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The Numerical Age of the Upper Frasnian (Upper Devonian) Kellwasser Horizons: A New U-Pb Zircon Date from Steinbruch Schmidt (Kellerwald, Germany)

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ABSTRACT

Steinbruch Schmidt (Kellerwald, Germany) is probably the world's most famous locality that exposes the late Frasnian anoxic events (Kellwasser Crisis). A 3-cm-thick bentonite layer (Bed 36) is intercalated between the two Kellwasser horizons and can be precisely assigned to the Upper *rhenana* conodont Zone. Twenty-four U-Pb ID-TIMS (isotope dilution–thermal ionization mass spectrometry) analyses of single zircons extracted from this layer yielded 17 concordant results, which form an elongated cluster along concordia. The oldest point of this cluster is assumed to represent the lowest amount of Pb loss, and its $^{206}\text{Pb}/^{238}\text{U}$ age of 377.2 ± 1.7 Ma is regarded as the eruption age of the bentonite. This dating allows the geochronological fixing of the anoxic Kellwasser horizons, which characterize one of the most important biotic crises of the Palaeozoic era, immediately prior to the Frasnian-Famennian boundary. A biochronometric timescale of the mid-Frasnian to late Famennian interval is established based on all recently published Late Devonian U-Pb ID-TIMS ages. The age of the Frasnian-Famennian boundary can be estimated by interpolation to 376.1 ± 1.6 Ma.

Introduction

Most of our knowledge concerning the duration of changes in the geological record depends on time interpolation between relatively few isotopic ages, which are determined with different isotope systems and various analytical methods. Only a few of these ages are precise and biostratigraphically well controlled. On the basis of such ages, vastly differing in quality, comparatively imprecise and flawed timescales have been constructed in the past. Consequently, a methodically consistent data set of biostratigraphically well-constrained isotopic ages is the most important precondition for the construction of accurate timescales. High initial parent/daughter element ratios and a unique dual decay scheme make U-Pb zircon ages more precise and reliable than most other isotopic ages and thus inherently superior for timescale calibration.

The late Frasnian anoxic events (Kellwasser ho-

rizons) are closely linked to the Frasnian-Famennian mass extinction, which represents one of the five major bioevents of the Phanerozoic (Sepkoski 1982). The causes of this faunal crisis have been a matter of a long and vigorous debate (see reviews and references in Schindler 1990; Buggisch 1991; McGhee 1996; Racki and House 2002). However, no isotopic ages have been acquired so far that fix the timing and duration of these events. Previous estimates for the age of the Frasnian-Famennian boundary range from 364.0 to 376.5 Ma (Harland et al. 1990; Fordham 1992; Young 1995; Sandberg and Ziegler 1996; Tucker et al. 1998; Compston 2000). This uncertainty is due to the lack of reliable and biostratigraphically well-constrained isotopic ages in the Late Devonian. To date, the age of 376.5 Ma of Tucker et al. (1998) represents the best estimate for the age of the Frasnian-Famennian boundary because it was interpolated between two isotope dilution–thermal ionization mass spectrometry (U-Pb ID-TIMS) zircon ages of the mid-Frasnian (Belpre Ash, 381.1 ± 1.3 Ma) and the late Famennian (Piskahegan Group, 363.6 ± 1.6 Ma).

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Locality

Steinbruch Schmidt is located at the eastern margin of the Rhenish Massif near the village of Bad Wildungen, Germany. It is one of the best investigated outcrops of the Kellwasser lithologies (Devleeschouwer et al. 2002 and references therein). Bed 36 bentonite is 3 cm thick and situated 0.5 m above the top of the Lower Kellwasser horizon and 2.7 m below the top of the Lower Kellwasser horizon and 2.7 m below the Frasnian-Famennian boundary (Schindler 1990; fig. 1). Steinbruch Schmidt is also a reference section for conodont zonation, spanning the interval from the Lower *rhenana* to the Middle *triangularis* Zone (Ziegler and Sandberg 1990). Bed 36 bentonite can be fixed accurately in the middle part of the Upper *rhenana* Zone (fig. 1).

Analytical Methods

U-Pb zircon ages were acquired at the Zentrallabor für Geochronologie, Institut für Mineralogie, Universität Münster, Germany. Zircons were separated from Bed 36 bentonite using refinements of the procedure outlined by Winter (1981). Grains were then handpicked and sorted according to color, quality, and morphology. Zircons selected for analysis were inclusion and crack free, normal to long prismatic (2 : 1 to 6 : 1), clear, and colorless. Following air abrasion (Krogh 1982), a few samples were additionally leached for 2 hr in concentrated HF-HNO₃ (4 : 1) at 80°C. After washing in 7N HNO₃ at 80°C for 25 min, individual grains were placed in multi-sample Teflon microcapsules and dissolved for 4 d in concentrated HF-HNO₃ (4 : 1) at 180°C. Subsequently, dissolved zircons were spiked with a mixed ²³³U-²⁰⁵Pb tracer solution, dried at 80°C, re-dissolved in 6N HCl, and equilibrated at 180°C for 1 d. After drying at 80°C, the samples were loaded on a single Re filament using a mixture of silica gel and 6N HCl-0.25N H₃PO₄. Isotope ratios of Pb and U were measured with a Daly-type detector in ion-counting mode on a VG Sector 54 thermal ionization mass spectrometer (TIMS). Low radiogenic Pb concentrations required isotope analyses by peak jumping. Pb and U (as UO₂⁺) were run sequentially on the same filament at temperatures of 1250–1350°C and 1300–1450°C, respectively. Average measured precision (2σ error) of Pb isotope ratios was 0.8% (²⁰⁶Pb/²⁰⁴Pb), 0.18% (²⁰⁷Pb/²⁰⁶Pb), and 0.06% (²⁰⁶Pb/²⁰⁵Pb). Ratios were corrected individually for mass fractionation by 1.4 ± 0.8‰/amu (²⁰⁷Pb/²⁰⁶Pb, ²⁰⁶Pb/²⁰⁵Pb), and 1.5 ± 0.8‰/amu (²⁰⁶Pb/²⁰⁴Pb), based on multiple analyses of NBS 982 standard. ²³³U/²³⁸U ratio was measured more pre-

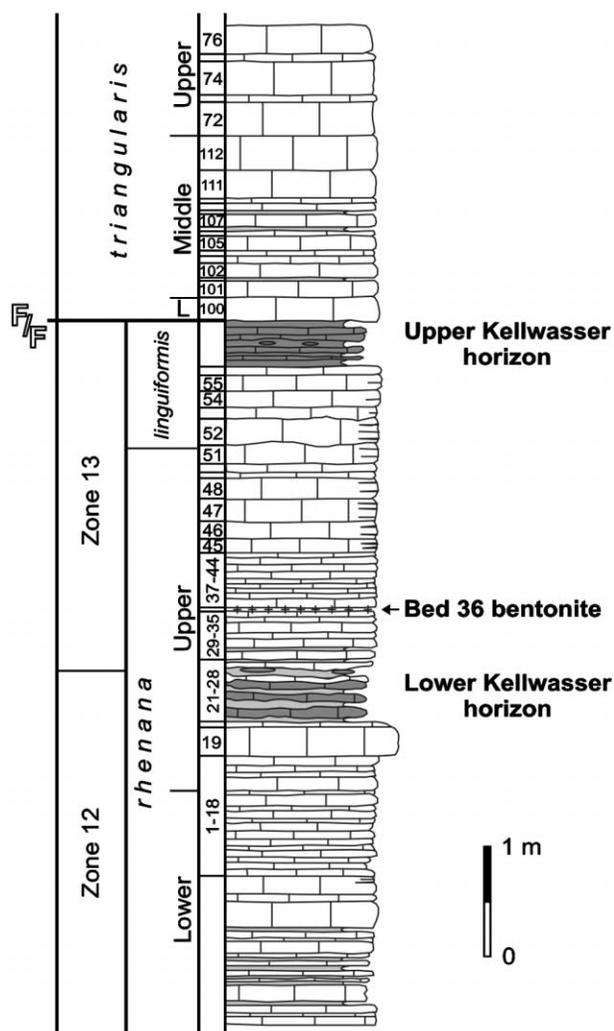


Figure 1. Lithologic column of Steinbruch Schmidt including conodont stratigraphy and position of Bed 36 bentonite (redrawn from Schindler 1990). Conodont stratigraphy based on Ziegler and Sandberg (1990).

cisely than 0.13% (2σ), and mass fractionation was corrected by 2.0 ± 0.8‰/amu based on analyses of NBS 500 standard. For each charge of samples, the maximum Pb blank was assumed to be equivalent to the amount of the common Pb in the most radiogenic sample. Small sample volumes and low concentrations of U were responsible for low radiogenic Pb amounts of 4.8 to 41.6 pg. Therefore, the highest possible reduction of the Pb blank was necessary. Due to extremely careful sample handling, the blank could be decreased in the course of the study from 3 to 1 pg. In all samples ²⁰⁶Pb_{rad}/²⁰⁶Pb_{tot} is sufficiently high to obtain precise ²⁰⁶Pb/

^{238}U ages. The U blank was too small to be measured and was thus assumed to amount to 20% of the individual Pb blank based on experience with the analysis of milligram-sized samples. Uncertainties in $^{206}\text{Pb}/^{238}\text{U}$, $^{207}\text{Pb}/^{235}\text{U}$, and $^{207}\text{Pb}/^{206}\text{Pb}$ ages were calculated using the algorithm of Ludwig (1980), and errors are quoted at the 2σ level. Software Isoplot/Ex, rev. 2.49 (Ludwig 2001) was used for the concordia plot.

Results

The zircon population of Bed 36 bentonite shows an astonishing homogeneity with euhedral, normal-to-long prismatic shapes (fig. 2A, 2B) and concentric growth zoning (fig. 2C, 2D) as expected for magmatic zircons. Inherited zircons have not been detected by cathodoluminescence investigations, but two (samples 1.3 and 2.8) of 24 analyzed

grains must have contained older cores as documented by their significant older ages (table 1). Concerning the effects of Pb loss, a slight but statistically significant trend toward concordance was observed in samples that were subjected to stronger air abrasion and additional HF leaching.

$^{206}\text{Pb}/^{238}\text{U}$ ages were determined with a typical precision of 0.4%. However, due to small sample weights and low concentrations of ^{207}Pb (at this age more than 18 times lower than ^{206}Pb), errors of the $^{207}\text{Pb}/^{235}\text{U}$ (mean 1.2%) and $^{207}\text{Pb}/^{206}\text{Pb}$ (mean 5.8%) ages are much higher (table 1). Seventeen of 24 analyses are concordant and form an elongated cluster on the concordia plot (fig. 3). $^{206}\text{Pb}/^{238}\text{U}$ ages of these points range from 359.2 to 377.2 Ma (table 1), scattering significantly outside their analytical error. This indicates that at least some of the zircons are affected either by inheritance or Pb loss. Inheritance can most likely be excluded (see above), and the two conspicuously older analyses are clearly resolved from the cluster. Therefore, and because

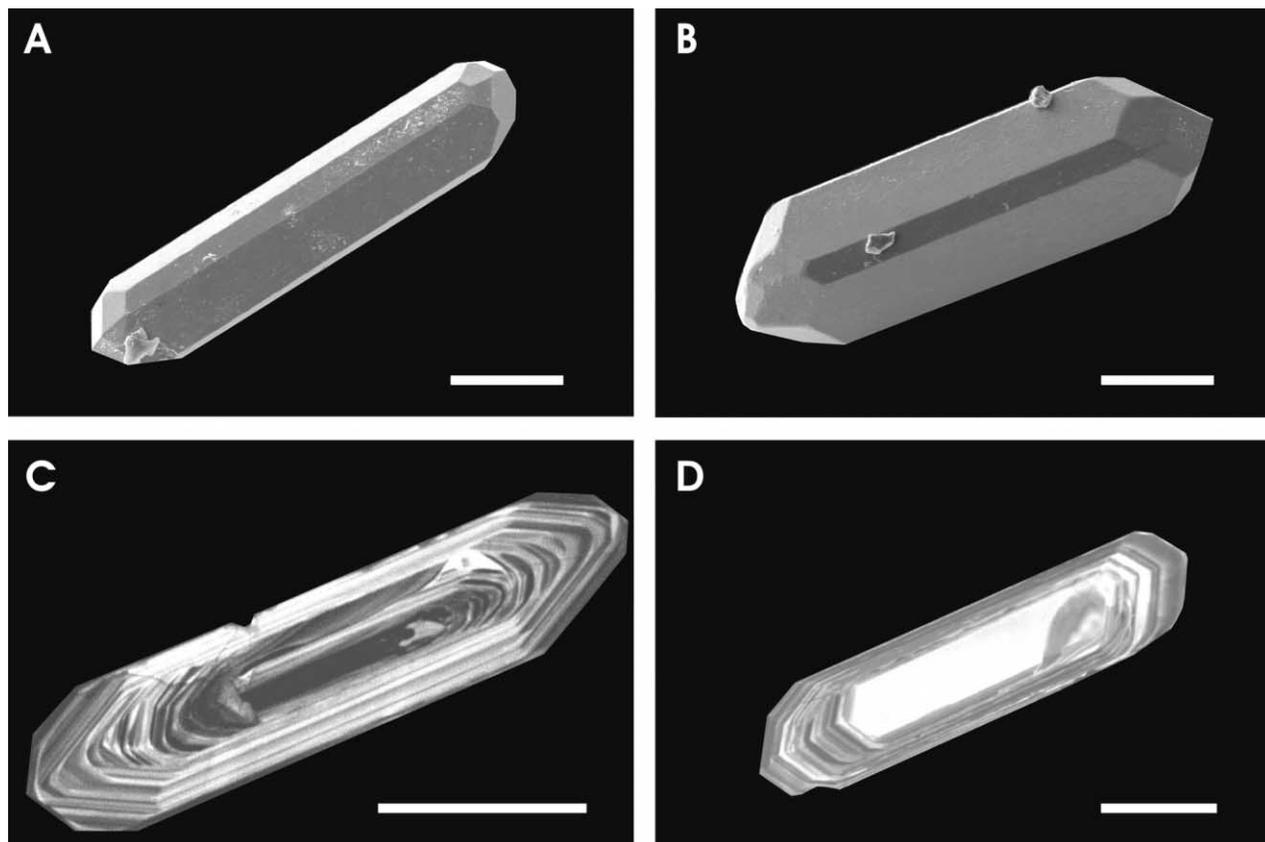


Figure 2. Scanning electron microscope and cathodoluminescence photographs of typical zircons from Bed 36 bentonite. Scale bar is 50 μm . A, Long prismatic zircon with $\{100\} > \{110\}$ and $\{101\} > \{211\}$. B, Normal prismatic zircon with $\{110\} > \{100\}$ and $\{101\} = \{211\}$. C, D, Zircons showing concentric growth zoning.

Table 1. U-Pb Zircon Data of Steinbruch Schmidt Bed 36 Bentonite

Sample ^a	Concentration				Isotopic ratio				Isotopic age (Ma)		
	Weight (μg) ^b	U (ppm) ^c	Pb rad. (pg) ^c	Pb com. (pg) ^d	²⁰⁶ Pb/ ²⁰⁴ Pb ^e	²⁰⁶ Pb/ ²³⁸ U (\pm %) ^f	²⁰⁷ Pb/ ²³⁵ U (\pm %) ^f	²⁰⁷ Pb/ ²⁰⁶ Pb (\pm) ^f	²⁰⁶ Pb/ ²³⁸ U ^f	²⁰⁷ Pb/ ²³⁵ U ^f	²⁰⁷ Pb/ ²⁰⁶ Pb ^f
1.0. 180 μm , 3.5 : 1, -a ^g	1.7	222	12.2	.7	372	.05839 (.4)	.43742 (.8)	.05433 (34)	365.8 \pm 1.4	368.4 \pm 2.9	385 \pm 14
1.0. 180 μm , 4 : 1, -a	1.3	132	6.8	.3	207	.05906 (.5)	.44159 (1.7)	.05423 (78)	369.9 \pm 1.7	371.4 \pm 6.2	380 \pm 33
1.3. 150 μm , 3 : 1, -a ^h	1.4	255	18.2	6.1	181	.06713 (.4)	.53172 (1.1)	.05745 (53)	418.8 \pm 1.6	432.9 \pm 4.7	509 \pm 20
1.4. 160 μm , 5 : 1, -a	.8	439	24.6	1.2	314	.05906 (.4)	.44226 (1.1)	.05431 (52)	369.9 \pm 1.4	371.8 \pm 4.1	384 \pm 21
1.5. 230 μm , 6 : 1, -a	.9	308	17.1	2.0	213	.05775 (.4)	.43231 (1.1)	.05430 (51)	361.9 \pm 1.6	364.8 \pm 4.1	383 \pm 21
1.6. 150 μm , 3 : 1, -a	1.4	218	11.6	<.1	361	.05868 (.4)	.43726 (.9)	.05405 (40)	367.6 \pm 1.4	368.3 \pm 3.4	373 \pm 17
1.7. 230 μm , 3 : 1, -a ^g	2.1	156	10.4	8.7	120	.05807 (.5)	.43926 (1.8)	.05486 (87)	363.9 \pm 1.6	369.7 \pm 6.6	406 \pm 36
2.3. 160 μm , 2.5 : 1, +a	1.7	155	8.5	<.1	437	.05896 (.4)	.43845 (.9)	.05393 (42)	369.3 \pm 1.5	369.2 \pm 3.5	368 \pm 18
2.8. 210 μm , 5 : 1, +a ^h	2.0	113	7.4	<.1	445	.07037 (.4)	.54361 (.8)	.05603 (34)	438.4 \pm 1.8	440.8 \pm 3.5	454 \pm 13
3.1. 150 μm , 3 : 1, +a, hf ^g	.7	704	40.3	.9	674	.05840 (.4)	.43712 (.6)	.05428 (21)	365.9 \pm 1.5	368.2 \pm 2.1	383 \pm 8
3.3. 210 μm , 4 : 1, +a, hf	1.8	89	4.8	.4	301	.05851 (.4)	.44102 (1.4)	.05467 (66)	366.5 \pm 1.5	371.0 \pm 5.1	399 \pm 28
3.4. 150 μm , 3 : 1, +a, hf	.8	211	11.5	<.1	372	.05963 (.4)	.44608 (1.0)	.05425 (44)	373.4 \pm 1.5	374.5 \pm 3.7	382 \pm 18
3.5. 160 μm , 2.5 : 1, +a, hf ^g	1.7	318	18.0	6.7	228	.05214 (.4)	.39411 (1.8)	.05482 (93)	327.6 \pm 1.3	337.4 \pm 6.0	405 \pm 37
3.6. 150 μm , 3 : 1, +a, hf	1.0	273	18.3	6.7	140	.05812 (.6)	.43340 (2.3)	.05408 (113)	364.2 \pm 2.1	365.6 \pm 8.5	375 \pm 48
3.8. 150 μm , 3 : 1, +a, hf	.9	462	29.5	6.7	198	.05805 (.4)	.43591 (1.3)	.05446 (63)	363.8 \pm 1.4	367.4 \pm 4.7	390 \pm 26
4.1. 150 μm , 3 : 1, +a, hf	.8	546	35.8	6.7	218	.06006 (.4)	.44860 (1.6)	.05417 (81)	376.0 \pm 1.5	376.3 \pm 6.0	378 \pm 33
4.3. 150 μm , 3 : 1, +a, hf	.9	301	16.5	<.1	438	.05816 (.4)	.43189 (.7)	.05386 (28)	364.4 \pm 1.3	364.5 \pm 2.5	365 \pm 12
4.4. 180 μm , 3.5 : 1, +a, hf	1.5	173	9.7	<.1	580	.05931 (.4)	.44270 (.7)	.05413 (30)	371.4 \pm 1.4	372.2 \pm 2.7	377 \pm 12
4.5. 150 μm , 3 : 1, +a, hf	1.0	130	7.2	.3	288	.06001 (.4)	.44803 (1.3)	.05415 (58)	375.7 \pm 1.6	375.9 \pm 4.7	377 \pm 24
5.0. 210 μm , 5 : 1, +a, hf	2.1	148	8.9	1.4	518	.05960 (.4)	.44247 (.8)	.05385 (35)	373.2 \pm 1.4	372.0 \pm 2.9	365 \pm 14
5.1. 160 μm , 2.5 : 1, +a, hf	1.5	123	7.0	<.1	716	.05970 (.4)	.44614 (.8)	.05420 (35)	373.8 \pm 1.4	374.6 \pm 3.0	379 \pm 15
5.2. 180 μm , 3.5 : 1, +a, hf	1.8	109	6.3	.3	639	.05966 (.4)	.44650 (.7)	.05427 (32)	373.6 \pm 1.4	374.8 \pm 2.8	382 \pm 13
5.4. 150 μm , 3 : 1, +a, hf	.8	192	11.2	.5	430	.05949 (.4)	.44426 (1.0)	.05416 (48)	372.5 \pm 1.4	373.3 \pm 3.8	378 \pm 20
5.5. 210 μm , 4 : 1, +a, hf ^g	1.7	296	17.6	3.9	394	.05681 (.4)	.42635 (.8)	.05443 (37)	356.2 \pm 1.4	360.6 \pm 3.0	389 \pm 15

^a 200 μm = length of grain; 3 : 1 = length : width ratio; -a = slightly abraded (~10% reduction of crystal volume); +a = strongly abraded (~30% reduction); hf = HF leached.

^b Estimated (\pm 30%) from crystal length and shape.

^c Calculations based on estimated weight.

^d Corrected for Pb blank and spike contribution.

^e Measured value corrected for mass discrimination and fractionation.

^f Ratios and ages corrected for mass discrimination and fractionation, spike, Pb blank, and initial common Pb (at 376 Ma from Stacey and Kramers 1975).

^g Denotes discordant zircon analysis caused by significant recent Pb loss; not displayed in the concordia plot.

^h Denotes zircon analyses with isotopic evidence for inheritance; not displayed in the concordia plot.

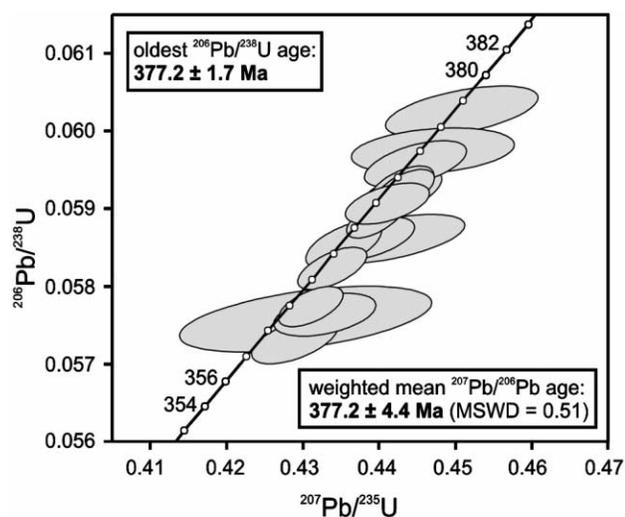


Figure 3. Concordia plot of 17 concordant single-zircon analyses of Bed 36 bentonite.

individual error ellipses are more or less homogeneously distributed within the cluster, the simplest explanation of the age scatter is Pb loss. If this interpretation is correct, the oldest point represents the lowest amount of Pb loss, and its $^{206}\text{Pb}/^{238}\text{U}$ age of 377.2 ± 1.7 Ma can be regarded as the time of eruption of the bentonite. This age is confirmed by the identical weighted mean $^{207}\text{Pb}/^{206}\text{Pb}$ age (377.2 ± 4.4 Ma, MSWD = 0.51) based on all concordant points.

Timescale Implications

The age of 377.2 ± 1.7 Ma of Bed 36 bentonite fits well into the timescale of Tucker et al. (1998), which is also based on U-Pb ID-TIMS zircon ages (fig. 4). It is bracketed by the mid-Frasnian Belpre Ash (381.1 ± 1.3 Ma) and the late Famennian Piskahegan Group (363.6 ± 1.6 Ma). The age of the latter has meanwhile been confirmed by a U-Pb monazite age obtained from the Nordegg Tuff of the late Famennian Exshaw Shale in southwestern Canada (Richards et al. 2002). If all these Late Devonian U-Pb ID-TIMS ages are treated equally, a recalibration and more precise rendering of the mid-Frasnian to late Famennian interval is possible. For this purpose, approximately time-linear biostratigraphic scales are used here for interpolation between the isotopic ages. The biostratigraphic scales were obtained from conodont-bearing sec-

tions with homogeneous lithologies that reflect more or less uniform rock accumulation rates. These in turn are the precondition for an approximately linear record of time. A proportionate presentation of biozonal lengths as shown here is much more reliable than bioschemes based on the assumption that biozones represent equal time intervals. For interpolation, the biostratigraphic scales have now been spanned between each two successive isotopic ages: at first, between the ages of the Belpre Ash and Steinbruch Schmidt and subsequently between the Steinbruch Schmidt age and the coeval Piskahegan Group/Exshaw Formation ages (fig. 5). The result is a biochronometric timescale that allows the assignment of a numeric age with its error to each biozone boundary in the mid-Frasnian to late Famennian interval (fig. 5). The numerical age of the Frasnian-Famennian boundary is estimated by interpolation to 376.1 ± 1.6 Ma. We regard a biochronometric timescale constructed in this manner as the closest approach to the “true” timescale that can be achieved based on the currently available data.

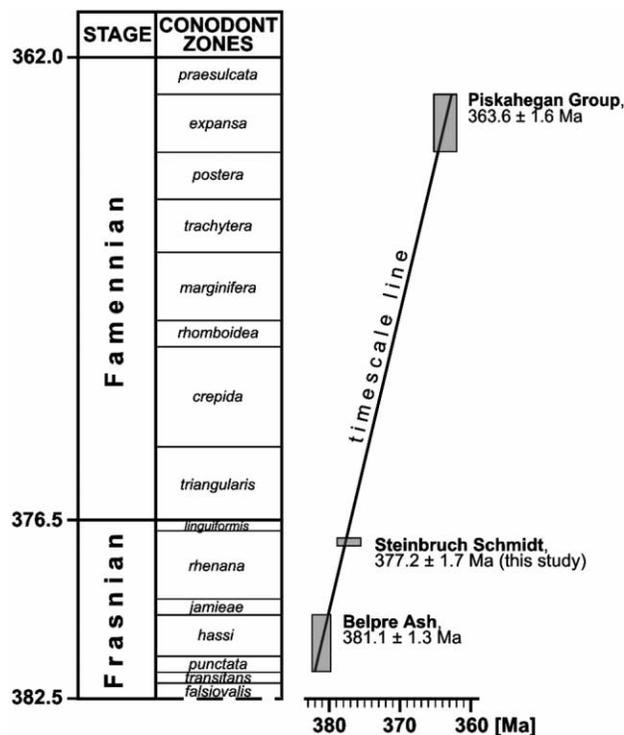


Figure 4. Plot of the U-Pb zircon age of Steinbruch Schmidt into the timescale (redrawn) of Tucker et al. (1998).

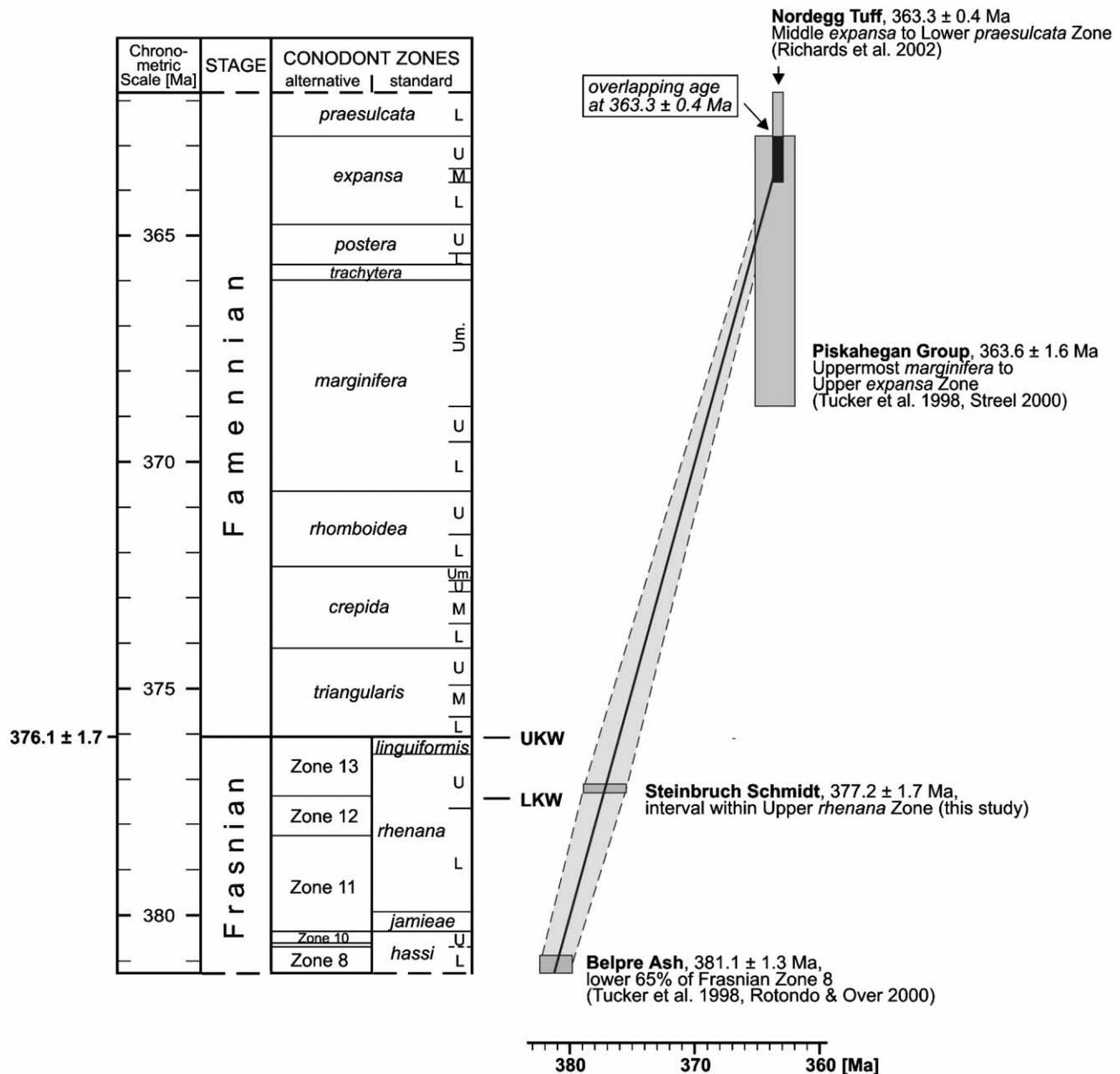


Figure 5. Late Frasnian to late Famennian biochronometric timescale. The biostratigraphic scale was obtained from the conodont sections at the Montagne Noire in France (*hassi* to *crepida* Zone; Klapper 1997) and at Lali in China (*crepida* to *praesulcata* Zone; Ji and Ziegler 1993). Correlation of standard with alternative conodont zones of the Frasnian is based on Klapper and Becker (1999). LKW = Lower Kellwasser horizon, UKW = Upper Kellwasser horizon.

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