Present state of global wetland extent and wetland methane modelling: methodology of a model inter-comparison project (WETCHIMP)


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Abstract. The Wetland and Wetland CH\textsubscript{4} Intercomparison of Models Project (WETCHIMP) was created to evaluate our present ability to simulate large-scale wetland characteristics and corresponding methane (CH\textsubscript{4}) emissions. A multi-model comparison is essential to evaluate the key uncertainties in the mechanisms and parameters leading to methane emissions. Ten modelling groups joined WETCHIMP to run eight global and two regional models with a common experimental protocol using the same climate and atmospheric carbon dioxide (CO\textsubscript{2}) forcing datasets. We reported the main conclusions from the intercomparison effort in a companion paper (Melton et al., 2013). Here we provide technical details for the six experiments, which included an equilibrium, a transient, and an optimized run plus three sensitivity experiments (temperature, precipitation, and atmospheric CO\textsubscript{2} concentration). The diversity of approaches used by the models is summarized through a series of conceptual figures, and is used to evaluate the wide range of wetland extent and CH\textsubscript{4} fluxes predicted by the models in the equilibrium run. We discuss relationships among the various approaches and patterns in consistencies of these model predictions. Within this group of models, there are three broad classes of methods used to estimate wetland extent: prescribed based on wetland distribution maps, prognostic relationships between hydrological states based on satellite observations, and explicit hydrological mass balances. A larger variety of approaches was used to estimate the net CH\textsubscript{4} fluxes from wetland systems. Even though modelling of wetland extent and CH\textsubscript{4} emissions has progressed significantly over recent decades, large uncertainties still exist when estimating CH\textsubscript{4} emissions: there is little consensus on model structure or complexity due to knowledge gaps, different aims of the models, and the range of temporal and spatial resolutions of the models.

1 Introduction

In order to study the importance of wetlands in the global water and carbon cycle a variety of hydrological and biogeochemical models have been developed over the last three decades. The first studies of global-scale wetland CH\textsubscript{4} modelling appeared twenty-five years ago (Matthews and Fung, 1987). Matthews and Fung (1987) combined vegetation, soil and fractional inundation maps along with estimates of CH\textsubscript{4} flux intensity to generate a map of global wetland distribution and an annual wetland emissions estimate of $\sim$100 Tg CH\textsubscript{4} yr\textsuperscript{-1}. Aselman and Crutzen (1989) soon followed developing their own wetland distribution datasets, and assumed CH\textsubscript{4} emission flux rates, yielding a wetland emissions estimated range of 40–160 Tg CH\textsubscript{4} yr\textsuperscript{-1}. These early approaches are limited by uncertainties inherent in upscaling point measurements to large regions, and an inability to predict changes to wetland systems due to changes in climate and hydrology because of the use of static wetland extent and simple scaling-based estimates of CH\textsubscript{4} emissions.

As an attempt to circumvent these limitations, process-based modelling of global CH\textsubscript{4} emissions from wetland systems was first pioneered by Fung et al. (1991) followed by Christensen and Cox (1995), Christensen et al. (1996), and Cao et al. (1996). While those early global studies used the static wetland maps of Matthews and Fung (1987), they differed in their approach to simulate the CH\textsubscript{4} emissions. Christensen and Cox (1995) was the first study to introduce a formulation for oxidation and a soil vertical discretization in a one-dimensional, single-column model. The simple approach of Christensen et al. (1996) estimates net CH\textsubscript{4} emissions as a fraction of heterotrophic respiration calculated by an equilibrium vegetation model (BIOME2) giving a climate sensitive, but perhaps simplistic CH\textsubscript{4} emissions estimate. A more mechanistic approach was adopted by Cao et al. (1996) whose CH\textsubscript{4} emission model assumes substrate supply to methanogens is controlled by plant primary productivity and soil organic matter decomposition. Methane production is then modelled as a function of soil temperature, soil organic matter decomposition, water table position, and a fixed ratio of CH\textsubscript{4} production to decomposed organic carbon. Oxidation of produced CH\textsubscript{4} is assumed to scale with gross primary productivity (GPP) and to increase to a fixed fraction under non-inundated conditions (Cao et al., 1996).

None of these initial modelling studies performed transient simulations and the models originally accounted for hydrologic variability only in the vertical dimension, not spatially within model grid cells. Other process-based models soon followed but were not applied on a global scale, at least initially (Walter et al., 1996; Potter, 1997; Walter and Heimann, 2000). These initial papers included mechanistic modelling of such processes as diffusive, aerenchymal, and ebullition gas and oxygen transport. More recent work has devoted much effort to improving modelling of these processes (Segers and Leffelaar, 2001; van Bodegom et al., 2001a,b; Zhuang et al., 2006) and other controls on CH\textsubscript{4} production such as pH (Zhuang et al., 2004). Oxidation in theoxic portion of the soil, water column, and rhizosphere has also been parameterized (Ridgwell et al., 1999; Segers and Leffelaar, 2001; Zhuang et al., 2006; Curry, 2007, 2009). Model simulations have also moved on from equilibrium-only simulations to transient simulations (Walter et al., 2001a,b; Shindell et al., 2004; Gedney et al., 2004; Zhuang et al., 2006). Regional- to global-scale models have now been applied for the recent past (Ringleval et al., 2010; Hodson et al., 2011; Spahni et al., 2011; Riley et al., 2011), more distant past climates (Kaplan, 2002; Valdes et al., 2005; Hopcroft et al., 2011; Singrayer et al., 2011; Beerling et al., 2011), and to project responses to future climate change (Shindell et al., 2004; Gedney et al., 2004; Bohn et al., 2007; Bohn and Lettenmaier, 2010; Ringleval, 2011). Wetland and wetland CH\textsubscript{4} models are now becoming included in intermediate complexity (Shindell et al., 2004; Gedney et al., 2004; Avis et al., 2011) and comprehensive (Riley et al., 2011) global climate and earth system models.
A number of models have also integrated approaches allowing for dynamic wetland response to climate changes. Approaches to simulate wetland distribution in order to study the interaction between climate and free water bodies were developed by Coe (1997, 1998) and Krinner (2003). The earliest attempt at wetland modelling for the purpose of estimating wetland CH₄ emissions was designed to estimate wetland emissions during the Last Glacial Maximum (LGM, ~ 21 ka; Kaplan, 2002). The simple scheme of Kaplan (2002) used threshold values of slope and soil moisture content to define wetland areas, with the soil moisture calculated by an equilibrium vegetation model (BIOME4); an approach adopted by other models (Shindell et al., 2004; Weber et al., 2010; Avis et al., 2011). Later schemes used land cover datasets to outline peatland regions (Wania et al., 2009a, 2010; Spahni et al., 2011), and/or satellite-derived inundation datasets to prescribe wetlands either directly (Hodson et al., 2011; Ringeval et al., 2010), or indirectly (Ringeval et al., 2011; Riley et al., 2011). Other wetland distribution schemes use internally calculated water table positions or soil moisture thresholds to locate wetlands (Chen et al., 2012).

In this context, the Wetland and Wetland CH₄ Intercomparison of Models Project (WETCHIMP) was designed to offer the first multi-model comparison highlighting similarities and differences between modelling approaches and results. The advantage of using a multi-model comparison is that the range of the current state-of-the-art model estimates for wetland extent and CH₄ emissions can be studied in parallel. This approach allows us to study the sources of uncertainties and spatial and temporal differences in model behaviour. The results of this multi-model comparison are presented in Melton et al. (2013). In this paper, we provide the technical background for WETCHIMP, presenting details of the modelling protocol (Sect. 2), descriptions of the models as used for WETCHIMP (Sect. 3), conceptual comparisons of the models involved and results from the model default simulation to illustrate the differences between models are presented in Sect. 4. Following the definitions set out in Melton et al. (2013), we define wetlands for the purpose of large-scale modelling as grid cells, or fractions thereof, where the land surface has inundated, or saturated, conditions. Peatlands are a form of wetlands characterized by fixed extent, at least on timescales of decades, and contrasting hydrologic and nutrient regimes between dry nutrient-poor bogs and wet nutrient-rich fens (Melton et al., 2013). Inundated areas are assumed to be wetlands (unless masked out with a rice agriculture or lake dataset) with the water table above, or at the soil surface, but do not include areas that are unsaturated at the soil surface but saturated at depth.

## 2 Modelling protocol

The models participating in WETCHIMP followed a common modelling protocol outlined in Table 1 and adhered to it as closely as possible; divergences from the modelling protocol are described in the individual model description section.

<table>
<thead>
<tr>
<th>Name</th>
<th>Experiment</th>
<th>Climate years</th>
<th>Description</th>
</tr>
</thead>
<tbody>
<tr>
<td>Experiment 3-opt</td>
<td>Optimized</td>
<td>User-defined</td>
<td>Model run with user selected “optimal” configuration. Comparison phase is 1993–2004.</td>
</tr>
<tr>
<td>Experiment 4-CO₂</td>
<td>Atmospheric [CO₂] sensitivity</td>
<td>1901–1931</td>
<td>From the model state at end of Experiment 1-equil simulation, apply a globally uniform step increase in [CO₂] to reach SRES A2 2100 levels (857 ppmv). Run model until equilibrium is re-established.</td>
</tr>
<tr>
<td>Experiment 5-T</td>
<td>Temperature sensitivity</td>
<td>1901–1931</td>
<td>From the model state at end of equilibrium run, apply a step increase in air temperature reflecting mean SRES A2 2100 increase (multi-model mean SAT warming for 2090 to 2099 relative to 1980 to 1999: ≈ +3.4 °C). Run model until equilibrium is re-established.</td>
</tr>
<tr>
<td>Experiment 6-P</td>
<td>Moisture sensitivity</td>
<td>1901–1931</td>
<td>From the model state at end of transient equilibrium run, a step increase in precipitation to reflecting mean SRES A2 2100 increase (30 yr average 2071 to 2100 relative to 1961 to 1990: ≈ +3.9 %b). Run model until equilibrium is re-established.</td>
</tr>
</tbody>
</table>

* Each modelling group used their own criteria for what equilibrium meant: LPJ-WhyMe, LPJ Bern, SDGVM used the stability of the soil C pool; UVic-ESCM used soil moisture and temperature variables; DLEM specified an upper limit for inter-annual changes in total ecosystem C storage (< 0.1 g C m⁻²), soil moisture (< 0.1), and nitrogen storage (< 0.1 g N m⁻²). LPJ:WSL used soil and vegetation carbon. IAP-RAS is an equilibrium model and thus does not require spin-up. b As the IPCC AR4 report does not contain a globally averaged number for the mean precipitation change, this value is from the IPCC TAR Report of 2001.
Table 2. List of experiments (described in Table 1) performed by each of the participating modelling groups. “Prognostic” indicates that a model located at least part of its wetlands based either on an inversion to GIEMS and modelled hydrological state or used GIEMS as a spatial or temporal constraint. “Prescribed” means that the model used a distribution map for wetlands and “simulated” means that the model did not input any kind of wetland observational data to locate wetlands (see Sect. 4.1).

<table>
<thead>
<tr>
<th>Model</th>
<th>Areal extent</th>
<th>Wetlands</th>
<th>CH4 fluxes</th>
<th>Experiments performed</th>
<th>Contact</th>
</tr>
</thead>
<tbody>
<tr>
<td>CLM4Me</td>
<td>global</td>
<td>prognostic</td>
<td>simulated</td>
<td>1, 2, 3, 4, 5, 6</td>
<td>W. J. Riley</td>
</tr>
<tr>
<td>DLEM</td>
<td>global</td>
<td>prognostic</td>
<td>simulated</td>
<td>1, 2, 3, 4, 5, 6</td>
<td>H. Tian</td>
</tr>
<tr>
<td>IAP-RAS</td>
<td>global</td>
<td>prescribed</td>
<td>simulated</td>
<td>1, 2, 3, 4, 5, 6</td>
<td>A. V. Eliseev</td>
</tr>
<tr>
<td>LPJ-Bern</td>
<td>global</td>
<td>prognostic</td>
<td>simulated</td>
<td>1, 2, 3, 4, 5, 6</td>
<td>R. Spahni</td>
</tr>
<tr>
<td>LPJ-WhyMe</td>
<td>35–90° N</td>
<td>prescribed</td>
<td>simulated</td>
<td>1, 2, 3, 4, 5, 6</td>
<td>R. Wania</td>
</tr>
<tr>
<td>LPJ-WSL*</td>
<td>global</td>
<td>prescribed</td>
<td>simulated</td>
<td>1, 2, 3, 4, 5, 6</td>
<td>E. L. Hodson</td>
</tr>
<tr>
<td>ORCHIDEE</td>
<td>global</td>
<td>prognostic</td>
<td>simulated</td>
<td>1, 2, 3, 4, 5, 6</td>
<td>B. Ringleval</td>
</tr>
<tr>
<td>SDGVM</td>
<td>global</td>
<td>simulated</td>
<td>simulated</td>
<td>1, 2, 3, 4, 5, 6</td>
<td>P. O. Hopcroft</td>
</tr>
<tr>
<td>UVic-ESCM</td>
<td>global</td>
<td>simulated</td>
<td>n/a</td>
<td>1, 2, 3, 4, 5, 6</td>
<td>C. A. Avis</td>
</tr>
<tr>
<td>UW-VIC</td>
<td>W-Siberia</td>
<td>prognostic</td>
<td>simulated</td>
<td>1, 2, 3, 4, 5, 6</td>
<td>T. Bohn</td>
</tr>
</tbody>
</table>

* LPJ-WSL uses the “prognostic” approach for Experiment 3-opt, using GIEMS as guidance for the wetland distribution.

As briefly described in Melton et al. (2013), WETCHIMP consisted of six experiments, including both a transient simulation and several equilibrium state simulations with step-changes to climate forcing. The first experiment (“Experiment 1-equil”) was an equilibrium simulation under repeating 1901–1931 climate and a carbon dioxide concentration ([CO2]) of 303 ppmv. The second experiment (“Experiment 2-trans”) was a transient historical simulation from 1901–2009, using observed climate and atmospheric CO2 concentration ([CO2]), with the final (equilibrium) state of Experiment 1-equil as its initial state. Some models require observed fractional inundation values as an input; these were provided and cover the period 1993–2004 by the Global Inundation Extent from Multi-Satellites (GIEMS) dataset (Prigent et al., 2007; Papa et al., 2010). Thus, the period 1993–2004 was selected from Experiment 2-trans for comparison of model results. A third experiment (“Experiment 3-opt”) was run for the same time period as Experiment 2-trans, but allowed the models to run under user-defined optimal configurations (e.g. running coupled into an earth system model or using different meteorological forcing or remotely sensed inundation datasets than those common to Experiment 2-trans).

The remaining three experiments applied step-changes to each model’s equilibrium state from Experiment 1-equil. The fourth experiment (“Experiment 4-CO2”) applied an instantaneous increase in atmospheric [CO2] to 857 ppmv (SRES A2 year 2100 levels from IPCC, 2000) while holding the other meteorological inputs identical to Experiment 1-equil; this perturbed simulation was then run until each model had reached a new equilibrium state. Experiment 5 (“Experiment 5-T”) investigated the effect of an instantaneous increase of +3.4 °C in surface air temperature (SAT). The magnitude of this increase was chosen from the SRES A2 year 2100 multimodel mean SAT warming for the period 2080–2099 relative to 1980–1999 (Meehl et al., 2007). The final experiment (“Experiment 6-P”) examined model responses to changes in precipitation with an instantaneous increase of +3.9 % (SRES A2 2100 level; 30 yr global average for 2071–2100 relative to 1961–1990) (Prentice et al., 2001). In all cases, the step increases were applied to all months and grid cells uniformly. While actual changes in climate are projected to vary in both space and time, these uniform changes are suitable for the purpose of sensitivity tests (Melton et al., 2013). An overview of which groups conducted which simulations is shown in Table 2.

All data are freely available for download on http://arve.epfl.ch/pub/wetchimp, please send request for a username and password to joe.melton.sci@gmail.com.

2.1 Datasets

2.1.1 Climate data

The CRU (Climate Research Unit) TS3.1 time series (Mitchell and Jones, 2005; Jones and Harris, 2008) was used for monthly climate forcing data and – dependent on the model – precipitation, 2 m air temperature, percentage cloud cover, number of wet days, and vapour pressure were used from this dataset. Models that required data with a higher temporal resolution used the CRUNCEP (CRU and National Centre for Environmental Prediction) data, which is the correction of the 6 hourly NCEP reanalysis by the CRU TS3.1 data (Viovy and Ciais, 2011). CRUNCEP provides incoming long- and short-wave radiation, air specific humidity (used to compute the relative humidity), pressure, total precipitation, temperature, and the zonal and meridional components of the wind. UVic-ESCM (University of Victoria – Earth System Climate Model) used surface winds and diurnal temperature range from the NCEP reanalysis directly.
2.1.2 Soil and wetland distribution data

The soil data used in WETCHIMP are given in Table 3 and are allocated to each model in Table 4. The model requirements for soil data are too broad to accommodate a uniform soil dataset easily. Soil datasets used, and model treatment of soil textural information, are thus considered part of the wetland model itself.

There are several wetland distribution maps that were used for our simulations: (i) remotely sensed inundation area from GIEMS (Prigent et al., 2007; Papa et al., 2010) (Fig. 1), (ii) northern peatland distribution from NCSCD (Northern Circumpolar Soil Carbon Database; Tarnocai et al., 2007, 2009) (Fig. 2), and (iii) peatland distribution for the West Siberian Lowlands (Sheng et al., 2004) (Fig. 3). In addition, some groups made use of the rice distribution dataset by Leff et al. (2004) (Fig. 4) and the GICEW water bodies and land ice dataset (Fig. 5) to exclude areas from their wetland distribution map.

2.1.3 Global Inundation Extent from Multi-Satellites (GIEMS)

As the GIEMS dataset is used extensively by several models, and forms a comparison for the model outputs in Melton et al. (2013), it will be described in more detail here. The GIEMS dataset (Fig. 1) is a global, multi-year product quantifying the monthly variations of the distribution and extent of episodic and seasonal inundations, wetlands, rivers, lakes and irrigated agriculture at 0.25° resolution at the equator. GIEMS is derived from a complementary suite of satellite observations including passive microwave observations (SSMI (special sensor microwave/imager) emissivities), active microwave observations (ERS scatterometer), along with AVHRR-NDVI (Advanced Very High Resolution Radiometer-normalized difference vegetation index). The complete methodology is described in detail in Prigent et al. (2007) and Papa et al. (2010) and is briefly summarized here. First, an unsupervised classification of the three sources of satellite data is performed, and pixels with satellite signatures likely related to inundation are retained. For each inundated pixel, the monthly fractional coverage by open water is obtained using the passive microwave signal and a linear mixture model with end-members calibrated with scatterometer observations to account for the effects of vegetation cover. We use here the dataset available at a monthly timescale for the period 1993–2004. More detailed information concerning the seasonal and inter-annual behaviour of GIEMS dataset can be found in Prigent et al. (2012) for the global-scale analysis and in Papa et al. (2006) and Papa et al. (2008) for the tropical and boreal regions, respectively.

3 Participating models

In this section, we describe each model briefly and refer to published papers where more detailed information can be found. Modifications to models that were made after recent publications or specifically for WETCHIMP are described in the respective model section. An overview of which models performed which experiments is given in Table 2 and a list of forcing data for each model is provided in Table 4. The models follow the prescribed modelling protocol listed in Table 1 unless otherwise stated in the respective model description.
The changes in the code include: (i) the calculation of below-ground root mass for determining aerenchyma area now uses the time-lagged (1 yr decay time) belowground-to-aboveground NPP (net primary productivity) ratio, instead of the instantaneous one, and (ii) in calculating the water availability for permafrost vegetation, root fraction is weighted over all soil layers down to last year’s maximum active layer depth, rather than the instantaneous active layer depth, thereby causing a slight delay in growth in the spring.

### 3.1 CLM4Me

The version of CLM4Me used for this project is described in Riley et al. (2011), and is incorporated into the Community Land Model 4 (CLM4; Lawrence et al., 2011), the land-surface component of CESM1 (Community Earth System Model 1). Using the hydrology, soil carbon cycling, and soil thermal physics predicted in CLM4, net CH$_4$ fluxes are computed separately in inundated and non-inundated areas in each grid cell, including uptake of atmospheric CH$_4$. The reaction and transport equations for CH$_4$ and oxygen (where applicable) include production, consumption, aerenchyma transport, ebullition, and diffusion.

#### 3.1.1 WETCHIMP set-up

The CH$_4$ model code deviates slightly from that described in Riley et al. (2011); these changes resulted in less than a 5% difference from the global budget presented in Riley et al. (2011). The changes in the code include: (i) the calculation of below-ground root mass for determining aerenchyma area now uses the time-lagged (1 yr decay time) belowground-to-aboveground NPP (net primary productivity) ratio, instead of the instantaneous one, and (ii) in calculating the water availability for permafrost vegetation, root fraction is weighted over all soil layers down to last year’s maximum active layer depth, rather than the instantaneous active layer depth, thereby causing a slight delay in growth in the spring.
Table 4. A list of the models that contributed simulations to WETCHIMP. The “Wetland types” gives a quick overview of what kind of wetlands are used or simulated by each model. The explanations for the wetland types and soil datasets, as well as the full references are given in Table 3.

<table>
<thead>
<tr>
<th>Model</th>
<th>Resolution (lon × lat)</th>
<th>Wetland types</th>
<th>Climate forcing</th>
<th>Soil data</th>
</tr>
</thead>
<tbody>
<tr>
<td>CLM4Me</td>
<td>2.5° × 1.9°</td>
<td>Simulated inundated area based on predicted water table and runoff and an inversion to GIEMS</td>
<td>CRUNCEP</td>
<td>IGBP-DIS</td>
</tr>
<tr>
<td>DLEM</td>
<td>0.5° × 0.5°</td>
<td>Mixture between prescribed and simulated, rice mask by Leff</td>
<td>CRUNCEP</td>
<td>IGBP-DIS</td>
</tr>
<tr>
<td>IAP-RAS</td>
<td>0.5° × 0.5°</td>
<td>Olson data for wetlands (bogs/mires, swamps, heaths/moorlands, tundra)</td>
<td>CRU3.1</td>
<td>Peat in peatlands, loam elsewhere</td>
</tr>
<tr>
<td>LPJ-Bern</td>
<td>0.5° × 0.5°</td>
<td>Peatlands from NCSCD, inundated wetlands from GIEWS, rice mask by Leff</td>
<td>CRU3.1</td>
<td>HWSD</td>
</tr>
<tr>
<td>LPJ-WHyMe</td>
<td>0.5° × 0.5°</td>
<td>Peatlands from NCSCD</td>
<td>CRU3.1</td>
<td>n/a³</td>
</tr>
<tr>
<td>LPJ-WSL</td>
<td>0.5° × 0.5°</td>
<td>Inundated area from GIEWS, rice mask by Leff for all experiments except 3b</td>
<td>CRU3.1</td>
<td>FAO</td>
</tr>
<tr>
<td>ORCHIDEE</td>
<td>1° × 1°</td>
<td>Simulated, but annual mean over 1993–2004 adjusted to mean of GIEWS</td>
<td>CRUNCEP</td>
<td>FAO for mineral</td>
</tr>
<tr>
<td>SDGVM</td>
<td>0.5° × 0.5°</td>
<td>All simulated</td>
<td>CRU3.1</td>
<td>ISLSCP I</td>
</tr>
<tr>
<td>UVic-ESCM</td>
<td>3.6° × 1.8°</td>
<td>All simulated</td>
<td>CRU3.1+NCEPc</td>
<td>ISLSCP II</td>
</tr>
<tr>
<td>UW-VIC</td>
<td>100 km³</td>
<td>Simulated lakes and peatlands</td>
<td>CRUNCEP</td>
<td>FAO for mineral soils, Sheng for peatland fraction</td>
</tr>
</tbody>
</table>

* LPI-WHyMe is a peatland only model, thus no “soil” data is required. * LPI-WSL Exp. 3 is a mix between prescribed (GIEWS) and simulated inundation area based upon an empirical relationship between simulated water runoff and GIEMS. * Surface winds and diurnal temperature are taken from the NCEP reanalysis. * 100 km polar azimuthal equal area grid (EASE grid), resampled to 0.5° × 0.5°.

The CLM model requires a number of forcings in addition to the lower atmospheric boundary conditions and fluxes specified in the CRUNCEP forcing. For all the experiments except Experiment 3-opt, the standard CLM4 year 2000 conditions are used for atmospheric [CO2] experienced by plant stomata (except Experiment 4-CO2), atmospheric nitrogen deposition, atmospheric aerosol deposition, and vegetation distributions. For Experiment 3-opt, the same configuration as in Riley et al. (2011) is used, namely a spin-up to “1850” conditions using the 1850 [CO2], nitrogen, aerosols, and vegetation distributions, and then repeated 1948–1972 (Qian et al., 2006) corrected-NCEP forcing. A transient simulation from 1850–2004 is run using transient data for [CO2], nitrogen, aerosols, and vegetation, using repeated 1948–1972 forcing through 1972, at which point the model is switched to actual-year forcing through 2004.

The model is run at 1.9° × 2.5° resolution and the standard CLM4 datasets are used, except that the default CLM4 1 × 10⁶ km² of inland non-vegetated wetland area that were used in Riley et al. (2011), were eliminated. As described in that paper, the CLM4Me model requires three parameters at each grid cell to calculate the inundated fraction as a function of the modelled water table and lagged surface runoff, based on an inversion to Prigent et al. (2007) satellite observations for 1994–1998. For WETCHIMP, the parameters generated in a previous model described in Riley et al. (2011) were used (similar to Experiment 3-opt); however, the parameters were not re-optimised with the CRUNCEP forcing, hence the CRUNCEP 1990s inundated area (e.g. Experiment 2-trans) may differ from that simulated in Experiment 3-opt.

3.2 DLEM

The Dynamic Land Ecosystem Model (DLEM) is a process-based model that simulates daily carbon, water and nitrogen fluxes and pool sizes for land and riverine ecosystems. These pools and fluxes are influenced by changes in atmospheric chemistry ([CO2], ozone concentration and nitrogen deposition), climate, land cover and land use change, management practices (e.g. irrigation, fertilization, rotation, and technology improvement), and other disturbances (e.g. fire, hurricane, insects, disease, and forest harvest) (Tian et al., 2010, 2011a,b, 2012). For WETCHIMP, the disturbance submodel...
and the influence of ozone chemistry were not used due to a lack of spatially explicit driving data.

The soil biogeochemistry module simulates CH$_4$ uptake in upland ecosystems and emissions in wetland ecosystems. The mechanisms and algorithms for simulating CH$_4$ fluxes have been described in Tian et al. (2010, 2011b); Xu et al. (2010). DLEM requires input datasets for daily climate (average, maximum, and minimum air temperature, precipitation, gross radiation, and relative air humidity), atmospheric composition ([CO$_2$], nitrogen deposition and ozone), annual land use information, soil condition information (soil texture, pH, and soil depth), and topographic data (elevation, slope, and aspect).

Wetlands are defined as those areas that are inundated or saturated by surface water at a frequency and duration sufficient to support vegetation growth, which leads to five wetland types: (i) rice paddy, (ii) permanent herbaceous wetland, (iii) permanent woody wetland, (iv) seasonal herbaceous wetland and (v) seasonal woody wetland. The distribution map for different wetland types are determined based on the data from Stillwell-Soller et al. (1995); Aselman and Crutzen (1989) and Lehner and Döll (2004). DLEM simulates water transport to rivers based upon catchments, but does not explicitly move water through grid cells and thereby does not influence conditions in neighbouring grid cells. The version of DLEM used to simulate CH$_4$ fluxes for WETCHIMP has been described by Tian et al. (2010, 2011a); Xu et al. (2010). The CH$_4$ exchanges between ecosystems and the atmosphere are a combination of CH$_4$ production, oxidation, and transport from soil/water to the atmosphere. DLEM only considers CH$_4$ production from dissolved organic carbon (DOC), which is indirectly controlled by environmental factors including soil pH, temperature, soil texture and soil moisture content.

### 3.2.1 WETCHIMP set-up

The input data except those specifically mentioned in the experimental designs were kept constant at 1993 levels during

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**Fig. 5.** Fractional grid cells covered by permanent water bodies or ice are not considered to be wetlands (GICEW, http://luh.sr.unh.edu/). The Greenland ice sheet is masked out.
the WETCHIMP simulations. DLEM was run at a daily time step and a 0.5° × 0.5° spatial resolution. For experiments 1, 4, 5, and 6, DLEM used wetland extent data in 1993 as input, while for experiments 2 and 3, transient wetland extent data was used. Before running it at the global scale, DLEM was first calibrated against field experimental and observational data. The parameters and their values were described in Tian et al. (2010) and Xu et al. (2010). The water cycle module in DLEM used for the WETCHIMP simulations differs from the previous version used in DLEM publications (e.g. Tian et al., 2012). The main difference can be found in the water cycle module, which simulates the dynamics of inundated surface water extent and water table position Liu et al., 2012. The new water cycle module uses components of previously published models, TOPMODEL (Topography based hydrological model; Beven and Kirkby, 1979), SIMTOP (Simple TOPMODEL; Niu et al., 2005) and CLM (Oleson et al., 2008), to improve its soil and surface water dynamics (Liu et al., 2012).

To integrate the GIEMS inundation fraction and rice paddy (Leff et al., 2004) datasets with the DLEM-simulated wetland distribution, some of the DLEM parameterizations were adapted for a semi-prognostic approach to determine wetland area. DLEM separately simulated the extent of two major natural wetland types: permanent and seasonal wetlands. For the permanent natural wetland distribution, areas of the GIEMS dataset that were continually inundated during the growing season (May to August) of 1993–2004 were regarded as permanent wetlands. In these areas, the soil moisture was prescribed at saturation. Excluding areas of permanent wetlands, seasonal wetlands were determined using the DLEM prognostic parameterizations as discussed previously. Thus minimum annual wetland area is controlled by the GIEMS dataset, but daily and seasonal wetland area dynamics above this were determined by internal DLEM model dynamics and controlled by the maximum wetland extent data. The maximum (or potential) wetland extent data was generated based on GIEMS data by taking the largest inundated percentage in each grid cell during 1993–2004. DLEM did not specifically separate peatland from other wetland types. Instead, we separated the five wetland types (listed above) – based on the soil moisture condition and land cover types – for which the CH₄ module is run. The remaining non-inundated fraction in each grid cell was treated as upland land cover. We separately ran each fraction of the grid cells. Before simulations lakes and rivers were masked out using the GWLD (Global Wetland and Lakes Database) dataset (Lehner and Döll, 2004) and rice with the Leff dataset (Leff et al., 2004).

3.3 IAP-RAS model

The present version of the IAP-RAS (Institute of Applied Physics – Russian Academy of Sciences) wetland CH₄ emission module is described by Mokhov et al. (2007). The module consists of two parts; one for soil temperature calculations, and a second for calculations of CH₄ emissions. Soil temperature is calculated based on the model by Kudryavtsev et al. (1977). This model represents a generalized solution of Stephan’s problem with annual temperature variations at the lower boundary of the atmosphere, while accounting for the influence of snow and moss cover. Moss cover is prescribed as a layer of 10 cm thickness in presence of boreal and tundra vegetation. The depth of seasonal thaw/freeze depends on the annual variation of the near-surface temperature and precipitation. The influence of the effect of snow metamorphosis is ignored. The seasonal thaw depth was assessed based on the thickness of the active layer and temperature of the soil surface. Only soil layers to a certain limit depth were included in the calculations. In the standard version, the depths of 15 and 60 cm were used for tropical and extratropical zones, respectively. Similar depth values are obtained for the organic carbon content in soil at the characteristic peat density of 200 kg m⁻³ based on data from http://soils.usda.gov/use/worldsoils/mapindex/soc.html. Deeper layers were ignored in calculations of CH₄ emissions by wetlands. The amount of water in wetlands is considered to be always sufficient for inundation. Methane emissions are calculated based on the empirical model of Christensen and Cox (1995).

3.3.1 WETCHIMP set-up

For WETCHIMP simulations, the model was run at 0.5° × 0.5° resolution using CRU TS3.1 dataset as climate forcing (Mitchell and Jones, 2005). Wetland areal extent was prescribed according to the CDIAC (Carbon Dioxide Information Analysis Center) NDP017 dataset, also known as the Olson database (http://cdiac.esd.ornl.gov/ndps/ndp017.html). In this dataset, only areas with ecosystem codes 44 (“bog/mire of cool or cold climates”), 45+72 (“warm and hot wetlands”), 64 (“heath and moorland”), and 53 (“tundra”) are considered as wetlands. The inclusion of tundra regions as methane–producing area was specifically for WETCHIMP; earlier applications of IAP-RAS model neglected their contribution. Outside of wetlands, soil thermophysical parameters are homogeneously prescribed as loam everywhere.

3.4 LPJ-Bern

LPJ-Bern is a subsequent development of the Lund-Potsdam-Jena dynamic global vegetation model (Sitch et al., 2003; Joos et al., 2004; Gerber et al., 2003) that combines process-based, large-scale representations of terrestrial vegetation dynamics, soil hydrology (Gerten et al., 2004; Wania et al., 2009a), human induced land use changes (Strassmann et al., 2008; Stocker et al., 2011), permafrost and peatland establishment (Wania et al., 2009a,b) and simulation of biogeochemical trace gas emissions, such as CH₄ (Wania et al., 2010; Spahni et al., 2011; Zürcher et al., 2013).
### 3.4 WETCHIMP set-up

The CH$_4$ model within the LPJ-Bern version differs slightly from the LPJ-WHYMe (LPJ Wetland Hydrology and Methane) CH$_4$ model that was used in Wania et al. (2010) and Spahni et al. (2011). The main differences with respect to CH$_4$ emissions concern peatland modelling, global carbon cycle parameters and input data. The differences between the model as used in this study and Spahni et al. (2011) (and thus to LPJ-WHYMe version 1.3.1, Wania et al., 2010) are described below ordered by CH$_4$ source and sink category.

LPJ-Bern uses a different ebullition mechanism for CH$_4$ emissions from peatlands, which includes variations in partial pressure of CO$_2$ (Zürcher et al., 2013). The carbon balance over all layers is now preserved after every gas diffusion time step, whereas in LPJ-WHYMe a correction factor for carbon balance is applied at the end of the year. The possible plant functional types in peatlands are limited to flood-tolerant graminoids and Sphagnum mosses. Additionally, the prescribed fractional peatland cover per grid cell is taken from NCSCD (Tarnocai et al., 2007, 2009). NCSCD covers histels and histosols in the northern high-latitudes with a total area of 2.7 × 10$^6$ km$^2$, which is larger than the extent (2.06 × 10$^6$ km$^2$) used in Spahni et al. (2011). The global scaling factor used by Wania et al. (2010) to account for the lack of microtopography in the model is thus reduced from 75 to 26% to constrain CH$_4$ emissions from peatlands in 2004 to 28.2 Tg CH$_4$ yr$^{-1}$ (Spahni et al., 2011).

For CH$_4$ emissions in inundated areas the GIEMS monthly fractional inundation dataset for 1993–2004 was averaged by month (e.g., mean January, mean February, etc.). The fractional area of inundation per grid cell is further limited by the fraction of land available, i.e. land not covered by permanent water or ice (GICEW, http://luh.sr.unh.edu/). In grid cells containing peatlands (35–90°N), the inundated fraction was assumed to include peatlands. If the inundated fraction exceeds the constant peatland fraction, the difference is interpreted as the inundated fraction of mineral soils. This is different to the treatment in Spahni et al. (2011), where the inundated fraction was explicitly set to zero north of 45°N. The fraction of inundated areas was further divided into natural wetlands and rice agriculture using the scheme as described in Spahni et al. (2011). For these two categories the CH$_4$ to CO$_2$ conversion tuning parameter was adjusted to obtain total CH$_4$ emissions in 2004 of 81.3 Tg CH$_4$ yr$^{-1}$ for natural wetlands and of 43.1 Tg CH$_4$ yr$^{-1}$ for rice agriculture (Spahni et al., 2011).

For CH$_4$ emissions in wet mineral soils, the above changes were included and the remaining non-inundated and non-peatland land cover was taken as fractional area of mineral soils. These mineral soils can either function as a CH$_4$ source or sink, depending on their soil moisture (Spahni et al., 2011). For this study the CH$_4$ to CO$_2$ conversion factor – a global scaling factor – for CH$_4$ emissions from wet mineral soils was scaled to obtain emissions of 63.1 Tg CH$_4$ yr$^{-1}$ for 2004. For the CH$_4$ uptake the concentration-to-flux tuning factor was reduced to reach a global consumption of 25.8 Tg CH$_4$ yr$^{-1}$ (Spahni et al., 2011).

While the peatland fraction is a separate tile in each grid cell with its own carbon and soil water pools, the other three CH$_4$ source types and the sink share the same tile. So, for the non-peatland areas, there is no interaction between water table position and vegetation growth. Compared to Spahni et al. (2011), an updated soil type map based on the World Harmonized Soil Database (FAO/IIASA/ISRIC/ISSCAS/JRC, 2009) was used by selecting the dominant soil type in each 0.5° × 0.5° grid cell. However, soil properties for the corresponding 9 soil types were not changed to previous simulations (Spahni et al., 2011).

When model results were compared to other WETCHIMP models a unique feature was observed in CH$_4$ emissions from northern peatlands as simulated by LPJ-Bern. Emissions for the years 1998, 1999 and 2001 are extremely high for some sites compared to average peatland emissions (Melton et al., 2013). Investigating the model output showed the high emissions in those years originate from CH$_4$ gas reservoirs within and below frozen peat layers. Two processes are responsible for the establishment of these CH$_4$ gas reservoirs. First, in LPJ-Bern, frozen peat layers act as a barrier for gas diffusion thus allowing CH$_4$ to accumulate beneath. Second, if environmental conditions are suitable methanogenesis can occur in unfrozen layers, regardless whether layers above are frozen. Normally, CH$_4$ production in deeper layers is negligible, but for the WETCHIMP simulations a considerable amount of CH$_4$ gas accumulated beneath a frozen layer during the model spin-up procedure (1000 yr). This stored CH$_4$ could not escape to the atmosphere until the year 1998 in the transient run, the first year showing an exceptional boreal warming that led to thawing of peat layers and burst emissions of CH$_4$. While the process as such could be questioned, the magnitude of CH$_4$ emissions is certainly too large, as can be concluded from the model–data comparison in the Hudson Bay lowlands (Melton et al., 2013).

### 3.5 LPJ-WHYMe

Methane emissions for peatlands north of 35°N were simulated using LPJ-WHYMe (Wania et al., 2009a,b, 2010; Spahni et al., 2011). Location and fractional cover of peatlands are taken from NCSCD (Tarnocai et al., 2007, 2009). Monthly air temperature, precipitation, percentage cloud cover and number of wet days are taken from CRU TS3.1 to force all simulations. The data from 1901–1930 are used repeatedly to spin-up the model for 1000 yr before running it transiently, either for 1901–1930 or 1901–2009.

#### 3.5.1 WETCHIMP set-up

Instead of running LPJ-WHYMe only once for an average topographical microform, two parallel runs were executed for
each experiment. One run represents the wetter microform, lawns, which emit more CH$_4$, and the other one represents the drier microform, hummocks. The model modifications to approximate these two microforms include setting the daily drainage rate to 0.2 mm (lawns) and to 0.6 mm (hummocks), whereas it was 0 mm in Wania et al. (2010). These modifications lower the water table position in hummocks compared to lawns. The vegetation for hummocks is restricted to *Sphagnum* mosses, whereas lawns are able to grow any plant functional type depending on the water table position (Wania et al., 2009b). Methane emissions from the two parallel runs are averaged under the assumption that hummocks and hollows cover approximately the same surface area.

### 3.6 LPJ-WSL

The LPJ-WSL CH$_4$ model used in this analysis is the same as presented in Hodson et al. (2011), but has been recalibrated to a new set of regional CH$_4$ fluxes as noted below. The wetland CH$_4$ flux ($E$) at each 0.5° × 0.5° grid cell (x) and monthly time step ($t$) is calculated as a linear function of two scaling factors ($r_{CH4:C}$ and $f_{ecosys}$), wetland extent (A) and heterotrophic respiration ($R_{het}$) according to the following equation:

$$E(x,t) = r_{CH4:C} \cdot f_{ecosys}(x) \cdot A(x,t) \cdot R_{het}(x,t).$$ (1)

The notation in Eq. (1) has been modified from Hodson et al. (2011) to follow Table 5 (to convert between the notation in Eq. (1) and Hodson et al. (2011): $r_{CH4:C} = \beta \cdot f_{ecosys} = F; A = S$).

Together, $r_{CH4:C}$ and $f_{ecosys}(x)$ comprise the scaling ratio $F(x)$, which converts C to CH$_4$ fluxes and is a function of two weighted-regional scaling factors, one representing tropical (T) and another representing boreal (B) wetland climates (with temperate conditions represented as a combination of tropical and boreal). This approach allows the model to account for broad ecosystem differences in CH$_4$-emitting capacity between wetland types (Eq. 2). The weighting of wetland type (i.e. tropical vs boreal) in each grid cell is calculated based on surface temperature (Eq. 3).

$$F(x) = r_{CH4:C} \cdot f_{ecosys}(x) = \sigma(x) \cdot F_T + (1 - \sigma(x)) \cdot F_B$$ (2)

$$\sigma(x) = \exp((T(x) - T_{max})/8),$$ (3)

where $T(x)$ is the mean near-surface temperature between 1960 and 1990, and $T_{max} = 303.55$ °K. $\sigma(x)$ cannot exceed 1. Equations 2 and 3 correct unintentional omissions in both equations as written in Hodson et al. (2011).

### 3.6.1 WETCHIMP set-up

For WETCHIMP, we constrained the scaling ratios, $F_T$ and $F_B$, by minimizing the error between our model fit, inverse modelling results (Spahni et al., 2011) and a regional flux estimate from the Hudson Bay Lowlands (Pickett-Heaps et al., 2011), yielding $F_T = 0.152$ and $F_B = 0.049$. Total global wetland and rice fluxes were constrained at 215.8 Tg CH$_4$ yr$^{-1}$, wetland and rice fluxes north of 45° N at 39.6 Tg CH$_4$ yr$^{-1}$ (Spahni et al., 2011), and wetland and rice fluxes from 50–60° N and 75–96° W at 2.3 Tg CH$_4$ yr$^{-1}$ (Pickett-Heaps et al., 2011). These are different constraints than were used in Hodson et al. (2011).

$R_{het}$ was calculated using the LPJ-WSL dynamic global vegetation model (DGVM), based on the LPJv3.1 DGVM (Sitch et al., 2003; Gerten et al., 2004). The monthly climatology inputs (precipitation, mean temperature, cloud cover, wet days) were taken from CRU TS3.1 and the non-gridded annual CO$_2$ concentration inputs to LPJ-WSL are described in Hodson et al. (2011). In addition, as in Hodson et al. (2011), soil texture was prescribed from the Food and Agriculture Organization (Zobler, 1986), using a 2-soil layer hydrological model with a total soil depth of 1.5 m. For scenarios 2 and 3, a 1000 yr spin-up was implemented by recycling the first 30 yr of climate data (1901–1930) with preindustrial CO$_2$ concentrations to equilibrate soil and vegetation carbon pools, followed by a transient climate and CO$_2$ simulation running from 1901–2005. For scenarios 1, 4, 5, and 6, instead of using a repeating 1901–1930 climate, first a 12 monthly mean annual dataset was created and repeated until equilibrium is reached. The default soil respiration equation in LPJ was used for all scenarios except experiment 3.

We used the same temperature and moisture dependent equation as in Hodson et al. (2011), which allows the soil respiration to drop to zero when soil moisture is zero. For experiment 3, we used the soil moisture–respiration equation from Sitch et al. (2003) that fixes the soil respiration to a minimum of 0.25 in the upper one meter of soil, even when soil moisture drops to values below 0.25.

For all experiments except experiment 3, wetland extent (A) represents natural wetlands and lakes and is the original, monthly varying GIEMS inundation dataset processed to 0.5° × 0.5° spatial resolution with rice growing regions removed (Leff et al., 2004). For experiments 1, 4, 5, and 6, this wetland area satellite product was averaged across all years from 1993–2004 to create a 12 month mean wetland area product. For experiment 3, a combined satellite and model product was used, which is described in Hodson et al. (2011).

### 3.7 ORCHIDEE

The ORCHIDEE (Organising Carbon and Hydrology in Dynamic Ecosystems) model (Krinner et al., 2005) has been implemented with a wetland CH$_4$ emissions scheme. This version of ORCHIDEE has been previously used to simulate the evolution of wetland CH$_4$ emissions under future climate change (Koven et al., 2011) and to study the feedback between climate, atmospheric CH$_4$, and CO$_2$ (Ringleval et al., 2011). Simulations of ORCHIDEE for the current time period have also been performed and evaluated against top-down simulations to investigate the role of wetlands in the
current atmospheric CH$_4$ concentration growth rate (Pison et al., 2013).

The wetland CH$_4$ emissions, $E_{CH_4}(g,t)$, are computed for each grid cell, $g$, and for each time step, $t$, through the following equation:

$$E_{CH_4}(g,t) = \sum_{WTP_i} S_{WTP_i}(g,t)D_{WTP_i}(g,t),$$

(4)

where $S_{WTP_i}$ is the fraction of $g$ covered by a wetland whose water table position (WTP) is equal to WTP$_i$. $D_{WTP_i}$ is the CH$_4$ flux (i.e. g CH$_4$ m$^{-2}$ of wetland) time$^{-1}$) for a wetland whose water table position is equal to WTP$_i$. $S_{WTP_i}$ and $D_{WTP_i}$ are, respectively, computed by (i) the coupling between a TOPMODEL (Beven and Kirkby, 1979) approach and ORCHIDEE, and (ii) the coupling between a (slightly modified) version of the Walter et al. (2001a) model and ORCHIDEE.

The main modification to the Walter et al. (2001a) model, as described in Ringeval et al. (2010), concerns the methanogenesis substrate. A fraction, $\alpha$, of the natural labile carbon pool computed by ORCHIDEE is used to estimate the methanogenesis substrate. The $\alpha$ parameter also encompasses the methanogenesis base rate at the reference temperature (see Ringeval et al., 2013). $\alpha$ has been optimized against three sites then extrapolated on all grid cells sharing the same vegetation type (boreal, temperate or tropical).

In contrast to LPJ-WHYMe and its derivatives, ORCHIDEE did not implement wetland-specific PFTs (plant functional types). Instead, a fraction of the mean natural labile carbon pool over the grid cell is used to estimate the substrate supply.

For the computation of $S_{WTP_i}$ in each grid cell, TOPMODEL allows distribution of the mean water deficit computed by ORCHIDEE according to the sub-grid topographic index distribution. This leads to the diagnostic of the grid-cell fraction with a null deficit. The mean deficit over the grid cell is computed from a gap to the field capacity (and not to the saturation, which cannot be reached in ORCHIDEE). The saturated wetland extent is computed from these “field capacity extents” using a shift of the topographic index distribution into each grid cell. The value of this shift is the same for all grid cells and has been chosen to obtain a global coverage by wetlands close to 4%. The simulated wetland extent has been evaluated both through the induced modification on the simulated river flows and against the GEMS data (Ringeval et al., 2012). The TOPMODEL approach is used to simulate not only the saturated wetland extent but also the wetland extent with a WTP below the soil surface.

### 3.7.1 WETCHIMP set-up

In the simulations performed for the WETCHIMP intercomparison, the WTP$_i$ values chosen for each grid cell are: 0, −3 and −9 cm. A value of water table position equal to 0 means that the water is at the soil surface while negative values corresponds to water table position below the soil surface. Thus, for each time step, three fractions of each grid cell are given: (i) a fraction covered by a saturated wetland, (ii) a fraction covered by a wetland with a mean WTP equal to −3 cm (i.e. where the deficit is between 0 and −6 cm) and (iii) a fraction covered by a wetland with a mean WTP equal to −9 cm (i.e. where the deficit is between −6 and −12 cm).

### Table 5. A general description of the dependencies of CH$_4$ production, CH$_4$ oxidation (does not include atmospheric CH$_4$ oxidation) and CH$_4$ flux. All of the fluxes are modulated by a CH$_4$-producing area (see Fig. 6). Parameters and variables used in the models were harmonized where possible, but identical names do not indicate identical values in the different models.

<table>
<thead>
<tr>
<th>Model</th>
<th>CH$_4$ production ($P$)</th>
<th>CH$_4$ oxidation ($O$)</th>
<th>Atmospheric CH$<em>4$ oxidation ($O</em>{atm}$)</th>
<th>Net CH$_4$ flux ($F$)</th>
</tr>
</thead>
<tbody>
<tr>
<td>CLM4Me</td>
<td>$P = R_{net}/CH_4/C/soil/ft/\Delta Q_{10}$</td>
<td>$O = O_{max}/f(\Delta Q_{10})/f(CH_4)/f_{soil}/Q_{10}$</td>
<td>$O_{atm} = R_{max}/f(\Delta Q_{10})/f(CH_4)/f_{soil}/Q_{10}$</td>
<td>$F = (P - O)/transport - O_{atm}$</td>
</tr>
<tr>
<td>DLEM</td>
<td>$P = P_{max}/C_{labile}/f/ft/\Delta Q$</td>
<td>$O_{trans} = O_{trans,max}/f(plant)/trans/ft$</td>
<td>$O_{trans} = O_{soil,max}/f(CH_4)/f_{trans}/ft/oxid/\Delta Q_{10}$</td>
<td>$F = (P - O)/transport - O_{atm}$</td>
</tr>
<tr>
<td>IAP-RAS</td>
<td>$P = fT$</td>
<td>–</td>
<td>–</td>
<td>$F = P/soil \Delta Q_{10}$</td>
</tr>
<tr>
<td>LPJ-Bern peat</td>
<td>$P = R_{net}/CH_4/C/root/fWTP$</td>
<td>$O = f(\Delta Q_{10})/f(CH_4)/f_{soil}/Q_{10}$</td>
<td>–</td>
<td>$F = P/soil \Delta Q_{10}$</td>
</tr>
<tr>
<td>LPJ-Bern wetlands</td>
<td>$P = R_{net}/CH_4/C$</td>
<td>–</td>
<td>–</td>
<td>$F = P$</td>
</tr>
<tr>
<td>LPJ-Bern rice</td>
<td>$P = R_{net}/CH_4/C$</td>
<td>–</td>
<td>–</td>
<td>$F = P$</td>
</tr>
<tr>
<td>LPJ-Bern wetsoils</td>
<td>$P = R_{net}/CH_4/C/soil/fWTP$</td>
<td>–</td>
<td>$O_{atm} = f(CH_4)/f_{soil}/Q_{10}$</td>
<td>$F = P - O_{atm}$</td>
</tr>
<tr>
<td>LPJ-WHYMe</td>
<td>$P = R_{net}/CH_4/C/root/fWTP$</td>
<td>$O = f(\Delta Q_{10})/f(CH_4)/f_{soil}/Q_{10}$</td>
<td>–</td>
<td>$F = (P - O)/transport$</td>
</tr>
<tr>
<td>LPJ-WL</td>
<td>$P = R_{net}/CH_4/C/soil/fWTP$</td>
<td>–</td>
<td>–</td>
<td>$F = P$</td>
</tr>
<tr>
<td>ORCHIDEE</td>
<td>$P = R_{n}/C_{labile}/fWTP/f/\Delta Q_{10}$</td>
<td>$O = f(WTP)/f(CH_4)/f_{soil}/Q_{10}$</td>
<td>–</td>
<td>$F = (P - O)/transport$</td>
</tr>
<tr>
<td>SDGVM</td>
<td>$P = R_{net}/CH_4/C/root/fWTP/\Delta Q_{10}$</td>
<td>$O = 0.9P$</td>
<td>–</td>
<td>$F = P - O$</td>
</tr>
<tr>
<td>UW-VIC</td>
<td>$P = R_{f}/f_{soil}/ft/\Delta Q_{10}$</td>
<td>$O = f(CH_4)/f_{soil}/Q_{10}$</td>
<td>–</td>
<td>$F = (P - O)/transport$</td>
</tr>
</tbody>
</table>
As in Ringeval et al. (2011) the wetland extent is corrected to subtract the systematic biases of the model by normalizing the mean yearly wetland extent to the GIEMS data (i.e. both the seasonal and year to year variability come from TOPMODEL). In the WETCHIMP simulations, a $Q_{10}$ equal to 3 (close to the mean value in Ringeval et al., 2010) has been chosen for all grid cells. As in Ringeval et al. (2013), the reference temperature for methanogenesis is defined as the mean surface temperature computed by ORCHIDEE when forced by the 1960–1991 CRUNCEP climatology.

### 3.8 SDGVM

The SDGVM (Sheffield Dynamic Global Vegetation Model; Woodward et al., 1995; Beerling and Woodward, 2001) was used in conjunction with a modified version of the Cao et al. (1996) wetland emissions model to perform the WETCHIMP simulations. The modelling set-up follows the approach of Singarayer et al. (2011), however in that study an equilibrium approach was taken wherein the vegetation and CH$_4$ models were forced with averaged (30 yr) climatologies from a series of general circulation model simulations. For WETCHIMP, a transient approach was required whereby the models were forced sequentially with monthly climatic inputs. The SDGVM and CH$_4$ models were therefore run in a similar manner as described by Hopcroft et al. (2011), in which a transient set-up of the model is described.

SDGVM requires monthly mean inputs of surface air temperature, precipitation and relative humidity, as well as a map of soil texture and global annual mean atmospheric CO$_2$ concentrations. The CRU vapour pressure was converted to relative humidity using standard formulae, whilst the temperatures and precipitation were used without modification.

In the CH$_4$ model, the supply of carbon substrate for methanogenesis is taken to scale with 1 m soil heterotrophic respiration as simulated in SDGVM. 1 m is the model's soil root depth. CH$_4$ production, $M_{prod}$, is then given by

$$M_{prod} = R_{het} P_0 f_w WTP f_T (T),$$

where $R_{het}$ is the soil heterotrophic respiration rate (g C m$^{-2}$ month$^{-1}$) from SDGVM, $P_0$ represents the fraction of decomposed matter converted to CH$_4$ under optimal conditions (0.47, see Cao et al., 1996). $f_w$ and $f_T$ are dimensionless scaling functions, which parameterize the effects of water table position (WTP, in cm, positive up from the soil surface) and temperature ($T$ in °C) on emission rates. These are given by

$$f_w (WTP) = 0.383 e^{0.096 WTP} \quad WTP \leq 10 \text{ cm};$$

$$f_T (T) = \frac{e^{0.04055 T}}{3.375}, \quad 5 \degree C < T \leq 30 \degree C,$$

where $f_w$ is 1.0 for WTP > 10 cm and $f_T$ is 1.0 for $T > 30 \degree C$ and 0.0 for $T \leq 5.0 \degree C$. Emissions thus increase with increasing water table depth (up to 10 cm). $f_w$ follows observations from Roulet et al. (1992) and Eq. (7) implies a $Q_{10}$ value of 1.5. 90 % of CH$_4$ produced is assumed to be oxidised. The water table position in each grid cell is calculated from the SDGVM simulated 1 m total soil moisture content using the relations from Cao et al. (1996) for tundra (their Eqs. 15, 16) and a constant global soil porosity.

SDGVM also includes the biogeochemical coupling between the above- and below-ground carbon and nitrogen cycles. This is an important feature of modelling realistic changes in land surface primary production, which depends on, and should be in line with, realistic biological and anthropogenic nitrogen fixation rates (Hungate et al., 2003). In SDGVM, litter production influences soil carbon and nitrogen pools via the Century soil nutrient cycling model (Parton et al., 1993), which in turn feedback to influence the primary production of vegetation; details are provided elsewhere (Beerling and Woodward, 2001).

#### 3.8.1 WETCHIMP set-up

The model output saved from SDGVM simulations includes annual NPP, soil carbon content, monthly heterotrophic respiration, soil moisture content, and GPP. The monthly outputs of CH$_4$ emissions and water table position were saved from the CH$_4$ model. The experiment protocol called for monthly and annual maximum wetland area (mmax_weta, amax_weta). Wetland area in this model is not used by the emissions model, which is instead a function of water table position (see Eq. 5). CH$_4$ emitting area could be used as a proxy for wetland area, but this would include grid cells with a water table far from the surface and with very small CH$_4$ fluxes. Since most of the other models in the inter-comparison were parameterized using inundated area from satellite observations, the SDGVM output was tailored to be somewhat comparable. The wetland area was here taken as simulated inundated area. Since there is no sub-grid hydrology in the model, in each grid cell this area will either be 0 or the area of the grid cell. However, the model includes a correction for sub-grid orography based on the ETOPO5 dataset, which is applied to CH$_4$ emissions. The orographic correction, $F_{corr}$ is computed as:

$$F_{corr} = \frac{0.01 - S_{max}}{0.01},$$

where $S_{max}$ is the maximum gradient between a grid cell and its 8 nearest neighbours. This correction was also applied to the calculated inundation area. Two climate-dependent conditions on CH$_4$ emissions are also currently used within the model: (i) the monthly air temperature must be above 5 °C, and (ii) if the temperature in a given grid cell during the current year is always > 0 °C, then in a given month, the evapotranspiration must not exceed precipitation. These two conditions were both also used to correct the simulated inundated area, so that it is more directly relevant to the simulated CH$_4$ emissions.
3.9 UVic-ESCM

The UVic Earth System Climate Model (Weaver et al., 2001; Meissner et al., 2003) is an intermediate complexity climate model consisting of several coupled components: an energy–moisture balance atmospheric model, a three-dimensional ocean general circulation model, a dynamic–thermodynamic sea ice model and a land surface scheme, which includes a representation of vegetation dynamics. The model was recently modified to include a representation of permafrost and global wetlands (Avis et al., 2011). Wetlands in the UVic model are determined using empirically derived threshold values for unfrozen soil moisture content and terrain slope so that wetlands are present where ground is sufficiently wet and flat (Kaplan, 2002). Wetlands are either “on or off” in a particular grid cell. If they satisfy the soil moisture criterion they are “on” and they occupy the fraction of the grid cell with the requisite terrain slope. Organic and mineral soil properties were specified by using the ISLSCP-II (International Land-Surface Climatology Project II) datasets as the model is presently incapable of generating the observed high soil carbon values in northern high latitudes (none of the participating models is able to couple soil biogeochemical with thermal characteristics yet).

3.9.1 WETCHIMP set-up

The UVic model is nominally run in a fully coupled configuration with coupling between atmospheric, ocean, land surface and other model components. For the purpose of participating in WETCHIMP, the land surface scheme was decoupled from the other model components and run in an offline configuration. This offline configuration uses the monthly CRU data to drive the land surface scheme. The model smoothly interpolates between these CRU fields to obtain data for a particular time step. Some of the atmospheric code from the coupled model’s energy–moisture balance model (Weaver et al., 2001) was retained to calculate factors such as incoming solar radiation, which needs to be specified to the land surface scheme and was not specified in the monthly CRU datasets.

Surface air temperature, precipitation and relative humidity were specified from the CRU data, while surface winds and the diurnal temperature range were drawn from the NCEP reanalysis. Incoming long- and short-wave radiation are calculated within the model’s atmospheric module.

The land surface scheme also has the capacity to simulate vegetation dynamics using the TRIFFID (Top-down Representation of Interactive Foliage and Flora Including Dynamics) dynamic vegetation model (Cox, 2001). As vegetation parameters were derived based on the coupled model climate, for the purposes of the WETCHIMP simulations, the vegetation distribution and characteristics were fixed and set equal to their mean year-1900 values from an equilibrated version of the coupled model, rather than adjusting vegetation parameters to fit the CRU climate data. Monthly mean fields of vegetation fraction, leaf area index (LAI), vegetation height and litterfall are then obtained from the mean year-1900 model output. The vegetation dynamics are consequently switched off for the offline run, and vegetation parameters for a given model time step are specified by smoothly interpolating between these monthly fields. Non-plant competition based vegetation parameters do remain interactive in the model. For example, plant stomata still open and close in response to factors like changing soil moisture or atmospheric CO$_2$ concentration.

For the equilibrium run, the model was forced repeatedly with 1901–1931 data for 2000 yr. This time period was found to be sufficient to allow for equilibration of soil moisture and temperature variables. When applying the CO$_2$ temperature and moisture runs, the model was forced with 1901–1931 spin-up period was repeated for an additional 2000 yr to allow a new equilibrium to be established.

3.10 UW-VIC

The University of Washington team used the Variable Infiltration Capacity (VIC) model, version 4.1.2, with some extensions specifically tailored for the modelling of boreal peatlands described in Bohn et al. (2007) and Bohn and Lettenmaier (2010). UW-VIC is a large-scale hydrologic model that balances the water and energy budgets of the land surface at an hourly time step and spatial resolutions ranging from 1° to 5 km. Most of the model physics are described in Liang et al. (1994). Land cover is represented as a collection of “tiles”, each containing a different plant functional type, overlaying a single soil column divided into 3 hydrologic layers, down to varying depths, but generally no deeper than 3 m. While UW-VIC does not track the storage of carbon in biomass, it computes NPP via a scheme taken from the BETHY (Biosphere Energy-Transfer and Hydrology) model (Knorr, 2000). The seasonal cycle of LAI is prescribed at each grid cell based on the MODIS (Moderate Resolution Imaging Spectroradiometer) LAI product (Myneni et al., 2002). Stomatal resistance is a function of day length, temperature, soil moisture, and vapor pressure deficit. UW-VIC models permafrost and the soil temperature profile via the finite difference scheme of Cherkauer and Lettenmaier (1999) with an exponential node distribution down to 50 m depth and a no-flux bottom boundary condition. Thermal properties of organic soil are also taken into account (Farouki, 1981). To account for dynamic surface water storage (lakes and seasonally flooded wetlands) UW-VIC’s lake/wetland model was employed (Bowling and Lettenmaier, 2010). This feature allocates one land cover tile to contain a combination of a lake (representing all lakes in the grid cell) and its surrounding catchment. Within the tile, the inundated area fraction is dynamic, changing as a function of storage and bathymetry. The extent of permanent lakes were assumed equal to the minimum annual observed inundation...
extent from the remotely observed inundation datasets (either Schroeder et al. (2010) or GIEMS, see Table 3). Thus, while the minimum and maximum possible extent of inundation within a grid cell are prescribed, the actual inundated extent is a dynamic function of environmental conditions. In the exposed portion of the tile, the water table position is assumed to have a distribution based on peatland microtopography: the peatland consists of a mix of hummocks (or ridges) and hollows (or pools), with the peat underneath hummocks up to 70 cm thicker than under the deepest points of the hollows. The fraction of peatland covered by hummocks is a calibrated parameter. Local water table position under any given point is computed as a function of soil moisture via the formulation described in Frolking et al. (2002). Methane emissions were computed for the lakes, inundated wetlands, and each point in the water table position distribution in the exposed wetlands as a function of water table position, soil temperature, and NPP via the model of Walter and Heimann (2000). A lake emission rate of 375 mg CH$_4$ m$^{-2}$ d$^{-1}$ was originally assumed during the ice-free season and half that rate during ice-covered season. This rate was found to be in error due to an artefact in the satellite data used to parameterize the lake emission rates. A value of 0 mg CH$_4$ m$^{-2}$ d$^{-1}$ is now used. The new lake emission rate results in lake CH$_4$ emissions that more closely resemble observations from the area. The influence of the new value can be observed by comparing from Melton et al. (2013) Fig. A1j (375 mg CH$_4$ m$^{-2}$ d$^{-1}$) and Fig. 8j (0 mg CH$_4$ m$^{-2}$ d$^{-1}$).

### 3.10.1 WETCHIMP set-up

For these simulations, each grid cell was separated into two parts: an upland fraction, underlain by mineral soils, with soil textures supplied by the FAO Digital Soil Map of the World (Batjes, 1997); and a lake/wetland fraction, underlain by peat soils, with peat depths given by the database of Sheng et al. (2004) and other characteristics taken from Letts et al. (2000). Simulations were run separately for each portion of the grid cell. The lake/wetland portion of each grid cell was determined as the superset of the Sheng et al. (2004) peatland map; wetlands, wet tundra, and croplands (so that nearby lakes could have a surrounding catchment) given by the Bartalev et al. (2003) land cover classification; and lakes given by the Global Lake and Wetland Database (GLWD; Lehner and Döll, 2004). Bathymetries for the lake/wetlands were estimated by combining lake size distributions from the GLWD; average lake depths from literature for bog pools, Arctic thaw lakes, and other boreal lakes; and topography of surrounding wetlands from the ASTER (Advanced Spaceborne Thermal Emission and Reflection Radiometer) (Hayakawa et al., 2008) and STRM (Shuttle Topography Radar Mission; Digital Elevation Model) (Farr and Kobrick, 2000) DEMs. When lake storage increased beyond the bounds of the “permanent” lake, it was allowed to flood the surrounding wetlands, with drainage rate controlled by a calibrated parameter. Both this parameter and the area fraction of hummocks within the peatland were calibrated to optimize the match with global inundation datasets. For the optimized runs (Experiment 3), the global daily AMSR-E/QuickSCAT-based dataset of Schroeder et al. (2010) was used; for all other runs, the GIEMS dataset was used (Pingen et al., 2007; Papa et al., 2010). Parameters for the CH$_4$ emissions model were calibrated to optimize the in situ observations of Glagolev et al. (2010) across West Siberia. For experiments other than experiment 3, UW-VIC’s parameters were calibrated to match the June/July/August average inundation given by the GIEMS dataset over the period 1993–2004 and used for years outside of that range.

### 4 Results and discussion

We provide two conceptual overviews of the participating models highlighting similarities as well as differences between the chosen approaches. These overviews are designed to assist discussions of the differences in modelling results (Melton et al., 2013), but they also represent the first attempt at conceptually describing the state-of-the-art approaches used in wetland extent and wetland CH$_4$ modelling. For the conceptual figures describing the modelling approaches, we defined two variables of interest: the CH$_4$ producing area (Fig. 6) and CH$_4$ flux (Fig. 7). We used these metrics to explore the dominant processes responsible for differences between the models.

#### 4.1 CH$_4$ producing area

We use the term “CH$_4$ producing areas” (MPAs) to include all terrestrial areas that may produce CH$_4$ biogenically. We include wet mineral soils, presently only simulated by LPJ-Bern, that are proposed to function as a CH$_4$ source or sink depending on the soil moisture level. The participating models use a large diversity of methods to determine MPAs (Fig. 6). We identified the features of the models that we found most strongly controlled the MPAs and visualized the concepts of the models.

The starting points to locate MPAs are either “Prescribed constant wetland extents”, “Remotely-sensed inundation” or a “Hydrological model” (Fig. 6). The simplest case of estimating MPAs is where “Prescribed constant wetland extents” are taken from annually, non-varying distribution maps, and are used without modifications. This approach is applied by LPJ-WHYMe and LPJ-Bern (peatlands), which use the northern peatland map from NCSCD (Tarnocai et al., 2007, 2009), and IAP-RAS, which uses the Olson dataset for global MPA location. A similar approach takes seasonally varying “Remotely-sensed inundation” to prescribe MPAs. LPJ-Bern wetlands uses an averaged monthly mean extent from GIEMS, while LPJ-WSL (all experiments except 3) uses the GIEMS dataset without modification. A further step up the
complexity ladder is LPJ-Bern wetsoils, the most basic wetland extent that uses model output. LPJ-Bern wetsoils uses a “Hydrological model” to derive “Soil moisture” and “Unsaturated” MPAs. “Unsaturated” means that the pore-space in the soil is not completely filled with water. This could be the case when – even though a water table position is calculated – it is below the surface or when the soil moisture is estimated as a homogeneous average over the soil depth and its values do not reach saturation. Next, we include approaches that comprise of “Topography” in addition to “Hydrological model” as an additional factor to locate “Unsaturated” areas (UVic-ESCM). SDGVM uses a similar approach to UVic-ESCM but simulates “Water table position” before determining “Unsaturated” as well as “Saturated/inundated” MPAs. CLM4Me, DLEM, ORCHIDEE, UW-VIC and LPJ-WSL (Exp. 3) all use “Remotely sensed inundation” (GIEMS) data in their approaches, but they use these data in different ways: e.g. ORCHIDEE guides the mean simulated wetland extent over the 1993–2004 period and CLM4Me uses the GIEMS dataset to invert for parameters that allow the hydrological state (i.e. water table depth and runoff) to be used to determine wetland extent. More details on the use of GIEMS can be found in the description of each model (Sect. 3). Once the “Water table position” is determined, CLM4Me, DLEM, ORCHIDEE, and UW-VIC identify the MPAs that are either “Unsaturated” or “Saturated/inundated” while LPJ-WSL (Experiment 3-opt) determines MPAs which are “Saturated/inundated” only. The UW-VIC model is the most complex model and takes advantage of all of the features included in Fig. 6, using the fractional peatland cover by Sheng et al. (2004) only as maximal boundaries, rather than as a fixed map.

From the conceptual overview (Fig. 6) one can see that the only two models that simulate MPAs without the guidance of other wetland-relevant observations are the UVic-ESCM and the SDGVM. The difference between these two models is that the UVic-ESCM uses only soil moisture (as well as topography) to find “Unsaturated” areas, whereas the SDGVM also calculates the water table position to find “Unsaturated” as well as “Saturated/inundated” areas. As the UVic-ESCM model was designed to identify wetland areas, not specifically MPAs (it presently has no CH₄ model), the model uses grid-cell mean unsaturated soil moisture values and terrain slope as a mean of approximating saturated areas.

An additional area of uncertainty that should be noted is the influence of anthropogenic changes to the land surface. Models that explicitly use the GIEMS dataset account for rice agriculture by masking out those regions (Table 3) while also implicitly including areas of human alteration such as wetland drainage, conversion to farmland, etc. Models that independently simulate wetland extent will not be sensitive to these alterations and this could lead to an overestimate of wetland area in these regions. Small lakes could also contribute to an overestimated wetland area for some models. Presently only LPJ-Bern and UW-VIC masks these lakes (Table 3).

4.2 CH₄ flux

The second variable we analyse in detail is the “CH₄ flux” calculation by nine out of the ten participating models – the UVic-ESCM does not yet include CH₄ fluxes. Figure 7 shows which pools and processes models consider to determine CH₄ flux. All models but IAP-RAS base their CH₄ production on some kind of carbon flux, where two groups can be distinguished – one that uses “Wetland PFTs” and one that uses “Upland PFTs” to simulate vegetation net primary production (NPP); only DLEM utilizes NPP (and also GPP) simulated by both types of PFTs for CH₄ production. The UW-VIC model uses NPP in the algorithm for CH₄ production, ORCHIDEE uses a fraction of the most labile of the “Litter + soil C” pool and all remaining models use “Heterotrophic respiration” as the basis for their “CH₄ production” (see also Table 5). LPJ-Bern, LPJ-WHyMe and DLEM add “Exudates”-derived carbon to the “Heterotrophic respiration” calculation.

All models calculate CH₄ production and half of the models consider “Transport” mechanisms such as ebullition, plant-mediated transport, and diffusion to derive “CH₄ fluxes” (Fig. 7, Table 5). This table also gives insight into which models include oxidation of soil-derived CH₄ and how they combine production and oxidation rates to simulate the final net CH₄ flux. Only three of the models include atmospheric CH₄ oxidation (CLM4Me, DLEM, LPJ-Bern wet soils). Thus, we include soil oxidation of atmospheric methane in Table 5 for completeness, but we excluded soil oxidation of atmospheric methane from model results as far as possible in order to compare gross CH₄ fluxes. The separation of gross CH₄ fluxes and atmospheric CH₄ uptake fluxes was not completely feasible in CLM4Me as the CH₄ uptake occurs implicitly in the reaction-transport solution, although for Melton et al. (2013) an estimate was determined to allow better comparison between models. Across the models, the complexity of equations for CH₄ production covers a wide range. The IAP-RAS model is the simplest model relating CH₄ production only to temperature, whereas all other models use some estimate of the available carbon flux rate. All but two models (IAP-RAS and LPJ-WSL) include some kind of soil-derived CH₄ oxidation, which can be as simple as reducing production by a fixed fraction (SDGVM) or by including up to five different terms in the equation (Table 5).

4.3 Methane producing area and methane flux

Mean annual maximum extent of wetland area and mean annual CH₄ fluxes for Experiment 1-equil are shown in Fig. 8. Total MPAs and CH₄ emissions for each model are listed in the bottom left corner of each sub-figure. Two of the models are regional models (LPJ-WHyMe and UW-VIC), the rest
are global models in which the UVic-ESCM does not simulate CH$_4$ fluxes. CLM4Me, DLEM, LPJ-Bern (without mineral soils), LPJ-WSL, and ORCHIDEE share similar wetland distributions due to their varying degrees of reliance on remotely sensed inundation data (see Fig. 6). The similarity of approaches is also reflected in the total MPA of these models (CLM4Me: 6.8 × 10$^6$, DLEM: 7.9 × 10$^6$, LPJ-Bern without mineral soils: 7.9 × 10$^6$, LPJ-WSL: 7.4 × 10$^6$, and ORCHIDEE: 9.2 × 10$^6$ km$^2$). Two models (IAP-RAS and LPJ-Bern with wet mineral soils) stand out visually because of their large areas of 80–100% MPA per grid cell. The IAP-RAS model uses a binary approach — either a grid cell is a wetland or it is not — resulting in a total MPA of 20.3 × 10$^6$ km$^2$, which is an entirely prescribed amount. Given the definition of the wet mineral soils as CH$_4$ source, the LPJ-Bern wet mineral soils map should be interpreted as a map of “potential CH$_4$ emissions in at least one month per year”. Since LPJ-Bern does not use a sub-grid-cell hydrology for wet mineral soils to estimate the CH$_4$ production capacity, the wet mineral soils component of LPJ-Bern is also a binary approach. However, the extent of wet mineral soils in a given grid cell can be reduced by peatland area and inundated area so that they jointly sum to 100%, but as soon as a grid cell qualifies as wet mineral soils, the MPA of that grid cell is 100%. This approach leads to the largest total MPA of 76.6 × 10$^6$ km$^2$ of the WETCHIMP models.

The only two models that use an explicit water balance scheme to simulate wetland extent without relying on wetland or inundation datasets are the SDGVM and the UVic-ESCM (Fig. 8). They show a similar spatial distribution but differ notably in Eastern Siberia, the western United States and northern Canada. SDGVM uses soil moisture content to first diagnose water table position and then MPAs, whereas UVic-ESCM uses soil moisture directly to derive MPAs. The differences between these two models could be related to parameterization of permafrost (present in UVic-ESCM, but absent in SDGVM) and other soil physics or hydrology parameters, i.e. hydraulic conductivity, porosity, etc. Further, the percentages of grid cells covered by wetlands are generally higher for SDGVM than for the UVic-ESCM leading to a higher overall wetland area of 34.9 × 10$^6$ km$^2$ vs 14.9 × 10$^6$ km$^2$, respectively.

Of the two regional models, LPJ-WHYMe uses a fixed peatland distribution (Fig. 8e) whereas the UW-VIC model uses the most sophisticated method of all participating models to simulate saturated and unsaturated wetland areas in the West Siberian Lowlands (Fig. 8j). A comparison focused on the West Siberian Lowlands is planned to evaluate the differences between a highly regionalized model like the UW-VIC model and the rest of the WETCHIMP models (T. Bohn, personal communication, June 2012).

Simulated CH$_4$ fluxes of nine of the participating models are shown on the right hand side in Fig. 8. Methane fluxes ranged from 0 to over 250 g CH$_4$ m$^{-2}$ of wetland yr$^{-1}$ with CLM4Me, DLEM, LPJ-WSL, and ORCHIDEE showing

![Fig. 7. Conceptual overview of the pathways from the carbon source to “CH$_4$ flux” (CH$_4$ flux per m$^2$ of wetland) in the participating models. Each model (and in some cases, version) is represented by a different colour. The flow of a particular model starts with the model’s name. The hatching of the lines indicates that the model uses wetland-specific PFTs, whereas “Upland PFTs” indicates that the model uses the already existing PFTs used for upland ecosystems. “NPP” stands for net primary production, “Exudates” are root exudates carbon pool. All models use some sort of temperature dependence when calculating “Heterotrophic respiration”. The models then calculate “CH$_4$ production” and the oxidation based on the equations given in Table 5. Some models include the effect of “Transport” mechanisms explicitly, whereas others include transport only implicitly by either producing less CH$_4$ or oxidizing it before emitting it to the atmosphere. All models use some sort of temperature dependence when calculating NPP, heterotrophic respiration, and/or CH$_4$ production. In this figure, “LPJ-Bern nonpeat” includes both wetlands and wetsoils, which also incorporate plant exudates (graphical simplification).](https://www.geosci-model-dev.net/6/617/2013/)
widespread high fluxes (Fig. 8a, b, g, h) and IAP-RAS, LPJ-Bern, LPJ-WHyMe, SDGVM, and UW-VIC showing low fluxes (Fig. 8c, d, e, i, j). Of the five models that show widespread high fluxes, three base their CH$_4$ flux on upland PFTs (CLM4Me, LPJ-WSL, ORCHIDEE), and one on both wetland and upland PFTs (DLEM) (Fig. 7). Of the four models that show low CH$_4$ fluxes, three rely on wetland PFTs (LPJ-Bern, LPJ-WHyMe, and UW-VIC), one on upland PFTs (SDGVM) and one does not rely on PFTs at all (IAP-RAS). This could indicate a general tendency to higher CH$_4$ fluxes when upland PFTs instead of wetland PFTs are used to simulate NPP. Some of the models show higher fluxes
in the tropics than in the extra-tropics (CLM4Me, IAP-RAS, SDGVM), whereas others show equally high fluxes (DLEM, LPJ-Bern, LPJ-WSL, ORCHIDEE), which may be linked to the model-inherent temperature sensitivities of, e.g. NPP, heterotrophic respiration or CH₄ production, but without specific parameter sensitivity studies it is impossible to evaluate where the differences arise from.

The patterns of CH₄ fluxes do not always match the pattern of wetland distribution (e.g. compare wetland area and CH₄ fluxes for the Amazon from CLM4Me). Total CH₄ emissions for a grid cell are calculated as the product of fluxes and wetland area (except for CLM4Me, which also considers production in upland soils). Therefore, models may simulate similar global CH₄ emissions with completely different MPAs and CH₄ fluxes (e.g. CLM4Me: 186 Tg CH₄ yr⁻¹ vs SDGVM: 183 Tg CH₄ yr⁻¹, or IAP-RAS: 154.2 Tg CH₄ yr⁻¹ vs LPJ-Bern: 156.6 Tg CH₄ yr⁻¹ vs LPJ-WSL: 151.5 Tg CH₄ yr⁻¹) (Fig. 8). The comparability of these simulated global CH₄ emissions emphasizes the fact that most models are tuned to some degree towards a global total CH₄ emissions value, which allows the MPAs to vary more between the models than global CH₄ emissions. As highlighted in Melton et al. (2013), the fact that the models agree fairly well on global CH₄ emissions with very different MPAs and CH₄ fluxes underlines the importance of regional-scale observational estimates to constrain this dichotomy.
Table 6. Explanation of variable names used in Table 5.

<table>
<thead>
<tr>
<th>Variable name</th>
<th>Description</th>
</tr>
</thead>
<tbody>
<tr>
<td>$C_{labile}$</td>
<td>labile carbon pool</td>
</tr>
<tr>
<td>$O_{air}$</td>
<td>soil oxidation of atmospheric CH$_4$, i.e. CH$_4$ uptake</td>
</tr>
<tr>
<td>$O_{air,max}$</td>
<td>maximum soil oxidation of atmospheric CH$_4$ rate</td>
</tr>
<tr>
<td>$O_{soil}$</td>
<td>oxidation in the soil pore water</td>
</tr>
<tr>
<td>$O_{soil,max}$</td>
<td>maximum oxidation rate in the soil pore water</td>
</tr>
<tr>
<td>$O_{trans}$</td>
<td>oxidation associated with transport through plants</td>
</tr>
<tr>
<td>$O_{trans,max}$</td>
<td>maximum oxidation associated with transport through plants</td>
</tr>
<tr>
<td>$O_{max}$</td>
<td>maximum oxidation rate</td>
</tr>
<tr>
<td>$P_{max}$</td>
<td>maximum CH$_4$ production</td>
</tr>
<tr>
<td>$Q_{10}$</td>
<td>factor describing dependence on temperature</td>
</tr>
<tr>
<td>$R_{hetr}$</td>
<td>heterotrophic respiration</td>
</tr>
<tr>
<td>$R_0$</td>
<td>CH$_4$ production rate</td>
</tr>
<tr>
<td>$f_{CH4}$</td>
<td>function of pore water CH$_4$ concentration</td>
</tr>
<tr>
<td>$f_{atm[CH4]}$</td>
<td>function of atmospheric CH$_4$ concentration</td>
</tr>
<tr>
<td>$f_{ecosys}$</td>
<td>function of ecosystem type</td>
</tr>
<tr>
<td>$f_{GPP}$</td>
<td>function of the ratio of monthly to annual net primary production (NPP)</td>
</tr>
<tr>
<td>$f_{O2}$</td>
<td>function of pore water O$_2$ concentration, determined by rate of O$_2$ diffusion through soil water and aerenchyma</td>
</tr>
<tr>
<td>$f_{E}$</td>
<td>function of alternative electron acceptors</td>
</tr>
<tr>
<td>$f_{pH}$</td>
<td>function of pH value</td>
</tr>
<tr>
<td>$f_{plantrans}$</td>
<td>function of plant-mediated CH$_4$ transport</td>
</tr>
<tr>
<td>$f_{root}$</td>
<td>function of vertical root distribution</td>
</tr>
<tr>
<td>$f_{soil}$</td>
<td>function of soil type</td>
</tr>
<tr>
<td>$f_T$</td>
<td>function of temperature</td>
</tr>
<tr>
<td>$f_{e}$</td>
<td>function of soil moisture</td>
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<tr>
<td>$f_{transport}$</td>
<td>function of CH$_4$ transport</td>
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<tr>
<td>$f_{WTP}$</td>
<td>function of water table position</td>
</tr>
<tr>
<td>$r_{CH4:C}$</td>
<td>fraction of C converted to CH$_4$</td>
</tr>
<tr>
<td>$r_{O2}$</td>
<td>fraction of O$_2$ used for CH$_4$ oxidation</td>
</tr>
</tbody>
</table>

5 Summary and conclusions

WETCHIMP is the first multi-model comparison of wetland extent and wetland CH$_4$ emissions. Our analysis demonstrates how diverse modelling approaches, wetland definitions, and wetland extent can be, while still leading to comparable values of global CH$_4$ emissions. In terms of modelling CH$_4$ producing areas (MPAs), there are three main approaches: (i) the fixed MPA (IAP-RAS, LPJ-Bern (peatlands and wetlands), LPJ-WHyMe), (ii) the guided MPA (CLM4Me, DLEM, LPJ-WSL, ORCHIDEE, UW-VIC) and (iii) the fully simulated MPA (UVic-ESCM, SDGVM, LPJ-Bern wetsoils). Total MPA can vary significantly between models depending on their definitions, which also influences CH$_4$ fluxes, but does not have as much impact on the global CH$_4$ emissions. Achieving similar global CH$_4$ emissions with very different MPA distributions also means that the CH$_4$ fluxes between the models differ greatly. A wide range of parameterization complexity is used to simulate CH$_4$ fluxes in the participating models, which influences not just the present day flux but also its sensitivity to climate factors.

Each individual model’s approach needs to be considered carefully when interpreting results, especially past and future climate change experiments or sensitivity experiments such as those that were conducted as part of WETCHIMP (Melson et al., 2013). There are several factors that need to be kept in mind: a fixed wetland distribution as used by some models or a seasonally varying distribution based on presently observed patterns is unlikely to be representative of past or future conditions. Another limitation is the absence of wetland specific PFTs in most models. Models that lack wetland specific PFTs (i.e. CLM4Me, LPJ-WSL, SDGVM, LPJ-Bern non-peatland, ORCHIDEE, IAP-RAS) may overestimate NPP due to an unrealistic lack of plant stress that would be caused by inundation or nutrient limitation. We expect these models to show different responses to changes in temperature, precipitation and CO$_2$ fertilization than the models that include wetland specific PFTs (i.e. DLEM, LPJ-WhYMe, LPJ-Bern peatlands, UW-VIC). For example, changes in precipitation will affect wetland specific PFTs that grow under inundated conditions differently than upland plants. Also, the effect of CO$_2$ fertilization on wetland plants is still unclear (Berendse et al., 2001; Heijmans et al., 2001, 2002a,b; Boardman et al.,
2011) and therefore wetland NPP under CO₂ fertilization calculated by models that include wetland specific PFTs remains highly uncertain.

There are features that are still missing, or are crudely represented, in almost all of the models, partially due to the difficulties of simulating small-scale processes in large-scale models. Such features include (i) lateral transport of water and groundwater dynamics within (beyond the assumptions of the TOPMODEL formulation) and between grid cells, and explicit treatments of floodplains and mangroves; (ii) plant nutrients (nitrogen, phosphorus and sulphur) and their interactions (presently only SDGVM, DLEM, and CLM4Me include carbon–nitrogen interactions); (iii) microtopographical features such as lawns, hollows or hummocks and their impacts upon overall CH₄ dynamics; (iv) vertically resolved carbon pools and soil organic matter remineralization modelling; (v) permafrost−preserved carbon; (vi) feedbacks between peat or carbon dynamics and thermal and hydrological processes in soil; (vii) hydrology affected by thawing permafrost; (viii) wetland specific vegetation (improvements for boreal peatlands, introduction of tropical wetland PFTs); and (ix) anthropogenic disturbance (such as wetland drainage) and management (such as dams and reservoirs).

WETCHIMP provides a the first multi-model platform to explore the current knowledge, recent improvements, and necessary future developments of models simulating wetland extent and wetland CH₄ emissions. The design of future iterations of WETCHIMP will be focused on analysing and understanding the different uncertainties and sensitivities of participating models with the goal of greatly improving the performance of the models for both wetland and wetland CH₄ modelling. The simulations conducted in WETCHIMP are available (http://arve.epfl.ch/pub/wetchimp, please contact J. R. Melton for immediate access) and their use is encouraged to advance research in this area.

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