Spatial heterogeneity and lake morphology affect diffusive greenhouse gas emission estimates of lakes

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Most estimates of diffusive flux (F) of methane (CH4) and carbon dioxide (CO2) from lakes are based on single-point flux chamber measurements or on piston velocity (k) modeled from wind speed and single-point measurements of surface water gas concentrations (Caq). We analyzed spatial variability of F of CH4 and CO2, as well as Caq and k in 22 European lakes during late summer. F and k were higher in the lake centers, leading to considerable bias when extrapolating single-point chamber measurements to whole-lake estimates. The ratio of our empirical k estimates to wind speed-modeled k was related to lake size and shape, suggesting a lake morphology effect on the relationship between wind speed and k. This indicates that the error inherent to established wind speed models can be reduced by determining k and Caq at multiple sites on lakes to calibrate wind speed-modeled k to the local system. Citation: Schilder, J., D. Bastviken, M. van Hardenbroek, P. Kankaala, P. Rinta, T. Stötter, and O. Heiri (2013), Spatial heterogeneity and lake morphology affect diffusive greenhouse gas emission estimates of lakes, Geophys. Res. Lett., 40, 5752–5756, doi:10.1002/2013GL057669.

1. Introduction

Large amounts of methane (CH4) and carbon dioxide (CO2), two important greenhouse gases, are released to the atmosphere from inland waters [Bastviken et al., 2004; Cole et al., 2007]. Estimates show that these fluxes correspond to a large share of the terrestrial carbon sink [Bastviken et al., 2011; Tranvik et al., 2009]. Therefore, accurate assessments of CO2 and CH4 fluxes from inland waters are vital for reliable estimates of the terrestrial greenhouse gas balance [Battin et al., 2009]. From open water (i.e., the lake area free of emerging vegetation), lakes emit CO2 and CH4 in several ways, including ebullition (bubbling of nondissolved gases from the sediments) and by diffusive exchange (flux of dissolved gases across the air-water interface). Diffusive flux (F) dominates for highly soluble gases such as CO2, and F may account for up to 50% of the total CH4 flux [Bastviken et al., 2004]. Most estimates of diffusive flux of CO2 (FCO2) and CH4 (FC14) from aquatic ecosystems rely on measurements of surface water concentrations (Caq) and the following equation:

\[ F = k(C_{aq} - C_{eq}) \]  

where F is the diffusive flux (mmol m⁻² day⁻¹), k is the gas exchange coefficient or piston velocity (m d⁻¹ in equation (1); frequently and here expressed in cm h⁻¹), and Ceq is the theoretical surface water concentration (µM) when in equilibrium with air partial pressure (typically calculated from Henry’s Law) [Cole and Caraco, 1998]. Ceq is generally measured at one single location in a lake, usually the center [e.g., Bastviken et al., 2004; Sobek et al., 2005; Juutinen et al., 2009; Marotta et al., 2009]. In turn, k is frequently estimated from wind speed at 10 m height (U10) based on empirical relationships. Unfortunately, empirical relationships between k and U10 are only available from a few systems [Bade, 2009; Wanninkhof et al., 2009; Vachon et al., 2010], and it is unclear to what extent the general use of these models is valid. Gas accumulation measurements with floating chambers can also provide estimates of F in lakes [Cole et al., 2010]. However, such estimates are again often based on measurements in a single location within the lake.

Available equations for calculating k from U10 [e.g., Crusius and Wanninkhof, 2003; Cole and Caraco, 1998] lead to differing estimates of k at a given U10. For many values of U10 in the range 0–10 m s⁻¹, one model returns a k value which is twice as high as the other. Further, a comparison of the model by Cole and Caraco [1998] with the underlying data set shows that for different sites, k at a given U10 can differ twofold from the model prediction. Hence, the error inherent to wind speed models, combined with variation caused by the choice of model, results in an uncertainty in whole-lake estimates of F that can be substantial.

Only few studies provide information on how FC14 and FCO2 vary within lakes [e.g., Hofmann, 2013], and we know of none which specifically addresses the within-lake variability of k. Such studies are needed for evaluating the reliability of F estimated from single-spot measurements. Furthermore, they might reveal factors responsible for the variability of k at a given U10 as documented by SF6 tracer experiments [e.g., Cole and Caraco, 1998]. We measured within-lake spatial patterns in Caq of CH4 and of accumulation of CH4 and CO2 in floating chambers for 32 lakes in Europe. At each site, chambers were deployed at four locations across the surface of the lake. Based on these data, we estimated FC14 for chambers which were not affected by ebullition, resulting in a data set of 12 lakes with reliable FC14 estimates from four zones within the lake and a second data set of 22 lakes with at least one estimate in the zones nearest to shore and one in the more central zones. This in turn allowed an assessment of the
variation in $k$ within and between lakes. We estimated the bias in whole-lake $F_{\text{CH}_4}$ and $F_{\text{CO}_2}$ when scaling up from single-point estimates. Finally, we discuss mechanisms that can explain the within- and between-lake patterns of $k$ and the discrepancies between $k$ estimates from wind speed-based models and our observations.

2. Methods

2.1. Sample Collection and Gas Measurements

In August and September 2010 and 2011 we measured $\text{CH}_4$ and $\text{CO}_2$ accumulation into floating chambers, $C_{\text{aq}}$ of $\text{CH}_4$, and surface water temperatures at multiple locations on 32 small- to intermediate-sized lakes in Europe. At each lake, four groups of three replicate chambers were deployed for 6 h ($\sim 10:00$–16:00) on a transect from the nearshore zone to the central zone of each lake (zones A, B, C, and D), with increasing distance between chamber zones (see Figure S2 in the supporting information). The chamber group in zone A was always close to the shore line, just beyond emerging macrophytes (if any). Chambers representing zone D were always in the center of the lake. The group representing zone C was placed approximately halfway between A and D and the group in zone B halfway between A and C. Chambers that had captured ebullition as well as diffusive emissions were identified using our replicate flux measurements in each zone and the method described by Bastviken et al. [2004] (see supporting information). Chambers identified as having received diffusive emissions only were used to provide estimates of $F_{\text{CH}_4}$, $F_{\text{CH}_4}$, $C_{\text{aq}}$ of $\text{CH}_4$, and surface water temperatures were then used to calculate $k$ according to equation (1).

We adopted the floating chamber design and deployment methodology described by Bastviken et al. [2004] and Cole et al. [2010], which were shown to correspond well with multiple independent methods to determine $k$, implying negligible bias in chamber measurements [Cole et al., 2010; Gallfolk et al., 2013]. Samples for determining $C_{\text{aq}}$ were collected at each group of chambers following Bastviken et al. [2010]: 40 mL of water was sampled 5 cm below the surface with 60 mL syringes (Becton-Dickinson) and exposed to 20 mL of ambient air in the syringes by shaking for 60 s. The headspace, now equilibrated with the water in terms of gases, was injected into a glass vial (10 mm thick butyl rubber stopper; Apodan) prefilled with saturated brine solution. These samples were stored in vials until analysis. Ambient air was collected to correct for the background air concentration in the water sample extractions and in the chamber measurements. After approximately 6 h, 30 mL gas was taken from each chamber with a syringe. These samples were again transferred to and stored in glass vials. Within 60 days of sampling, $\text{CH}_4$ and $\text{CO}_2$ concentrations were analyzed in the laboratory by gas chromatography using a flame ionization detector with a methanizer (GC-FID; Shimadzu GC-8, PoropackN column). Tests initiated in 2006 confirmed that samples collected and stored as described here are stable for years. Samples from the Finnish lakes were measured on a gas chromatograph with an autosampler (Agilent 6890 N, PlotQ capillary column, with FID for $\text{CH}_4$ and Thermal Conductivity Detector for $\text{CO}_2$).

2.2. Data Analysis and Upscaling

For chambers not affected by ebullition, $F_{\text{CH}_4}$ can be determined from $\text{CH}_4$ accumulation in the chambers, accounting for chamber area, volume, and deployment duration and correcting for the decreasing concentration gradient due to increasing $\text{CH}_4$ concentration in the chambers [Cole et al., 2010]. The chamber deployment, sample handling, and storage were optimized for $\text{CH}_4$. Deployment was too long for providing estimates of $F_{\text{CO}_2}$ because $\text{CO}_2$ equilibrates much faster with the chamber headspace (typically within a few hours). Therefore, after 6 h concentrations of $\text{CO}_2$ in the chamber headspace will have reached values close to equilibrium with $p\text{CO}_2$ in the surface water. This also implies that final $\text{CO}_2$ concentrations in the chambers can be used to estimate $C_{\text{aq}}$ of $\text{CO}_2$ following Henry’s Law, which allowed us to estimate $F_{\text{CO}_2}$ (using $k$ of $\text{CH}_4$ transformed to $k$ of $\text{CO}_2$; Bade [2009]). We report $k$ as $k_{5000}$, the value corresponding to a gas with a Schmidt number of 600 ($\text{CO}_2$ at 20°C).

For analysis of within-lake patterns of $k$, $F$, and $C_{\text{aq}}$, and for scaling up, we partitioned the original data set into different subsets (Tables S1 and S2): A data set of 13 lakes that consists of lakes that yielded at least one estimate of $k$ for each sampling zone (subset 1). This subset was used to identify spatial patterns of $k$, $F_{\text{CH}_4}$, and $F_{\text{CO}_2}$ in the study lakes. The second subset consists of 24 lakes with at least one estimate of $k$ each in the central part of the lake (zones C + D) and the nearshore area (zones A + B) (subset 2). For these lakes, surface areas were digitized from national topographic maps. Using ArcMap 9.3 (Esri), the area belonging to the nearshore and the central zones was quantified. Nearshore area was defined as lying within a distance to shore corresponding to the mean distance of chamber groups B and C to shoreline. The central area consisted of the remaining surface. Estimates of $k$ and $F$ for nearshore and central areas from subset 2 were then used to scale up to whole-lake averages weighted by area. $U_{10}$ was provided by the respective national meteorological services.

One lake that fulfilled the selection criteria for subset 1 and two lakes that fulfilled the criteria for subset 2 were not included in further analysis. At Valkea-Kotinen, $\text{CH}_4$ accumulation rates were so low they resulted in $k$ values lower than commonly found in literature. Therefore, this lake was eliminated from both subsets. The wind speed data for lake Kaisajön (3.9 m s$^{-1}$) deviated strongly from our observations in the field (the lake was sampled on the windiest day of the campaign). Therefore, we concluded that the wind speed data did not reflect the local conditions, and this lake was eliminated from subset 2. Measurements from these lakes can be found in the supporting information (Tables S1 and S2) with those of the other lakes not incorporated in either subset (Table S3).

3. Spatial Patterns and Whole-Lake Estimates of $k$

Values of $k_{5000}$ were clearly higher in the central zone of the lake in 11 of the 12 lakes remaining in subset 1 (Figure 1). $C_{\text{aq}}$ of $\text{CH}_4$ was usually lower in the center, in agreement with recent findings by Hofmann [2013]. $F_{\text{CH}_4}$ showed the opposite pattern, however, with highest values in the central zones, which is in contrast with the conclusions by Hofmann [2013] that are based on wind speed-derived $k$ values. $C_{\text{aq}}$ of $\text{CO}_2$ and $F_{\text{CO}_2}$ were lowest nearshore and more elevated in the central zones (Figure 1).
were usually higher than average, again with a large overlap in range. Grouping chamber measurements into two groups instead of four allowed spatially resolved estimates of \( k_{600} \) and \( F \) to be calculated for 22 of our 32 study lakes. Whole-lake estimates of \( k_{600} \) based on this 22-lake data set are in the range of 1.2–6.3 cm h\(^{-1}\) \( /C_0 \). The relationship between \( U_{10} \) and our estimates of \( k_{600} \) is in agreement with the model by Cole and Caraco [1998] (Figure 2). Residuals of \( k_{600} \) to modeled values for our data set (range 1.9 to 1.7 cm h\(^{-1}\) \( /C_0 \), standard deviation (SD) 1.1 cm h\(^{-1}\) \( /C_0 \)) are similar to the residuals recalculated for the SF6 tracer studies that were originally used to develop this model (range 2.8 to 1.7 cm h\(^{-1}\) \( /C_0 \), SD 1.3 cm h\(^{-1}\) \( /C_0 \)) [Cole and Caraco, 1998, Figure 8a]. We therefore conclude that the chamber method we used for assessing \( k \) provides comparable results to tracer studies.

4. Implications for Whole-Lake Estimates of \( F \)

[12] We compared \( F \) derived from single-point chamber measurements to whole-lake estimates based on the spatially resolved 22-lake data set (the values are provided in Table S2). Values based on single locations ranged from 38 to 222% of whole-lake \( F_{CH_4} \). On average, inferences from near-shore measurements slightly underestimated \( F_{CH_4} \) (88 %), whereas those from the center provided overestimates (115%). Measuring \( C_{aq} \) in the lake center and then applying the model by Cole and Caraco [1998] yields \( F_{CH_4} \) amounting to 55–300% of our whole-lake estimates (average 110%) and 33–320% for \( F_{CO_2} \) (average 148%). Hence, this common way of estimating fluxes results in \( F_{CH_4} \) and \( F_{CO_2} \) being overestimated on average by 10% and 48%, respectively, in our data set. For individual lakes, the bias is highly variable, reflecting the large variability in the data behind the \( k \) versus \( U_{10} \) relationship.

5. Relationship of \( k \) With Lake Morphometry

[13] The relationship between whole-lake \( k_{600} \) and \( U_{10} \) is strong and well studied [e.g., Cole and Caraco, 1998; Crusius and Wanninkhof, 2003; Wanninkhof et al., 2009; Vachon et al., 2010]. Available wind speed-based models provide realistic estimates of the average whole-lake \( k_{600} \) expected at a given wind speed. However, estimates for individual lakes can still deviate considerably from this value (Figure 2) [Cole and Caraco, 1998]. We used the ratio between \( k_{600} \) modeled following Cole and Caraco [1998] (\( kCC \)) and \( k_{600} \) inferred from our chamber measurements (\( kS \)) to explore where the two methods lead to the most pronounced differences in \( k_{600} \). This ratio (\( kCC/kS \)) was highest for lakes in which \( k_{600} \) was higher in the center than near-shore (Figure 3a), suggesting that the spatial gradient in \( k_{600} \) may be a source of bias in wind speed-based estimates of \( k_{600} \). We consider the most likely explanation for the observed within-lake variability of \( k \) to be the proximity to shoreline and sheltering shoreline structures and vegetation. This potentially reduces the direct impact of wind at the lake.
surface and can explain low $k_{600}$ and $F$ recorded in nearshore zones (Figure 1). If this hypothesis is true, we would expect this sheltering to lead to a less pronounced relationship of $U_{10}$ with $k_{600}$ in nearshore areas than in the lake center, which is confirmed by our data (Figure 3b). Also, the proportion of the lake area influenced by sheltering effects of the shoreline is expected to be larger on small lakes. Therefore, we also expect higher $k_{600}$ in larger lakes. In our data, this is indeed the case (Figure 3c). Unfortunately, lake size is also correlated with $U_{10}$ in our data ($r = 0.49$, $p < 0.05$), so we cannot separate the relationship between $k_{600}$ and area from its relationship with $U_{10}$. However, the relationship of $k_{600}$ with lake area (Figure 3d) is absent or weak in nearshore areas ($r = 0.19$, $p = 0.39$), whereas it is strong in central zones ($r = 0.68$, $p < 0.001$). This again supports our interpretation that under similar wind conditions, nearshore areas are more protected.

If distance to shoreline influences $k_{600}$, both lake size and shape have the potential to affect the relationship of $k_{600}$ with $U_{10}$. However, both of these variables are presently rarely taken into account when producing wind speed-based estimates of $k_{600}$. That lake area can influence modeled $k_{600}$ is supported by its significant relationship with $kCC/ks$ in our data set (Figure 3e). This indicates that $k_{600}$ modeled following Cole and Caraco [1998] tends to provide overestimated values for small lakes. Similarly, a first analysis of our data in respect to the potential effects of lake shape reveals that more complex lake basins, which in our campaign were typically elongated in shape, were characterized by higher within-lake variability in $k_{600}$ (Figure 3f). Lakes with high variability in $k_{600}$ were also the ones in which the approach of Cole and Caraco [1998] returns higher values of $k_{600}$ than we observe (Figure 3a). Hence, lake shape also appears to influence whole-lake $k_{600}$.

Our interpretations suggest that the effect of $U_{10}$ on $k_{600}$ should be greater on larger lakes and lakes with simple shapes, in which sheltering effects of shoreline are limited. Large lakes provide the potential for a larger fetch, and lakes with more complex shorelines have shorter average distance to sheltering shoreline vegetation. We therefore expect the influence of $U_{10}$ on $k_{600}$ to be reduced in small, complex, and very elongated lakes. This also implies that wind speed-based models may have to be adapted for these kinds of systems.

Wanninkhof [1992] suggested lake size as an explanation for the discrepancy between existing SF$_6$ tracer studies, and Guérin et al. [2007] identified a relationship between basin size and $k$ for rivers and estuaries. Read et al. [2012] also stressed the importance of lake size for the convection component of $k$. Our data were collected at daytime, which means that the convection driven by heat loss during nights is less likely to have contributed to the observed spatial patterns. Our data support Wanninkhof’s [1992] suggestion and imply that wind speed models are only applicable to lakes with a size and shape similar to the lakes the model was developed on.

6. Conclusions

Our study shows distinct and profound spatial patterns in $F$ of greenhouse gases from lakes as well as in the driving parameters, $k$ and $C_{aq}$. It also reveals lake size and shape as

Figure 3. (a) Ratio of $k_{600}$ inferred from Cole and Caraco [1998] and our whole-lake $k_{600}$ estimates ($kCC/ks$) plotted versus the strength of spatial variability in $k_{600}$ ($k_{600}$/nearshore) ($r = 0.48$, $p < 0.05$). (b) The relationship between $U_{10}$ and $k_{600}$ for nearshore chamber group A (open circles; $r = 0.41$, $p = 0.1$) and central chamber group D (solid circles; $r = 0.67$, $p < 0.005$). (c) Whole-lake $k_{600}$ plotted against area ($r = 0.59$, $p < 0.005$). (d) Relationship between area and $k_{600}$ determined from the nearshore chamber groups A and B (open circles; $r = 0.19$, $p = 0.39$) and from the central chamber groups C and D (solid circles; $r = 0.68$, $p < 0.001$). (e) The $kCC/ks$ plotted versus area ($r = 0.53$, $p < 0.05$). (f) Lake shape complexity, here presented as shore development ($L_d$), and its relationship to $k_{600}$ (nearshore ($r = 0.49$, $p < 0.05$). $L_d$ is defined as the circumference of a lake divided by the circumference of a circle with the same area [Håkanson, 2004].
Furthermore, the equilibration of CH4 between the water and atmosphere is inexpensive, swift, and uncomplicated. CH4 appears to be more suitable than CO2 for determining error range is unsatisfying.

Lake system characteristics, such as lake size, lake shape, presence and height of vegetation, and convection into account vary from lake to lake. However, k values for individual lakes can be determined in situ, and the results can be used to calibrate established wind speed models to local conditions. Our data show that this can be done with floating chambers if properly designed [Cole et al., 2010] and with appropriate consideration of confounding factors like ebullition. The method is inexpensive, swift, and uncomplicated. CH4 appears to be more suitable than CO2 for determining k using chambers because CH4 is nearly always supersaturated and less affected by chemical processes or variables with diurnal variations such as primary production, respiration, and pH [Bastviken, 2009]. Furthermore, the equilibration of CH4 between the water and chamber headspace is often slow enough to allow chamber deployments over 24 h, yielding a daily average k accounting for different wind- and water-mixing conditions during day and night [Bastviken et al., 2010].

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References
SCHILDER ET AL.: SPATIAL HETEROGENEITY OF DIFFUSIVE FLUX