24*10^{-25}$ m³, measurements were done at room temperature (298 K), and the K_1 anisotropy constant for magnetite is $1.35*10^4$ J/m³ (Syono, 1965). For these parameters, $\tau_N = 5.6*10^{-9}s$.

570 The Brownian relaxation time has been described as $\tau_B = \left(\left(\frac{\gamma K}{M}\right) * \sqrt{KV/kT\pi} * \exp\left(-\frac{KV}{kT}\right)\right)^{-1}$, 571 where γ is the gyromagnetic constant, 1.76085963*10¹¹ (sT)⁻¹, and M the magnetisation (Jones and 572 Srivastava, 1989). We use 480 kA/m, i.e., the saturation magnetization of magnetite. For the 10 nm 573 particles in the fluid, this leads to $\tau_B = 1.5 * 10^{-9}s$.

574 These calculations indicate the both relaxation times have the same order of magnitude, and the 575 Brownian relaxation is slightly faster. At the same time, Jones and Srivastava (1989) state that for small particles (<10⁶ atoms), Néel's model provides more physically reasonable results. They also 576 found both relaxation times to be similar for particles with 30 nm diameter, which is different from 577 the results obtained here. Söffge and Schmidbauer (1981) report typical relaxation times of τ_B = 578 $10^{-3}s$ and $\tau_N = 10^{-9}s$ for a ferrofluid with magnetite particles of 10 nm diameter. These values 579 580 agree with the calculation shown here for τ_N , but not for τ_B . A reason for the discrepancy could be that they calculated $\tau_B = \frac{3\eta V}{kT}$, where η is the fluid viscosity. For the fluids used here, $\eta < 5*10^{-3}$ Pa*s 581 for EMG705, and η = 3*10⁻³ Pa*s for EMG909, leading to $\tau_B = 1.1 - 1.9 * 10^{-6} s$, which is closer to 582 583 their results, but still does not agree. Using this value, Brownian relaxation is significantly slower 584 than Néel relaxation, indicating that Néel relaxation likely dominates. However, the large variability 585 between calculations and studies indicates that these relaxation times need to be interpreted with care. Note that both relaxation times depend on particle size, so that the particle aggregation that 586 587 was observed in the ferrofluids over time will affect the results. Because particle aggregation

appears fastest in strongly diluted EMG909 ferrofluid, the strongest artefacts are expected there.
The lower stability of oil-based ferrofluid over time was also observed by Robion (pers. comm.). To
ensure a stable magnetic response over time, we would thus recommend using water-based
ferrofluid, or ferrofluid at higher concentration, in magnetic pore fabric studies.

592 Worm and Jackson (1999) describe the frequency-dependence of susceptibility due to viscous relaxation: $X = (\mu_0 V M_s^2) / (3kT(1 + (2\pi f)^2 \tau^2))$, where f is the measurement frequency. Using 593 $\tau_N = 5.6 * 10^{-9} s$, V = 5.24*10⁻²⁵ m³, and M_s = 480 kA/m results in no observable frequency-594 595 dependence at frequencies between 1-16 kHz. This may be related to the uncertainty in the 596 relaxation times, or indicate that the magnetic particle size is smaller than 10 nm. The observed 597 difference in frequency-dependence of EMG705 and EMG909, despite them having the same 598 nominal size of 10 nm, may indicate that their actual sizes differ from each other, with the effective 599 size of the EMG705 particles being smaller, and therefore showing a stronger frequency-dependence 600 than EMG909 and predicted for 10 nm particles.

4.3 Implications for the interpretation of MPF and impregnation efficiency data

602 Empirical relationships for MPFs focus on fabric orientation and anisotropy degree. The anisotropy 603 degree, and in particular the variation of the P-value (and analogously the L- and F-values) with 604 intrinsic fluid susceptibility are discussed here. For a given fluid susceptibility, the MPF anisotropy 605 degree increases nonlinearly with pore aspect ratio, and corresponding empirical relationships have 606 been used to efficiently estimate the average pore shapes or their alignment (Hrouda et al., 2000, 607 Jones et al., 2006). However, the measured P-value also increases nonlinearly with ferrofluid 608 susceptibility for any given pore aspect ratio. The effective intrinsic susceptibility depends on the 609 type and concentration of ferrofluid, as well as the measurement frequency. Thus, all published empirical relationships depend on the effective intrinsic susceptibility of the respective fluids, so that 610 a comparison between studies is not straightforward. For example, the MPF anisotropy degrees for a 611 given pore shape reported in Jones et al. (2006) tend to be higher than those of Pfleiderer and Halls 612

613 (1990), which are in turn larger than those in Pfleiderer and Halls (1993). A partial explanation for 614 this is that Jones *et al.* (2006) used a fluid with initial susceptibility $k_i = 1.09$, whereas Pfleiderer and 615 Halls (1993) used a fluid with $k_i = 0.8$ (the susceptibility of the fluid in Pfleiderer and Halls (1990) is 616 not specified). However, careful comparison between the literature data and models indicates that 617 the data by Jones *et al.* (2006) more closely matches the model for $k_{int} = 2$, and that of Pfleiderer and 618 Halls (1993) resembles a model with $k_{int} < 0.5$ (Fig. 11a,b). Thus, the comparison between measured 619 and modelled *P*-values indicates that the effective *k*_{int} of the EMG705 fluid in Pfleiderer and Halls 620 (1993) is lower, while the effective k_{int} of the EMG905 ferrofluid in the study by Jones *et al.* (2006) is 621 higher than the reported initial susceptibilities. A lower than expected value could be due to dilution 622 of the fluid, but the higher effective value can only be explained by a higher effective intrinsic 623 susceptibility, related to frequency-dependent susceptibility. The measured k_{int} at ~1 kHz ranges 624 from ~20 - ~35% of the expected value for EMG705 (using ZK1 and ZK2 measurements shown in Fig. 625 2), and ~125% - ~150% of the expected value for EMG909. Hence, the frequency-dependence of the 626 ferrofluid susceptibility strongly affects the measured P-values, as well as the empirical relationships 627 between pore shape and MPF anisotropy degree. The combined dependence of the measured 628 anisotropy degree on the effective k_{int} and the pore shape is summarized in Fig. 11c. The influence of 629 the measurement frequency is further illustrated in Fig. 11d, which shows that the expected P-value 630 for a prolate ellipsoidal void filled with pure EMG705 is >2.5. In contrast, at the typical ~1kHz 631 measurement frequency, the P-value of the same void and fluid is only ~1.5. At higher frequencies, 632 the P-values are even lower. The frequency-related decrease in P is comparable to lowering the 633 ferrofluid concentration to 1:25. These models additionally illustrate that higher fluid concentrations 634 and lower measurement frequencies make it possible to detect weak anisotropies that may be 635 below the detection limit if strongly diluted ferrofluids or high measurement frequencies are used. 636 The relevance of the ferrofluid concentration has been described previously (Jones et al., 2006, 637 Biedermann, 2019), and the results presented here clearly indicate that measurement frequency 638 may be equally relevant for the interpretation of MPF data.



Fig. 11: (a) The influence of ferrofluid susceptibility on the MPF – pore shape relationships, where
shaded areas reflect variability from rotationally oblate to rotationally prolate ellipsoidal pores with
a given axial ratio. Published data are shown on top of the models. (b) Expected MPF P-values as
function of fluid susceptibility. Data as in (a), but only datasets with known intrinsic susceptibility are
shown. Numbers next to the data points reflect their axial ratios. (c) Variation of anisotropy degree
as a function of pore axial ratio and ferrofluid intrinsic susceptibility combined, and effect of non-zero
measurement frequency (d) or ferrofluid concentration (e).

647 Although the fabric orientation is not directly affected by the frequency-dependence of the 648 ferrofluid susceptibility, the lower anisotropy degrees at higher frequencies may lead to lower signal 649 to noise ratios. This in turn can result in higher uncertainty and larger confidence ellipses of the 650 principal MPF directions. When noise levels are high, this results in unrealistically large P, L and F-651 values, and also large uncertainties in anisotropy shape (Biedermann et al., 2013), and this may 652 explain some of the deviations from the generally observed trends in the D.T. a, b and c series for 653 high frequencies. For most reliable estimations of pore fabrics, ferrofluids with high effective 654 intrinsic susceptibilities at the measurement conditions are preferred. For EMG705 and possibly 655 other water-based fluids, a frequency of ~1 kHz is favourable compared to 4 kHz or 16 kHz. If an 656 instrument with a comparable noise level at lower frequencies becomes available in the future, this 657 may provide even better MPF results. For EMG909 and potentially other oil-based fluids, the 658 frequency-dependence is less pronounced, so that measurements at different frequencies can be 659 more easily compared. In both cases, highly concentrated ferrofluids are preferable in that they 660 provide higher susceptibilities and stronger anisotropies that are easier to characterize.

Empirical relationships also exist between MPFs and permeability anisotropy. There, the
interpretation is more challenging, as in addition to the considerations above, permeability depends
on the volumetrically small connections between pores, while MPFs are dominantly defined by the
shape and orientation of larger pores. Anisotropic permeability also controls the migration of fluid
during impregnation, and may influence MPF results when impregnation efficiency is < 100%.
Further work is needed to fully characterize this effect.

In rock samples, impregnation efficiency, i.e., the percentage of the pore space that is filled with
ferrofluid, is a crucial parameter that defines the quality and reliability of MPF-based interpretations.
Impregnation efficiency has been determined using either the increase in mass, or the increase in
susceptibility after impregnation compared to the dry sample, and by visual inspection on cut
surfaces (Robion *et al.*, 2014, Parés *et al.*, 2016, Almqvist *et al.*, 2011). If impregnation efficiency is

672 evaluated based on susceptibility changes, then the measured susceptibility is compared to the 673 expected susceptibility computed as the product of pore space volume and initial susceptibility of 674 the (diluted) ferrofluid. This would be an accurate estimate if the effective susceptibility of the fluid 675 were equal to the initial susceptibility. However, when the effective susceptibility at measurement 676 conditions deviates from the initial susceptibility used to estimate impregnation efficiency, the 677 obtained results are misleading. The effective susceptibility of EMG705 at 1 kHz is ~35% of the initial 678 susceptibility stated in the fluid's technical specifications. Thus, when the entire pore space is 679 impregnated, but the difference between initial and effective susceptibility are not considered, the 680 calculated impregnation efficiency (35%) would significantly underestimate the proportion of 681 impregnated pore space. At the same time, when the effective susceptibility is higher than the initial 682 susceptibility, the impregnation efficiency would be overestimated unless effective susceptibilities 683 are used in the calculation. For EMG909, impregnation of 66-80% of the pore space seemingly leads 684 to an impregnation efficiency of 100%, when it is calculated based on the fluid's initial susceptibility. 685 This may lead to the wrong conclusion that oil-based ferrofluid is more efficient at impregnating rock 686 samples, and different ways of characterizing impregnation efficiency should be used in 687 combination. In fact, oil-based ferrofluid has been described as more efficient at impregnation than 688 water-based fluid, but based on weight changes (Robion et al., 2014).

689 4.4 Artefacts and recommendations for sample preparation

In addition to providing explanations for the large variability in published empirical relationships, the data presented here also give insight into some of the challenges that arise from working with colloidal suspensions rather than solid magnetite grains in rock samples. These include (1) the fluid not occupying the supposed space or moving out of that space after preparation, during the measurement or during storage, (2) trapping of air, (3) aggregation of the particles in the fluid, so that the susceptibility and magnetization vary within the fluid-occupied space. Some of these are specifically related to the type of synthetic samples investigated here, but others are more related to

- the fluid itself, and we will discuss how the observed artefacts may manifest in rocks, and formulate
- 698 some recommendations for studies on impregnated rock samples (Fig. 12).

(a) Ideal distribution of ferrofluid



Macroscale: Fluid with homogeneous properties

(b) Observed artefacts transferred to rocks



Fluid film at sample surface



Particle clusters at magnetite grains

(c) Related challenges to be investigated in rocks



Particles at grain surfaces



Nanoscale: Dispersed particles, non-magnetic carrier



Gravity affecting particle distribution within each pore



Particle aggregation and filtering



Impregnation hindered by narrow pore throats

- Fig. 12: (a) Ideal distribution of ferrofluid in a rock's pore space on the macroscale (homogeneous
- fluid, used for modelling MPFs) and nanoscale (dispersed magnetic nanoparticles in non-magnetic
- carrier fluid). (b) Conceptual sketches of how the artefacts observed in synthetic samples may
- transfer to rocks: Fluid at the sample-seal interface could translate to a film of ferrofluid at the
- sample surface; particle aggregation and sedimentation could translate to gravity affecting the
- distribution of nanoparticles within each pore, nanoparticles clustering around magnetite grains in
- the rock, or particle aggregation and filtering effects during the impregnation process. (c) Related
- challenges that may occur in rocks include particles clustering at grain surfaces, rather than being

distributed uniformly throughout the pore space, or parts of the pore space not being impregnated
 due to bottlenecks at narrow pore throats, again resulting in a non-uniform distribution of magnetic
 particles throughout the pore space.

711

712 Magnetic pore fabrics in rocks are interpreted assuming that the entire pore space is filled with a homogeneous fluid of constant susceptibility. On the nanoscale, this corresponds to a random 713 714 distribution of the magnetic nanoparticles throughout the pore space (Fig. 12a). A major issue we 715 observe in our synthetic samples is ferrofluid moving out of the voids and migrating along the 716 interface of the seal (tape, glue, plastic plate) and the sample cylinder, with the result that air 717 bubbles form inside the voids (cf. Fig. 3). This appears especially important for oil-based fluid that 718 destroyed all our seals over time. The migrated fluid at the interface naturally possesses strong 719 anisotropy as it is constrained by two flat surfaces in contact with each other. The strong shape 720 anisotropy of a small amount of migrated fluid may outweigh the weaker anisotropy of the fluid 721 inside the void, and thus dominate the measurements. Similarly, a film of ferrofluid on the surface of 722 a rock sample could significantly affect the measured MPF (Fig. 12b). This is especially problematic 723 when the sample is wrapped in e.g. plastic, which is advisable to protect the instrument from 724 contamination. Parés et al. (2016) encapsulated their samples in plastic boxes, preventing them from touching the walls using plastic spacers, to minimize buildup of a fluid film. Alternative approaches 725 726 may include advanced sealing of surface pores, e.g. using varnish, or the evaporation of the carrier 727 fluid after impregnation, to immobilize the particles. Surface effects are less prominent in larger 728 samples that have the additional advantage of being more representative of the rock volume. 729 However, large samples are harder to impregnate, so that smaller samples are preferred in MPF studies (Almqvist et al., 2011, Parés et al., 2016). Therefore, the selection of sample size will always 730 731 be a compromise between impregnation efficiency and the reduction of surface artefacts. 732 The air replacing the fluid in the synthetic samples diminishes the anisotropy of the fluid in the void,

as it decreases its aspect ratio. This effect was especially relevant for voids of small volume, i.e., the

734 ZK3, ZK4 and D05 series. Trapped air within the void can also move during a measurement, and be 735 located at different positions for each measurement direction. If this happens, different fluid shapes 736 and thus shape anisotropies are measured in each direction and that data should strictly not be used 737 to calculate one common anisotropy tensor. Some air bubbles moved during sample storage (cf Fig. 738 3). In rocks, trapped air may block entry to some pores, thus diminishing impregnation efficiency, 739 and if located in larger pores, cause deviations between the fluid distribution and the pore space. It 740 is possible for example, that gravity causes all nanoparticles, especially when aggregated, to 741 concentrate at the bottom of each pore, while air is located at the top of the pores. This 742 inhomogeneous particle and susceptibility distribution would affect the measured magnetic 743 anisotropy. More work is needed to investigate the distribution of the magnetite particles within a 744 single pore or throughout the pore space, and reduce these artefacts. 745 Particle aggregation and sedimentation is another process that leads to inhomogeneous distribution 746 of susceptibility and magnetization throughout the fluid, and may further affect the magnetic 747 properties due to increased interaction between particles. We have mainly seen aggregation in the 748 oil-based EMG909 ferrofluid at the lowest concentration used (1:50), leading to measured 749 anisotropies opposite to the shape of the void. This is important in light of previous studies that used 750 even lower concentrations of 1:100 to avoid saturating the signal of the susceptibility meter (Parés 751 et al., 2016), and who favoured oil-based ferrofluid over water-based ferrofluid due to its better 752 impregnation properties (Robion et al., 2014). A careful evaluation of impregnation properties, 753 instrument capabilities and potential artefacts is necessary prior to choosing the most appropriate ferrofluid. In rocks, particles may not only aggregate with each other, but could cluster at magnetite 754 755 grains, or other minerals with high susceptibility. Due to their larger diameter, particle aggregates 756 will have more difficulties migrating through narrow pore throats and impregnating small pores, 757 decreasing the number of pores that are reached, and diminishing the resolution and usefulness of 758 the MPF method. Our results illustrate that higher fluid susceptibilities lead to stronger anisotropies,

and that higher concentrations and water-based fluids may be less prone to particle aggregation
 compared to lower concentrations and oil-based fluids.

761 Additional challenges that may arise in rocks, but are not observed here in the synthetic samples are 762 that particles may be concentrated at the grain surfaces rather than distributed throughout the 763 pores, and that impregnation may be hindered in parts of the sample due to narrow pore throats 764 (Fig. 12c). The latter is particularly important when particles are aggregated and thus larger than the 765 nominal 10 nm diameter. Incomplete impregnation in the core of the sample is a recurring issue in 766 MPF studies (Almqvist et al., 2016; Robion et al., 2014). Parés et al. (2016) have prepared smaller 767 samples to overcome this issue. Additional characteristics of the rock, including pore tortuosity and 768 wettability of the minerals, which are known to affect fluid flow (Abdallah et al., 2007; Clennell,

1997; Ghanbarian *et al.*, 2013), will also affect the impregnation process.

770 Note that the quality of the magnetic data, as defined by confidence ellipses and F-statistics (Hext, 771 1963, Jelinek, 1977) or R_1 values from repeat measurements (Biedermann *et al.*, 2013), provides no 772 information on the presence or severity of any of these artefacts. Here, artefacts could be identified 773 from deviations between models and measurements, e.g. for the D05 samples, and because the 774 synthetic samples are transparent, the trapped air and migrated fluid along the sample-seal 775 interface are observed. It remains to be investigated how important similar artefacts are in rocks, 776 where they are not as easily identifiable. Most of these artefacts seem to worsen over time, so that 777 we would generally recommend to measure MPFs right after sample impregnation for best results.

5. Conclusions

This study investigated the magnetic susceptibility and anisotropy of ferrofluid-filled voids in
synthetic samples. Computations of shape and distribution anisotropies were compared to low-field
AMS measured at different frequencies. The measured anisotropy of voids filled with EMG705
water-based ferrofluid is generally lower than that predicted by models based on the initial
susceptibility of the fluid, and decreases with increasing measurement frequency. For EMG909 oil-

based fluid, the measurements agree more closely with the models, and there is a weaker
frequency-dependence. These observations can be explained by shape anisotropy being a function
of the fluid's effective intrinsic susceptibility, which shows a strong frequency-dependence for
EMG705, and weaker frequency-dependence for EMG909.

788 For a single pore of a given shape, the measured degree of magnetic anisotropy increases 789 nonlinearly with the fluid susceptibility. The effective intrinsic fluid susceptibility depends on the 790 type of ferrofluid and its concentration, and the measurement frequency. This frequency-791 dependence is important to take into account when defining and interpreting empirical relationships 792 between the MPF anisotropy degree and pore aspect ratios. Analogously, the fluid properties and 793 measurement conditions affect empirical relationships between MPF and permeability anisotropy, 794 and may complicate the determination and comparison of impregnation efficiency between 795 different fluids.

796 The details of the frequency-dependence vary between the two ferrofluids used here, EMG705 and 797 EMG909. It is possible that other oil-based ferrofluids have a similar frequency-dependence as 798 EMG909, and other water-based fluids behave analogously to EMG705, but it is also possible that 799 each fluid displays its own frequency-dependence. Hence, the initial susceptibilities as specified by 800 the manufacturer do not reflect the effective susceptibilities at measurement conditions, and there 801 is no universally applicable relationship between initial and effective susceptibilities. These findings 802 may explain in part or full the variability of reported empirical relationships. We therefore 803 recommend that measurement frequencies, and frequency-dependence of susceptibilities are 804 reported in future MPF studies, in addition to ferrofluid type and concentration. Detailed 805 information on magnetic properties of ferrofluids at actual measurement conditions of MPF studies 806 will help correct for frequency-related variations in MPF parameters, facilitating the comparison 807 between empirical relationships reported in different studies. The results shown here present an 808 important step towards a quantitative and robust interpretation of MPF data in terms of pore fabric

- 809 properties or permeability anisotropy. Thus, we believe that the method is more applicable to fluid
- 810 migration studies in the future, facilitating the characterization of reservoirs and aquifers, and
- 811 supporting convective flow, geothermal energy and CO₂ sequestration applications.
- 812 Higher effective intrinsic susceptibilities enable the description of weak pore fabrics, and simplify the
- 813 distinction between different fabric strengths. Optimal parameters can be achieved by selecting the
- 814 appropriate type and concentration of ferrofluid, as well as measurement frequency. A final
- 815 recommendation for future MPF studies is related to the migration of fluid and particle aggregation
- 816 that occur over time. Translated to rocks, these processes may cause artefacts leading to erroneous
- 817 pore fabric interpretation, including formation of a film of ferrofluid on the sample surface or
- 818 between the sample surface and sample holder, aggregation of particles inside pores, clustering of
- 819 nanoparticles at magnetite grains, or filtering effects during impregnation. To avoid these artefacts,
- 820 we consider it best to measure MPFs within a few days after impregnation.

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829 Data Availability

- 830 All data is reported in the paper and supplementary materials. The FinIrrSDA model used here for
- the predictions has been previously published (Biedermann, 2020), and the code is available on
- 832 https://zenodo.org/record/4040785.

833 References

- Abdallah, W., Buckley, J.S., Carnegie, A., Edwards, J., Herold, B., Fordham, E., Graue, A., Habashy, T.,
 Seleznev, N., Signer, C., Hussain, H., Montaron, B., Ziauddin, M., 2007. Fundamentals of
 wettability. *Oilfield Review* 19, 44-61.
- Almqvist, B.S.G., Mainprice, D., Madonna, C., Burlini, L. & Hirt, A.M., 2011. Application of differential
 effective medium, magnetic pore fabric analysis, and X-ray microtomography to calculate
 elastic properties of porous and anisotropic rock aggregates, *Journal of Geophysical Research-Solid Earth*, 116.

- Ayan, C., Colley, G., Ezekwe, E., Wannell, M., Goode, P., Halford, F., Joseph, J., Mongini, A.,
 Obondoko, G. & Pop, J., 1994. Measuring permeability anisotropy: The latest approach, *Oilfield Review*, 6, 24-35.
- 844 Bean, C.P. & Livingston, J.D., 1959. Superparamagnetism, *Journal of Applied Physics*, 30, S120-S129.
- Benson, P.M., Meredith, P.G. & Platzman, E.S., 2003. Relating pore fabric geometry to acoustic and
 permeability anisotropy in Crab Orchard Sandstone: A laboratory study using magnetic
 ferrofluid, *Geophysical Research Letters*, 30, 1976.
- Biedermann, A.R., 2019. Magnetic pore fabrics: the role of shape and distribution anisotropy in
 defining the magnetic anisotropy of ferrofluid-impregnated samples, *Geochemistry*, *Geophysics, Geosystems*, 20, 5650-5666.
- Biedermann, A.R., 2020. FinIrrSDA: A 3D model for magnetic shape and distribution anisotropy of
 finite irregular arrangements of particles with different sizes, geometries, and orientations,
 Journal of Geophysical Research: Solid Earth, e2020JB020300.
- Biedermann, A.R., Lowrie, W. & Hirt, A.M., 2013. A method for improving the measurement of low field magnetic susceptibility anisotropy in weak samples, *Journal of Applied Geophysics*, 88,
 122-130.
- Brown, W.F., 1959. Relaxational Behavior of Fine Magnetic Particles, *Journal of Applied Physics*, 30,
 \$130-\$132.
- Brown, W.F., 1963. Thermal Fluctuations of a Single-Domain Particle, *Physical Review*, 130, 16771686.
- Cañón-Tapia, E., 1996. Single-grain versus distribution anisotropy: a simple three-dimensional
 model, *Physics of the Earth and Planetary Interiors*, 94, 149-158.
- Cañón-Tapia, E., 2001. Factors affecting the relative importance of shape and distribution anisotropy
 in rocks: theory and experiments, *Tectonophysics*, 340, 117-131.
- Clark, D.A., 2014. Methods for determining remanent and total magnetisations of magnetic sources
 a review. *Exploration Geophysics*, 45, 271-304.
- 867 Clark, D.A. & Emerson, D.W., 1999. Self-Demagnetisation, *Preview*, 79, 22-25.
- Clennell, M.B., 1997. Tortuosity: a guide through the maze. In: Lovell, M.A., Harvey, P.K. (ed.)
 Developments in Petrophysics, *Geological Society of London Special Publication*, 122, 299 344
- Cnudde, V. & Boone, M.N., 2013. High-resolution X-ray computed tomography in geosciences: A
 review of the current technology and applications, *Earth-Science Reviews*, 123, 1-17.
- Coffey, W.T. & Kalmykov, Y.P., 2012. Thermal fluctuations of magnetic nanoparticles: Fifty years
 after Brown, *Journal of Applied Physics*, 112.
- Bearing, J.A., Dann, R.J.L., Hay, K., Lees, J.A., Loveland, P.J., Maher, B.A. & O'Grady, K., 1996.
 Frequency-dependent susceptibility measurements of environmental materials, *Geophysical Journal International*, 124, 228-240.
- B78 Dormann, J.L., 1981. Le phénomène de superparamagnétisme, *Revue de Physique Appliquée*, 16,
 B79 275-301.
- Esteban, L., Géraud, Y. & Bouchez, J.L., 2006. Pore network geometry in low permeability argillites
 from magnetic fabric data and oriented mercury injections, *Geophysical Research Letters*,
 33, L18311.
- Eyre, J.K., 1997. Frequency dependence of magnetic susceptibility for populations of single-domain
 grains, *Geophysical Journal International*, 129, 209-211.
- Ghanbarian, B., Hunt, A.G., Ewing, R.P., Sahimi, M., 2013. Tortuosity in porous media: a critical
 review. Soil Science Society of America Journal, 77, 1461-1477.
- Goya, G.F., Berquó, T.S., Fonseca, F.C. & Morales, M.P., 2003. Static and dynamic magnetic
 properties of spherical magnetite nanoparticles, *Journal of Applied Physics*, 94, 3520-3528.

Grégoire, V., Darrozes, P., Gaillot, P. & Nédélec, A., 1998. Magnetite grain shape fabric and distribution anisotropy vs rock magnetic fabric: a three-dimensional case study, *Journal of Structural Geology*, 20, 937-944.

- Grégoire, V., De Saint-Blanquat, M., Nédélec, A. & Bouchez, J.-L., 1995. Shape anisotropy versus
 magnetic interactions of magnetite grains: experiments and application to AMS in granitic
 rocks, *Geophysical Research Letters*, 22, 2765-2768.
- Hailwood, E.A., Bowen, D., Ding, F., Corbett, P.W.M. & Whattler, P., 1999. Characterizing pore fabrics
 in sediments by anisotropy of magnetic susceptibility analyses. *in Paleomagnetism and Diagenesis in Sediments*, pp. 125-126, eds. Tarling, D. H. & Turner, P. Geological Society,
 London, Special Publications.
- Hailwood, E. & Ding, F., 2000. Sediment transport and dispersal pathways in the Lower Cretaceous
 sands of the Britannia Field, derived from magnetic anisotropy, *Petroleum Geoscience*, 6,
 369-379.
- Hargraves, R.B., Johnson, D. & Chan, C.Y., 1991. Distribution anisotropy: the cause of AMS in igneous
 rocks?, *Geophysical Research Letters*, 18, 2193-2196.
- Hext, G.R., 1963. The estimation of second-order tensors, with related tests and designs, *Biometrika*,
 50, 353-373.
- Hrouda, F., 2011. Models of frequency-dependent susceptibility of rocks and soils revisited and
 broadened, *Geophysical Journal International*, 187, 1259-1269.
- Hrouda, F., Chadima, M., Jezek, J. & Pokorny, J., 2017. Anisotropy of out-of-phase magnetic
 susceptibility of rocks as a tool for direct determination of magnetic subfabrics of some
 minerals: an introductory study, *Geophysical Journal International*, 208, 385-402.
- Hrouda, F., Hanak, J. & Terzijski, I., 2000. The magnetic and pore fabrics of extruded and pressed
 ceramic models, *Geophysical Journal International*, 142, 941-947.
- Hrouda, F., Pokorný, J. & Chadima, M., 2015. Limits of out-of-phase susceptibility in magnetic
 granulometry of rocks and soils, *Studia Geophysica et Geodaetica*, 59, 294-308.
- Hrouda, F., Pokorný, J., Ježek, J. & Chadima, M., 2013. Out-of-phase magnetic susceptibility of rocks
 and soils: a rapid tool for magnetic granulometry, *Geophysical Journal International*, 194,
 170-181.
- Huang, T., Tao, Z., Li, E., Lyu, Q. & Guo, X., 2017. Effect of permeability anisotropy on the production
 of multi-scale shale gas reservoirs, *Energies*, 10.
- Humbert, F., Robion, P., Louis, L., Bartier, D., Ledésert, B. & Song, S.-R., 2012. Magnetic inference of
 in situ open microcracks in sandstone samples from the Taiwan Chelungpu Fault Drilling
 Project (TCDP), *Journal of Asian Earth Sciences*, 45, 179-189.
- ljeje, J.J., Gan, Q. & Cai, J., 2019. Influence of permeability anisotropy on heat transfer and
 permeability evolution in geothermal reservoir, *Advances in Geo-Energy Research*, 3, 43-51.
- Jackson, M., 2003-2004. Imaginary Susceptibility A Primer, *The IRM Quarterly*, 13, 1, 10-11.
- Jackson, M., Moskowitz, B., Rosenbaum, J. & Kissel, C., 1998. Field-dependence of AC susceptibility
 in titanomagnetites, *Earth and Planetary Science Letters*, 157, 129-139.
- Jelinek, V., 1977. The statistical theory of measuring anisotropy of magnetic susceptibility of rocks
 and its application.
- 930 Jelinek, V., 1981. Characterization of the magnetic fabric of rocks, *Tectonophysics*, 79, T63-T67.
- Jezek, J. & Hrouda, F., 2007. A program for magnetic susceptibility-equivalent pore conversion,
 Geochemistry Geophysics Geosystems, 8, GC001709.
- Jones, D.H. & Srivastava, K.K.P., 1989. A re-examination of models of superparamagnetic relaxation,
 Journal of Magnetism and Magnetic Materials, 78, 320-328.
- Jones, S., Benson, P. & Meredith, P., 2006. Pore fabric anisotropy: testing the equivalent pore
 concept using magnetic measurements on synthetic voids of known geometry, *Geophysical Journal International*, 166, 485-492.
- Joseph, A. & Mathew, S., 2014. Ferrofluids: synthetic strategies, stabilization, physicochemical
 features, characterization, and applications, *ChemPlusChem*, 79, 1382-1420.
- Joseph, R.I., 1966. Ballistic demagnetizing factor in uniformly magnetized cylinders, *Journal of Applied Physics*, 37, 4639-4643.

- Joseph, R.I., 1967. Ballistic demagnetizing factor in uniformly magnetized rectangular prisms, *Journal of Applied Physics*, 38, 2405-2406.
- Joseph, R.I., 1976. Demagnetizing factors in nonellipsoidal samples a review, *Geophysics*, 41, 1052 1054.
- Joseph, R.I. & Schlömann, E., 1965. Demagnetizing field in nonellipsoidal bodies, *Journal of Applied Physics*, 36, 1579-1593.
- Kosterov, A.A., Sergienko, E.S., Kharitonskii, P.V. & Yanson, S.Y., 2018. Low temperature magnetic
 properties of basalts containing near ~TM30 titanomagnetite, *Izvestiya, Physics of the Solid Earth*, 54, 134-149.
- Landis, E.N. & Keane, D.T., 2010. X-ray microtomography, *Materials Characterization*, 61, 1305-1316.
- Louis, L., David, C., Metz, V., Robion, P., Menendez, B. & Kissel, C., 2005. Microstructural control on
 the anisotropy of elastic and transport properties in undeformed sandstones, *International Journal of Rock Mechanics and Mining Sciences*, 42, 911-923.
- Muscas, G., Concas, G., Cannas, C., Musinu, A., Ardu, A., Orrù, F., Fiorani, D., Laureti, S., Rinaldi, D.,
 Piccaluga, G. & Peddis, D., 2013. Magnetic Properties of Small Magnetite Nanocrystals, *The Journal of Physical Chemistry C*, 117, 23378-23384.
- Nabawy, B.S., Rochette, P. & Géraud, Y., 2009. Petrophysical and magnetic pore network anisotropy
 of some cretaceous sandstone from Tushka Basin, Egypt, *Geophysical Journal International*,
 177, 43-61.
- 961 Néel, L., 1949. Influence des fluctuations thermiques sur l'aimantation de grains ferromagnétiques
 962 très fins, *Comptes rendus hebdomadaires des séances de l'Académie des sciences* T228, 664 963 666.
- 964 Odenbach, S., 2004. Recent progress in magnetic fluid research, *Journal of Physics: Condensed* 965 *Matter*, 16, R1135-R1150.
- 966 Osborn, J.A., 1945. Demagnetizing factors of the general ellipsoid, *Physical Review*, 67, 351-357.
- Panja, P., McLennan, J. & Green, S., 2021. Influence of permeability anisotropy and layering on
 geothermal battery energy storage, *Geothermics*, 90.
- Papaefthymiou, G.C., 2009. Nanoparticle magnetism, *Nano Today*, 4, 438-447.
- Parés, J., Miguens, L. & Saiz, C., 2016. Characterizing pore fabric in sandstones with magnetic
 anisotropy methods: initial results, *Journal of Petroleum Science and Engineering*, 143, 113120.
- Pfleiderer, S. & Halls, H.C., 1990. Magnetic susceptibility anisotropy of rocks saturated with
 ferrofluid: a new method to study pore fabric?, *Physics of the Earth and Planetary Interiors*,
 65, 158-164.
- Pfleiderer, S. & Halls, H.C., 1993. Magnetic pore fabric analysis: Verification through image
 autocorrelation, *Journal of Geophysical Research*, 98, 4311-4316.
- Pfleiderer, S. & Halls, H.C., 1994. Magnetic pore fabric analysis: a rapid method for estimating
 permeability anisotropy, *Geophysical Journal International*, 116, 39-45.
- Pfleiderer, S. & Kissel, C., 1994. Variation of pore fabric across a fold-thrust structure, *Geophysical Research Letters*, 21, 2147-2150.
- Pugnetti, M., Zhou, Y., Biedermann, A.R., 2021 Experimental improvements for ferrofluid
 impregnation of rocks using directional forced impregnation methods: results on natural and
 synthetic samples. EGU virtual general assembly.
- Robion, P., David, C., Dautriat, J., Colombier, J.-C., Zinsmeister, L. & Collin, P.-Y., 2014. Pore fabric
 geometry inferred from magnetic and acoustic anisotropies in rocks with various
 mineralogy, permeability and porosity, *Tectonophysics*, 629, 109-122.
- 988 Rosensweig, R.E., 1987. Magnetic fluids, *Annual Review of Fluid Mechanics*, 19, 437-463.
- Rosensweig, R.E., 1988. An introduction to ferrohydrodynamics, *Chemical Engineering Communications*, 67, 1-18.
- Sato, M. & Ishii, Y., 1989. Simple and approximate expressions of demagnetizing factors of uniformly
 magnetized rectangular rod and cylinder, *Journal of Applied Physics*, 66, 983-985.

- Sinan, S., Glover, P.W.J. & Lorinczi, P., 2020. Modelling the impact of anisotropy on hydrocarbon
 production in heterogeneous reservoirs, *Transport in Porous Media*, 133, 413-436.
 Söffge, F. & Schmidbauer, E., 1981. AC susceptibility and static magnetic properties of an Fe₃O₄
 ferrofluid, *Journal of Magnetism and Magnetic Materials*, 24, 54-66.
 Stephenson, A., 1971. Single domain grain distributions: 1. A method for the determination of single
 domain grain distributions, *Physics of the Earth and Planetary Interiors*, 4, 353-360.
- 999Stephenson, A., 1994. Distribution anisotropy: two simple models for magnetic lineation and1000foliation, Physics of the Earth and Planetary Interiors, 82, 49-53.
- 1001Stoner, E.C., 1945. The demagnetizing factors for ellipsoids, The London, Edinburgh, and Dublin1002Philosophical Magazine and Journal of Science: Series 7, 36, 803-821.
- Storesletten, L., 1998. Effects of anisotropy on convective flow through porous media. *in Transport Penomena in Porous Media*, pp. 261-283, eds. Ingham, D. B. & Pop, I. Pergamon (Elsevier),
 Oxford, UK.
- Syono, Y., 1965. Magnetocrystalline anisotropy and magnetostriction of Fe₃O₄–Fe₂TiO₄ series, with
 special application to rock magnetism, *Japanese Journal of Geophysics*, 4, 71-143.
- 1008Torres-Diaz, I. & Rinaldi, C., 2014. Recent progress in ferrofluids research: novel applications of1009magnetically controllable and tunable fluids, *Soft Matter*, 10, 8584-8602.
- Wang, C., Huang, Z., Lu, Y., Tang, G. & Li, H., 2019. Influences of reservoir heterogeneity and
 anisotropy on CO₂ sequestration and heat extraction for CO₂-based enhanced geothermal
 system, *Journal of Thermal Science*, 28, 319-325.
- Wang, G., Wei, X., An, H., Wang, F.-Y. & Rudolph, V., 2014. Modeling anisotropic permeability of coal
 and its effects on coalbed methane reservoir simulation. *in Proceedings of the 4th International Conference on Simulation and Modeling Methodologies, Technologies and Applications*, pp. 477-483.
- Willems, C.J.L., Nick, H.M., Donselaar, M.E., Weltje, G.J. & Bruhn, D.F., 2017. On the connectivity
 anisotropy in fluvial Hot Sedimentary Aquifers and its influence on geothermal doublet
 performance, *Geothermics*, 65, 222-233.
- 1020Worm, H.-U., 1998. On the superparamagnetic stable single domain transition for magnetite, and1021frequency dependence of susceptibility, *Geophysical Journal International*, 133, 201-206.
- Worm, H.-U. & Jackson, M., 1999. The superparamagnetism of Yucca Mountain Tuff, *Journal of Geophysical Research: Solid Earth*, 104, 25415-25425.
- Zhou, Y., Pugnetti, M., Foubert, A., Lanari, P., Neururer, C., Biedermann, A.R., 2021. Correlations of
 magnetic pore fabrics with pore fabrics derived from high-resolution X-ray computed
 tomography and with permeability anisotropy in sedimentary rocks and synthetic samples.
 EGU virtual general assembly.