Bacterial diversity and biogeochemistry of different chemosynthetic habitats of the REGAB cold seep (West African margin, 3160 m water depth)

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Abstract

The giant pockmark REGAB (West African margin, 3160 m water depth) is an active methane-emitting cold seep ecosystem, where the energy derived from microbially mediated oxidation of methane supports high biomass and diversity of chemosynthetic communities. Bare sediments interspersed with heterogeneous chemosynthetic assemblages of mytilid mussels, vesicomyid clams and siboglinid tubeworms form a complex seep ecosystem. To better understand if benthic bacterial communities reflect the patchy distribution of chemosynthetic fauna, all major chemosynthetic habitats at REGAB were investigated using an interdisciplinary approach combining porewater geochemistry, in situ quantification of fluxes and consumption of methane, as well bacterial community fingerprinting. This study revealed that sediments populated by different fauna assemblages show distinct biogeochemical activities and are associated with distinct sediment bacterial communities. The methane consumption and methane effluxes ranged over one to two orders of magnitude across habitats, and reached highest values at the mussel habitat, which hosted a different bacterial community compared to the other habitats. Clam assemblages had a profound impact on the sediment geochemistry, but less so on the bacterial community structure. Moreover, all clam assemblages at REGAB were restricted to sediments characterized by complete methane consumption in the seafloor, and intermediate biogeochemical activity. Overall, variations in the sediment geochemistry were reflected in the distribution of both fauna and microbial communities; and were mostly determined by methane flux.

1 Introduction

Cold seeps belong to the most productive ocean ecosystems, and are hence known as oases of life in the deep sea. They are fuelled by energy provided via microbial transformation of methane and higher hydrocarbons, which supports large biomasses of highly specialized chemosynthetic communities (Sibuet and Olu, 1998; Jørgensen...
Cold seep ecosystems are found on passive and active continental margins around the world’s oceans, but represent fragmented, isolated habitats of locally small areal coverage, similar to hydrothermal vents. Ever since the discovery of chemosynthetic communities at cold seep ecosystems (Paull et al., 1984; Kennicutt II et al., 1985; Suess et al., 1985) ecologists have been fascinated by the question of their interconnectivity and biogeography on global and regional spatial scales, just as for hydrothermal vent communities (Tunnicliffe, 1991; Sibuet and Olu, 1998; Tunnicliffe et al., 1998; Tyler et al., 2003; Vanreusel et al., 2009). The most conspicuous characteristics of cold seep ecosystems are seafloor emissions of hydrocarbon fluids and gases as primary, locally restricted energy source causing a specific adaptation of species to their habitat. Furthermore, cold seeps show a fragmented distribution along the global continental margins, leading to isolation of populations. Hence, studies dedicated to the geobiology of cold seeps have been seeking to answer the following overarching questions (i) how diverse and specific are microbial and faunal communities at cold seeps compared to the background environment (ii) to what extent are seep communities indicative of geological processes such as the strength of hydrocarbon seepage, and (iii) which factors control the diversity and dispersal rates of seep – associated organisms on both regional and global scales (among and within cold seeps).

These questions were central to this study of the bacterial communities of the giant pockmark cold seep REGAB (Charlou et al., 2004). Studies investigating seep microorganism have predominantly been focusing on the taxonomical identification of the key-microbial players mediating the core-energy producing processes i.e. Anaerobic Oxidation of Methane (AOM) coupled to Sulphate Reduction (SR) and sulphide oxidation, their phylogenetic affiliation, and relation of their abundances and activities to the local geochemistry (Boetius et al., 2000; Knittel et al., 2003; Heijs et al., 2007; Cambon-Bonavita et al., 2009; Knittel and Boetius, 2010; Orcutt et al., 2010). However, high-resolution studies investigating the structure of microbial communities as a whole, over a range of habitats within a single cold seep (m to km scale) are still rare, hence...
the abiotic and biotic factors that may directly influence the bacterial diversity at cold seeps remain elusive.

Here we tested the hypotheses that the bacterial diversity at cold seeps changes along geochemical gradients with sediment depth, and moreover that the availabilities of methane and sulphide as main energy sources at cold seeps structure the bacterial communities. Secondly, similar to what has been shown for small-size fauna (Van Gaever et al., 2009), we tested if the type and specific activity of chemosynthetic megafauna can exert a selective pressure and therefore influence bacterial diversity patterns at cold seeps. Therefore we combined porewater geochemistry, quantification of fluxes and consumption rates of methane and bacterial community fingerprinting of the major habitat types at REGAB, populated by different chemosynthetic organisms (mytilid mussels, vesicomyid clams or thiotrophic bacterial mats). The main aims were to better understand (i) if the bacterial community within a cold seep ecosystem varies according to habitat distribution on a “landscape” scale (ii) if the bacterial diversity is shaped by availability of energy, and (iii) if habitat partitioning by large symbiotic megafauna influences the structure of bacterial communities.

2 Material and methods

2.1 Description of sampling sites and sampling procedure

The REGAB pockmark, situated at 3160 m water depth on the Congo-Angola margin represents the second largest (800 m Ø and 15–20 m deep) single cold seep site known to date in the Eastern Atlantic (Sibuet and Olu-Le Roy, 2002; Ondrédas et al., 2005). REGAB has unusual sedimentological features resulting from the large amount of terrigenous input by the Congo River and the vicinity of one of the largest submarine canyons, the Congo deep-sea fan (Ondrédas et al., 2005 and references therein; Pierre and Fouquet, 2007). The pockmark formation has been related to a sudden release of overpressurized gas followed by a collapse of a sediment dome (Ondrédas
et al., 2005). The current escape of gas is assumed to occur along a deep pipe (300 m) rooted in a buried paleo-channel (Gay et al., 2003; Ondréas et al., 2005). The seafloor of the REGAB pockmark is characterized by high biomasses of diverse chemosynthetic megafauna which occur clustered as non-overlapping aggregations of mytilid mussels or vescomyid clams, or siboglinid tubeworms (Ondréas et al., 2005; Olu-Le Roy et al., 2007a).

Video surveys during this study in 2008 (METEOR 76/3b report http://www.dfg-ozean.de/de/berichte/fs_meteor/) revealed that the REGAB pockmark was dominated by the same three major types of megafauna assemblages already discovered during previous investigation of this cold seep in 1998, 2000 and 2001 (Ondréas et al., 2005; Sibuet and Vangriesheim, 2009; Fig. 2). Extensive carbonate cements and outcropping gas hydrates mark the central part of REGAB, where also strong venting of gas and surface precipitates of gas hydrate were observed. This area was densely populated by mussels and tubeworms attached to the carbonates, and was surrounded by soft sediments littered with clam shells, and patchily distributed beds of living clams. In comparison to the areal coverage of fauna estimated in 2000 (Olu-Le Roy et al., 2007a), the number and sizes of mussel patches associated with soft sediments seemed to have declined with time, as deduced from the Remotely Operated Vehicle (ROV)-video observations performed in 2008. However, the spatial distribution of clam species forming heterogeneous patches (Fig. 2) has remained more or less the same, indicating that relatively stable biogeochemical conditions prevail over time at REGAB.

Sediment sampling, in situ geochemical measurement and ROV QUEST (MARUM, Bremen, Germany) video observations at REGAB were performed during the M76/3b cruise in 2008, aboard the R/V Meteor. The eleven investigated sites cluster in three geographical regions of the pockmark (north, south and south-west) and included: 1 bare sediment site where venting of gas bubbles was observed (Gas), 1 bacterial mat (Bacter_N), 3 clam (Clam_N, Clam_S, Clam_SW) and 1 mussel patch (Mussel_S), as well as 6 associated sites devoid of symbiotic megafauna characterized by bare sediment (Bacter_N_Env, Clam_N_Env, Clam_S_Env, Mussel_S_Env, Clam_SW_Env). Push
coring and biogeochemical measurements were not possible in the central carbonate-hydrate site, due to the hard substrate at this location. The distances between the megafaunal patches and the adjacent barren sediment were in the range of 3–28 m. Approximate distances between the main sampling locations were: 100–150 m between N REGAB – S REGAB, 400–450 m between S REGAB and SW REGAB and 350–400 m between N REGAB and SW REGAB (Fig. 1).

In the northern part of REGAB (REGAB_N; Fig. 2) two main types of habitats, white thiotrophic bacterial mats (not found in the other parts of REGAB) and clam patches were sampled. The two investigated sites (Bacter_N and Clam_N) were found <36 m apart. The Clam_N patch was the largest of all investigated patches and was dominated by one clam species – *Calyptogena regab* (von Cosel and Olu, 2009). The seafloor at REGAB_N was littered with empty and broken shells of clams. Intermingled among the clam shells, high number of holothurians could be observed. Additional sampling of the bare sediments surrounding the clam patch (Clam_N_Env) and the bacterial mat (Bacter_N_Env) was also performed.

Two distinct types of megafauna assemblages dominated by either mussels or clams were detected in the southern part of REGAB (REGAB_S; Fig. 2). The different megafauna patches were found close to each other (20–30 m), however in no occasion they overlapped. Uniquely, at this site (Mussel_S) the mussels were surrounded by soft sediment, thus sampling was possible. *Bathymodiolus* sp. aff. *boomerang* (Olu-Le Roy et al., 2007a) was the mussel patch-forming species. Within the mussel patch individuals of the siboglinid polychaetes – *Escarpia southwardae* (Andersen et al., 2004) were visible, as well as numerous shrimps and occasionally galatheid crabs. The investigated clam patch (Clam_S) consisted of *Calyptogena regab* (von Cosel and Olu, 2009) species buried in dark sediment. No shrimps were observed inhabiting the clams’ assemblages. The larger area of the surrounding sediment at the mussel patch was overlain with shell and tubeworm debris compared to the sediment in the vicinity of the clam patch. Two additional sites, located in the vicinity of the respective megafauna assemblages (Clam_S_Env and Mussel_S_Env) were investigated during this study.
Clams were the only chemosynthetic organisms forming assemblages in the south-western part of REGAB (REGAB_SW; Fig. 2). Distinctly, the sampled patch (Clam_SW) at this part of REGAB was dominated by two species of vesicomyid clams – *Calyptogena regab* (von Cosel and Olu, 2009) and *Laubiericoncha chuni* (Thiele and Jaeckel, 1931; von Cosel and Olu, 2009) immersed in very dark sediment. No other vagrant megafauna were noticed to dwell in the Clam_SW clam site upon sampling. The sediment adjacent to the Clam_SW_Env was from one side covered by shell debris, and from the other more or less barren.

Sampling locations and sample labels are summarized in Table 1 and Supplement Table S1, and all related data are available in the PANGAEA database.

Push core (⌀ 8 cm, sediment height 10–20 cm) targeted sampling of the individual habitats was performed using the ROV QUEST camera system and the manipulator arm. Immediately after recovery, the push cores were transferred to a tempered room kept at in situ temperature (4°C) and subsampled for different type of analyses.

### 2.2 Biogeochemical measurements

#### 2.2.1 Porewater chemistry

Porewater was extracted in one centimetre resolution, from the top 0–20 cm sediment depth, using Rhizon moisture samplers (Seeberg-Elverfeldt et al., 2005; pore size 0.1 µm) inserted into holes of predrilled push core liners at all investigated sites (Table 1). The porewater was immediately subsampled for different types of analysis (total H₂S, SO₄²⁻, NH₄⁺, Fe²⁺, Mn²⁺, PO₄³⁻, Cl⁻, and Alkalinity). Description of the measurement procedure for Fe²⁺, Mn²⁺, PO₄³⁻ as well as pH can be found in the Supplement. Alkalinity was calculated from a volumetric analysis by titration of 1 ml of the porewater samples with 0.01 or 0.05 M HCl, performed on board. Concentration of ammonium was determined via the conductivity method on board (modified after Hall and Aller, 1992). Porewater subsamples (1.5 ml) were immediately fixed with 0.6 ml zinc acetate, and the total sulphide (∑H₂S = H₂S+HS⁻+S²⁻) concentration was determined.
photometrically (Cline, 1969) in the home laboratory. Sulphate and chloride subsamples were diluted 1:100 and stored frozen at −20 °C until further processing and determination of the concentration by ion chromatography (Metrohm IC Advanced Compact 861) at a flow rate of 0.7 ml min⁻¹.

To estimate the proportions of bioirrigation and advection on the transport of solutes, the transport-reaction model Explicite was applied (Zabel and Schulz, 2001; Kuester-Heins et al., 2010). For this purpose we used the porewater concentration profiles of sulphate and sulphide. Values for sulphate reduction were adjusted to the integrated rates determined by the radiotracer injection method.

2.2.2 Methane and sulphate consumption rates (AOM and SR)

Rates of AOM and SR were determined ex situ at all investigated habitats (Table 1). Immediately after recovery, the sediment push cores were subsampled on board in 3 replicates for each of the methods. Following the whole core injection method (Jørgensen, 1978), subcores (Ø 28 mm) were injected with either ¹⁴CH₄ or ³⁵SO₄ radiotracers at 1 cm intervals, according to the procedure described in Felden et al. (2010). Methane and sulphate concentrations were measured by gas chromatography (5890A, Hewlett Packard) and anion exchange chromatography (Waters IC-Pak anion exchange column, waters 430 conductivity detector) in the home laboratory, respectively. Turnover rates of methane and sulphate were determined in the home laboratory by scintillation counting according to Treude et al. (2003) and Kallmeyer et al. (2004), respectively.

2.2.3 Benthic chamber measurements (TOU and CH₄ efflux)

The Total Oxygen Uptake (TOU) and methane efflux through the sediment-water interface were determined in situ for most of the investigated habitats (Table 1) using a ROV – operated benthic chamber module (CHAM). Technical description of the benthic chamber module and details on the CH₄ measurement procedure can be found in Felden et al. (2010) and Duperron et al. (2011). Briefly, 284 cm² of sediment with
10–15 cm overlying water was incubated in situ and changes monitored over time by pre-programmed syringe sampling and optode measurements. Oxygen concentration in the enclosed bottom water was continuously measured with a help of an oxygen optode, and the TOU flux (mmol m$^{-2}$ d$^{-1}$) was calculated from the initial linear decrease in O$_2$ concentration versus time (Wenzhöfer and Glud, 2002; Felden et al., 2010). Methane concentrations were determined on water samples taken from the incubated water column, on board using a gas chromatograph (Agilent 6890N) as described previously (Niemann et al., 2009 and references therein). The methane efflux (mmol m$^{-2}$ d$^{-1}$) was calculated as the change of methane concentration in the enclosed bottom water over time (Felden et al., 2010).

2.2.4 Oxygen microsensor measurements

A modified version of a deep-sea microprofiler (MICP) was used to carry out high-resolution microsensor measurements (200 µm) for in situ determination of oxygen concentrations (Wenzhöfer et al., 2000; Lichtschlag et al., 2010a). Precise positioning and operation of the MICP was achieved with the ROV. Due to the fragile nature of microelectrodes, measurements of the oxygen inventory were restricted only to sites devoid of hard substrates and shell debris, such as the bacterial mat and the bare sediments surrounding the clam patches (Table 1). For each deployment the MICP carried 3 Clark-type O$_2$ microelectrodes with guard cathode and internal reference (Revsbech, 1989). The microsensors were calibrated by applying a two-point calibration (estimated from the O$_2$ concentration in the bottom water and in the anoxic part of the sediment). Bottom water was sampled by means of ROV–water sampler (KIPS) (Garbe-Schönberg et al., 2006) and the oxygen concentration was determined with the Winkler titration method (Grasshoff et al., 1999). Diffusive Oxygen Uptake (DOU) was calculated from the linear concentration gradient in the DBL (Diffusive Boundary Layer) by applying Fick’s first law of diffusion, as described in (Jørgensen and Des Marais, 1990). The diffusion coefficient of O$_2$ in seawater was corrected for salinity and temperature (Li and Gregory, 1974).
2.3  Bacterial community characterization

2.3.1  Bacterial cell numbers

The total number of single cells was determined at every station by applying the Acridine Orange Direct Count (AODC) method. Push cores were subsampled vertically with smaller cores (Ø 28 mm), which in turn were sliced into 2 cm sections. Samples were fixed in 4 % formaldehyde/seawater and stored at 4 °C. The AO-staining was done in the home laboratory as previously described (Meyer-Reil, 1983; Boetius and Lochte, 1996). For each sample, a minimum of two replicate filters and 30 grids per filter were randomly counted.

2.3.2  DNA samples

On board, sediment cores for DNA analysis were sliced in 2 cm intervals, down to 18 cm depth and stored at −20 °C for further analysis in the home laboratory. DNA was extracted from one gram of sediment with the Ultra Clean Soil DNA Kit (MoBio Laboratories, Inc., Carlsbad, CA, USA) following the manufacture’s recommendations for maximum yields. DNA was finally eluted in 100 µl 1 × TE buffer (Promega, Madison, WI, USA). DNA quality was confirmed with a gel electrophoresis (1 % agarose gel). Extracted DNA was quantified with a microplate spectrometer (Infinite® 200 PRO NanoQuant, TECAN Ltd, Switzerland).

2.3.3  Automated ribosomal intergenic spacer analysis (ARISA)

The bacterial community structure at REGAB was determined by Automated Ribosomal Intergenic Analysis (ARISA; Fisher and Triplett, 1999). Standardized amounts of DNA (10 µl) from each sample were amplified in triplicates using the forward FAM-labelled ITSF and reverse ITSReub primers (Cardinale et al., 2004). The PCR procedure, purification of PCR-products, capillary electrophoresis reaction, as well
as data transformation were carried out as described previously (Ramette, 2009). The ARISA peaks were binned, using a bin size of 2 bp, to account for slight peak shifts between runs and for peak size calling imprecision (Interactive Binner function, http://www.ecology-research.com, Ramette, 2009).

2.3.4 Statistical analyses

All statistical analyses were performed with the open-source software R (The R Project for Statistical Computing v.2.9.2 [http://www.r-project.org/]) using the "vegan" package (Oksanen et al., 2011). All statistical analyses were restricted to the 0–10 cm sediment depth. Correlation between differences in geochemical fluxes and β-diversity (calculated as Bray-Curtis and Jaccard distances) was tested applying the Mantel correlation test based on spearman ranking. Mantel p-values were corrected for multiple testing using the Bonferroni correction (Ramette, 2007). A reduced dataset, including the following sites: Clam_N, Clam_S, Clam_S_Env, Mussel_S, Mussel_S_Env, Clam_SW, Clam_SW_Env, and pooled samples within individual habitats (0–10 cm merged ARISA peaks) were used for this analysis. All other statistical analyses were performed on individual non-pooled ARISA samples derived from individual sediment depth samples. The bacterial community structure at REGAB was visualized by applying Non-metric MultiDimensional Scaling (NMDS) analysis. Separations of groups identified on the NMDS plot were tested for significance using the Analysis of Similarity (ANOSIM) test. Forward selection procedure was used to choose the simplest model that can explain most of the variation in the community. The effect of the geochemistry (porewater concentrations of CH_4, total H_2S, SO_4^{2-}, pH, Mn^{2+}, Fe^{2+}, Cl^-, PO_4^{3-} and alkalinity), space (calculated as geographic distances in m), sediment depth and clam presence on the bacterial community structure was assessed by canonical variation partitioning analysis, on a priori Hellinger-transformed response dataset (Legendre and Legendre, 1998; Ramette and Tiedje, 2007). Single and combined effects of factors were tested for significance by performing 999 Monte Carlo permutations. Clam presence/absence
was dummy-coded assuming that clams directly influence only the topmost 6 cm of sediment, based on their average shell length and distance reached by their foot.

Distance based test for homogeneity of multivariate dispersions (Anderson et al., 2006) was applied to test for the influence of the clams on the dispersion of bacterial β-diversity on smaller scales. The difference among dispersion groups was tested for significance using the Kruskal Wallis and Mann Whitney U test for unmatched samples. The later test was also used to assess if differences in shared OTUs among sites are statistically significant.

Geochemical data were standardized and normalized using the log-transformation prior to the analyses. Pairwise distances among samples were calculated using the Bray-Curtis dissimilarity index (Bray and Curtis, 1957) for the ARISA diversity data and Euclidean distances for geochemical data (for both, flux and concentration data) prior to the analyses mentioned above.

3 Results

3.1 Biogeochemistry of different habitats at REGAB

3.1.1 Total sulphide concentration and fluxes

Porewater concentrations of sulphide varied substantially between different habitats of REGAB, both in terms of ranges as well depth profiles (Fig. 3). The sulphide concentration increased with increasing sediment depth at most of the sampling sites. At the bacterial mat and the mussel habitat 0.1–2 mM of sulphide was detected already in the surface layers. At all investigated clam habitats, as well as the gas bubble site, sulphide was only detected below 3–9 cm sediment depth (Fig. 3). The maximum sulphide concentrations at the Muss_S and Bacter_N sites reached 9 and 11 mM, which was almost two times higher (3–6 mM) compared to the maximum concentrations measured at any of the clam-populated sites. Comparison of the sulphide fluxes revealed similar
patterns, with highest values measured at the Muss_S site, followed by slightly lower values at the Bacter_N site (Table 2). The lowest sulphide fluxes (5–9 mmol m⁻² d⁻¹) were associated with the clam patches (Clam_N, Clam_S, Clam_SW) (Table 2). The bare sediment sites surrounding the mussel, Clam_N and bacterial patches exhibited even higher ranges of sulphide fluxes, but at the Clam_SW_Env and Clam_S_Env sites no sulphide flux could be detected outside the clam bed (Fig. 3, Table 2).

### 3.1.2 Sulphate concentration and alkalinity

Sulphate concentrations decreased with depth at most of the habitats at REGAB due to the anaerobic oxidation of methane, matching well the increasing concentration of sulphide at the respective sites, as well as the increase in alkalinity (Fig. 3). At the Bacter_N and Mussel_S sites lowest concentrations of only 2–3 mM were measured at 15–16 cm sediment depth. In contrast, at all clam populated sites sulphate penetrated deep into the sediment and substantial amounts (19–24 mM) were measured up to 17 cm sediment depth. Sulphate was not completely depleted at any of the investigated sites at REGAB (Fig. 3). Most of the sulphate profiles at the bare sediment sites were similar to the ones measured within the megafauna patches. Exceptions were the Bacter_N_Env and Clam_N_Env sites, where minimum sulphate concentrations of 17 and 10 mM were measured. The sulphate depth profile at the Gas site was similar to the clam patches where sulphate concentrations decreased only after 5 cm sediment depth. pH was around 7.6 in the cores which did not show substantial degassing (see Supplement Fig. S1, also for other porewater constituents not further discussed here).

### 3.1.3 Ammonium concentrations

Ammonium was measured only at few of the investigated sites, due to restricted availability of porewater volumes (Fig. 3). At both locations (REGAB_S and REGAB_SW), NH₄ concentrations were higher inside the clam patches (Clam_S and Clam_SW) relative to the adjacent bare sediment (Clam_S_Env and Clam_SW_Env). Moreover, within
the clam patches ammonium was peaking in shallower depths (3–5 cm), while outside the patches ammonium steadily increased with depth. The ammonia content and depth distribution pattern at the Mussel_S_Env was similar to that of the bare sediment sites in the vicinity of the clam patches.

3.1.4 Sulphate consumption rates (SR)

The average integrated (0–10 cm) sulphate reduction rates measured at different habitats at REGAB confirmed the trend observed for the sulphide fluxes. The lowest rate of <1 mmol m\(^{-2}\) d\(^{-1}\) was observed at Clam_N, and the highest rate of 36 mmol m\(^{-2}\) d\(^{-1}\) at the Muss_S_Env (Table 2). Of the habitats populated by different chemosynthetic organisms, the Muss_S site exhibited highest average integrated SR rates of 28 mmol m\(^{-2}\) d\(^{-1}\), followed by slightly lower rates at the Bacter_N (23 mmol m\(^{-2}\) d\(^{-1}\)) and lowest integrated SR rates at the clam patches (<1–10 mmol m\(^{-2}\) d\(^{-1}\)). SR rate depth patterns differed substantially among habitats, ranging from surface peaks (Clam_S and Clam_SW) to deep sediment maxima (Bacter_N and Mussel_S_Env) (Fig. 3). The bare-sediment sites had lower SR rates than the adjacent fauna-populated sites, except for the conspicuous Clam_N_Env where an average integrated SR rate of 4.5 mmol m\(^{-2}\) d\(^{-1}\) was detected. Sulphate reduction rates at the Bacter_N_Env were an order of magnitude lower compared to the Bacter_N site.

3.1.5 Methane consumption rates (AOM)

The average integrated (0–10 cm) rates of AOM matched well the corresponding SR rates, and for most of the habitats a 1:1 ratio between methane oxidized and sulphate consumed could be observed as predicted by the AOM stoichiometry (Boetius et al., 2000). As already observed when integrated SR rates and sulphide fluxes were compared among fauna-populated habitats, highest integrated AOM average rates were measured at the Muss_S (19 mmol m\(^{-2}\) d\(^{-1}\)), followed by intermediate rates at the Bacter_N (9 mmol m\(^{-2}\) d\(^{-1}\)) sites and the lowest
rates at the clam patches (<1–6 mmol m$^{-2}$ d$^{-1}$) (Clam_N, Clam_S, Clam_SW) (Table 2). Highest peaks of AOM activity were measured in the surface layers (2.5 cm) at the Bacter_N site (208 nmol cm$^{-3}$ d$^{-1}$), mid-sediment peaks (4.5 cm) at the Muss_S site (465 nmol cm$^{-3}$ d$^{-1}$) and deep-sediment maxima (8.5 cm) at the Clam_SW site (67 nmol cm$^{-3}$ d$^{-1}$) (Fig. 3). At the Clam_N site the AOM rates remained low (<11 nmol cm$^{-3}$ d$^{-1}$) throughout the whole investigated sediment depth. The integrated average AOM rates of fauna-populated sites were similar to their adjacent bare sediment sites, despite the difference in the depth-distribution of AOM rates.

### 3.1.6 In situ CH$_4$ efflux

CH$_4$ effluxes were hardly detectable with only <1–3 mmol m$^{-2}$ d$^{-1}$ measured at the investigated clam habitats – both inside and outside the clam patches (Table 2). The opposite was true for the mussel-related site (Mussel_S_Env) where two orders of magnitude higher methane efflux was measured (334 mmol m$^{-2}$ d$^{-1}$). Within the mussel patch CH$_4$ efflux varied substantially – from 1 to 81 mmol m$^{-2}$ d$^{-1}$ (the lower value being measured during the beginning of the incubation and the higher towards the end of the same 3h-benthic chamber incubation), most probably reflecting temporal variations in the seepage of methane during the period of incubation.

### 3.1.7 In situ TOU measurements

The REGAB megafauna contributed substantially to the benthic oxygen uptake measured at all fauna-populated sites (Table 2). Within the clam patches TOU was on average one order of magnitude higher (50–590 mmol m$^{-2}$ d$^{-1}$) relative to the bare sediment sites (12–18 mmol m$^{-2}$ d$^{-1}$). The TOU measured at the bare sediment site (Muss_S_Env) adjacent to the mussel patch was approximately four times higher (70 mmol m$^{-2}$ d$^{-1}$) than the other bare sediment sites, and was almost as high as the TOU detected at the Muss_S site (94 mmol m$^{-2}$ d$^{-1}$).
3.1.8 In situ oxygen microsensor measurements

The Bacter_N site had the highest Diffusive Oxygen Uptake (DOU; 13 mmol m\(^{-2}\) d\(^{-1}\)) and shallowest Oxygen Penetration Depth (OPD; 2 mm) compared to other investigated sediment sites (Table 3). The diffusive fluxes of oxygen determined on bare sediments were not uniform, and fluxes changed drastically in dependence on the vicinity of the measurement site to a clam patch. In general measurements performed on bare sediment found more than 1 m away from a clam patch, showed deep OPD and low DOU (35–13 mm and 3 mmol m\(^{-2}\) d\(^{-1}\), respectively), while similar sites located app. 20 cm from a clam patch exhibited shallower OPD (2–3 mm) and much higher DOU (6–11 mmol m\(^{-2}\) d\(^{-1}\)), most probably revealing hotspots of activity in the surrounding of the clam patch.

3.2 Characterization of bacterial communities at REGAB

3.2.1 Bacterial cell numbers

The number of single cells varied substantially among habitats at REGAB. The Muss_S site had up to two times higher (2.7 \(\times\) 10\(^{10}\) cm\(^{-2}\) sediment) total integrated (0–10 cm) cell counts compared to the other fauna-populated sites (0.7–1.4 \(\times\) 10\(^{10}\) cm\(^{-2}\) sediment) (Table 2). Unusually low cell numbers were detected below the bacterial mat (0.7 \(\times\) 10\(^{10}\) cm\(^{-2}\) sediment). A sharp decrease in cell numbers with depth was detected at both mussel-related sites (Muss_S and Muss_S_Env), while at all other sites the cell numbers remained more or less constant over the entire investigated sediment depth (Fig. 3). No major differences could be observed among the populated sites and their respective adjacent bare sediment sites, except at REGAB_N area where the Clam_N site had a two times lower cell number integrated over depth (0.7 \(\times\) 10\(^{10}\) cm\(^{-2}\) sediment) compared to the Clam_N_Env site. The Muss_S_Env was the site with highest cell numbers (3.5 \(\times\) 10\(^{10}\) cm\(^{-2}\) sediment) at REGAB.
3.2.2 Bacterial community structure

The analysis of occurrence, abundance and distribution of bacterial types at the REGAB cold seep was based on defining Operational Taxonomic Units (OTU) representing relatively abundant bacterial populations as detected with the ARISA fingerprinting method (Brown and Fuhrman, 2005; Hewson and Fuhrman, 2006; Böer et al., 2009). Within individual habitats, all horizons sampled from the top 10 cm sediment depth shared on average only 30% of the OTUs, with maximum similarity detected at the Gas site (40%) and minimum at the Mussel_S (18%). In general, no depth related pattern in the OTU richness was observed, except for the mussel sites, Clam_SW and Clam_S_Env where a decline in the percentage of shared OTUs with increasing depth was revealed. Overall 450 unique OTUs were detected across all 11 sites investigated here, with the maximum number of unique OTUs per single site (401) in the top 10 cm layer at Bacter_N, and the minimum number (252) at Mussel_S_Env. All OTUs occurred at least at 2 sites, and 24% occurred at all sites. The megafauna and bacterial mat populated sites shared on average 74% of their OTUs with the adjacent bare sediments (Supplement Table S2). In general the adjacent sites had more OTUs in common compared to more distant sites (Supplement Table S2). The percentage of shared OTUs among samples at the clam-populated sites was not significantly different from the shared OTUs among samples at the non-clam-populated sites (Mann-Whitney U-test for unmatched samples \( W = 4, p = 1 \)) (Supplement Table S3). Finally, no clam habitat-specific bacterial signature was revealed, as shown by the comparison of shared OTUs by all clam patches versus the OTUs present at all other investigated sites at REGAB (Supplement Figs. S2 and S3).

The bacterial community structure of samples from two adjacent sites i.e. samples from within a patch and samples from the respective adjacent bare sediment, grouped very close to each other, (Fig. 4), indicating high similarities between the patches and the corresponding adjacent bare sediments (here defined as one habitat). Different habitats had different bacterial community structures (Table 4, Fig. 4). The mussel
habitats, characterized by highest fluxes and consumption rates of methane and sulphate (Table 2), had a very distinct bacterial community structure, which was significantly different from all other sites (Fig. 4, Table 4). The bacterial mat habitat, where intermediate levels of geochemical fluxes were detected (Table 2) had a rather similar bacterial community with the low-geochemical flux habitats, the Clam_N and Clam_SW habitats (Fig. 4, Table 4). Although, all clam habitats had similar low fluxes (Table 2), it was shown that the Clam_S habitat had a significantly different bacterial community structure from the other two clam habitats (Clam_N and Clam_SW) (Table 4, Fig. 4).

Bacterial β-diversity (change in community structure and/or composition among sites calculated using the Bray-Curtis index; Whittaker, 1960) at REGAB was significantly positively correlated to differences in CH$_4$ effluxes among sites, as revealed by the Mantel correlation test (Table 5). Marginally significant positive relations (the relation was not significant when corrected for multiple comparisons, applying the Bonferroni correction) was also revealed between β-diversity and differences in integrated AOM and SR rates, as well as the alkalinity fluxes, but not to any other single porewater parameter (Table 5).

Variation partitioning analysis performed on the full data set, including all porewater concentrations and ARISA samples from all sediment depths, showed that environmental variables comprised under “sediment geochemistry” (21%, $p = 0.001$), “space” – geographic distances among sites (7%, $p = 0.001$), and “sediment depth” (0 to 10 cm, in 1 cm horizons) (2%, $p = 0.012$) explained most of the variation in the bacterial community structure of REGAB habitats (Fig. 5a). The combined effect of “sediment geochemistry” and “space” accounted for additional 5% of the observed variations (Fig. 5a). Additional variation partitioning analysis aiming to disentangle the individual effects of the geochemical parameters considered above, revealed that methane sediment concentrations (5%, $p = 0.001$) significantly explained the highest portion of the variability in the bacterial community at REGAB including all sediment depth layers (Fig. 5b). In contrast, sulphide concentrations alone did not significantly account for the variation in the bacterial community, but the confounding effect with the
other geochemical parameters explained 6% of the observed diversity shifts (Fig. 5b). A small portion of the variation in the bacterial community structure at REGAB was related to shifts in alkalinity (1%, $p = 0.038$). Finally, the presence of clams had a very small direct overall effect, though not statistically significant, on the bacterial community structure at REGAB (1%, $p = 0.106$) (Fig. 5a). Accordingly, the comparison of the community dispersions among the clam-populated sites and the adjacent bare sediment sites revealed no significant differences (Supplement Table S4).

4 Discussion

The giant pockmark REGAB is among the best studied deep-water cold seeps in terms of biogeographical and geobiological processes shaping faunal communities (Sibuet and Olu-Le Roy, 2002; Sibuet and Vangriesheim, 2009; Cordes et al., 2010). REGAB is an endmember of the Atlantic Equatorial Belt (AEB) – the longitudinal connection from Costa Rica to the continental margin off West Africa comprising many types of reduced, chemosynthetic ecosystems. Comparison of REGAB fauna to that of other cold seep sites from the AEB has shown that water depth, rather than geographic distance shapes the distribution of megafauna, not only within biogeographic provinces (Sahling et al., 2003; Cordes et al., 2007), but also on much larger spatial scales (Sibuet and Olu, 1998; Sibuet and Olu-Le Roy, 2002; Olu et al., 2010). Furthermore, several of the key chemosynthetic species show a broad distribution, such as two amphi-Atlantic Bathymodiolus species complexes, indicating high past and/or present dispersal capabilities of these organisms across the AEB within bathymetric zones (Olu-Le Roy et al., 2007b). Evolutionary history of ocean basins, dispersal with water-masses and predation pressure have been assumed to be the broad-scale factors controlling the depth-stratification of seep chemosynthetic communities on a global scale.

On the regional to local scale, previous investigations have matched the distribution of the benthic communities to specific bottom water conditions indicative of methane seepage, and the presence of chemosynthetic megafauna (Olu-Le Roy et al., 2007a;
Cambon-Bonavita et al. 2009; Olu et al., 2009; Van Gaever et al., 2009; Menot et al., 2010). This study combines a detailed biogeochemical description of the chemosynthetic megafauna habitats with their bacterial biodiversity, and shows that different biological habitats at REGAB are linked to distinct biogeochemical regimes of the underlying sediment. The main aim was to evaluate the major factors shaping the structure of these communities and to gain a better understanding of the complexity and heterogeneity of cold seep ecosystems caused by the interplay of geochemistry, faunal and microbial distribution.

4.1 Methane flux and its subsurface microbial consumption shapes habitats at the REGAB giant pockmark

Previous studies of the composition of seep fauna communities on local and regional spatial scales revealed the role of energy availability – in the form of methane, sulphide and oxygen fluxes, along with the type of seafloor substrate (Olu-Le Roy et al., 2007a; Sahling et al., 2008; Olu et al., 2009; Van Gaever et al., 2009; Levin et al., 2010; Menot et al., 2010; Decker et al., 2012; Fischer et al., 2011; Ritt et al., 2011). Respective hypotheses have also been tested at the giant REGAB pockmark, which shows a high spatial heterogeneity and non-overlapping habitats dominated by different chemosynthetic megafauna (Olu-Le Roy et al., 2007a). Megafauna distribution and densities could be linked to local variations in energy availability, in particular to methane concentrations in the bottom water, and most probably underlying gas emissions (Olu-Le Roy et al., 2007a). Also the diversity and density of non-symbiotic megafauna, macro- and meiofauna was found to be influenced by the habitat heterogeneity caused by variations in bottom water geochemistry as well as by the presence and activity of symbiotic megafauna (Olu et al., 2009; Van Gaever et al., 2009; Menot et al., 2010). Here we tested this hypothesis both for the distribution of chemosynthetic megafauna as well as for the associated bacterial community using high resolution in situ and ex situ measurements of methane fluxes from the seafloor, microbial methane consumption (AOM) and associated biogeochemical activities.
Similar to other cold seeps, the REGAB pockmark comprises highly reduced, patchy habitats where due to the local upward transport of hydrocarbons oxygen is completely consumed within the first millimetres of seafloor (Beer et al., 2006; Girnth et al., 2011; Lichtschlag et al., 2010a; Menot et al., 2010; Grünke et al., 2011). The strongest gas venting in the form of bubble streams and outcropping hydrates were observed at the central carbonate cements (Fig 2a). Previous investigations of deep water cold seeps have shown that free gas may escape from the seafloor within the gas hydrate stability zone, even at such high pressure and cold temperature as at REGAB (Suess et al., 1999; Bohrmann et al., 2003; Greinert et al., 2006; Klaucke et al., 2006 Sauter et al., 2006; Fischer et al., 2011). Methane concentrations in the bottom waters were highest at the Mussel_S (3.6 µM), around 0.4 µM at Clam_S to and decreased to 0.2 µM at the clam habitat (Clam_SW) furthest away from the central gas vents. These values fall into the low range of values detected previously (Duperron et al., 2005; Olu-Le Roy et al., 2007a).

For the first time we measured in situ methane fluxes from the sedimentary seafloor at REGAB using benthic chambers. Highest methane efflux and also highest methane consumption rates were found at the mussel habitat of REGAB (Mussel_S and Mussel_M_Env). This *Bathymodiolus* type hosts sulphur- and methane-oxidizing endosymbionts and hence depends mostly on methane (Duperron et al., 2011). The methane and sulphate consumption rates in the sediments, as well as sulphide fluxes and to certain extent TOU were 0.5–4 times higher compared to other habitats at REGAB. The high fluxes concomitantly supported up to two times higher numbers of bacterial cells at this site relative to the other habitats, and a dense colony of mussels interspersed with tubeworms (Fig. 2b). In a previous study at REGAB, Cambon-Bonavita et al. (2009) detected highest abundances of ANME/SRB aggregates – a microbial consortium shown to mediate AOM coupled to SR processes (Boetius et al., 2000), in the sediments inhabited by mussels. Despite the high AOM rates, the mussel habitat was the only site where extensive seepage of gaseous and dissolved methane was observed, indicating substantial transport of methane to the bottom water. The AOM...
process removed only 6–20% of the total upward diffusing methane. Accordingly, the highest bottom water methane concentrations were also detected previously above mussel patches at REGAB (Charlou et al., 2004).

Sediments covered by bacterial mats, a common feature of many cold seeps (Treude et al., 2003; Niemann et al., 2006; Lessard-Pilon et al., 2010; Fischer et al., 2011; Grünke et al., 2011), were rather scarce and restricted only to the northern part of REGAB. The sediment below the bacterial mat exhibited intermediate levels of AOM, SR rates and H$_2$S fluxes, approximately 2 times higher than in the clam patches, but lower relative to the mussel patch. At this habitat, SR rates differed by an order of magnitude and SR and AOM rate did not match well, probably reflecting substantial spatial and temporal heterogeneity in the methane transport, as previously observed in relation with bacterial mats and other reduced habitats associated with hydrates (Treude et al., 2003; Lichtschlag et al., 2010b). However, both sites sampled within the bacterial mat habitat were characterized by low single cell counts in comparison to the other habitats at REGAB. In corroboration with our results, Cambon-Bonavita et al. (2009) found that ANME/SRB aggregates were the least abundant in the sediments covered by bacterial mat at the REGAB pockmark, which is a striking contrast to other cold seep settings (Lösekann et al., 2007; Girmth et al., 2011).

The clam patches were the most widely distributed sedimentary chemosynthetic habitats at REGAB. The clam species *Calyptogena regab* and *Laubiericoncha chuni* have sulphide-oxidizing symbionts and are not known to use methane directly (von Cosel and Olu, 2009; Krylova and Sahling, 2010). Vast areas within the REGAB pockmark were littered with shells of dead clams, indicating a long-term association and turnover of this bottom-dwelling, mobile chemosynthetic megafauna. The water depth at REGAB is above the calcite compensation depth, hence bivalve shells are not dissolved, and may accumulate over long times. Independent of their location within REGAB, all clam patches were associated to sediments with similar geochemistry. These sediments were characterized by lowest CH$_4$ and SO$_4$ consumption rates, as well as lowest H$_2$S fluxes. Very little to no methane was escaping the sediment at the clam
habitats, and AOM accounted for 50–80 % of the upward removal of methane. Lowest bottom water methane concentrations were detected above clam patches also during previous studies at REGAB (Olu-Le Roy et al., 2007a). A common feature of all clam habitats was the absence of sulphide from the topmost surface (up to 5 cm) sediment layers. This appears to be a universal characteristics for clam beds at cold seeps, as it has been also shown for other seeps throughout the world i.e. Northern California Seeps (Levin et al., 2003) Cascadia Convergent margin seeps (Sahling et al., 2002), Monterey Bay cold seeps (Barry et al., 1997), Makran accretionary prism (Fischer et al., 2012). Apparently, the bottom dwelling activity of the clams enables them to populate cold seep habitats with low gas fluxes and hence low microbial activity, so that they dwell the subsurface sediments to exploit rather deep peaks in sulphide production via AOM (Fischer et al., 2012).

The gas bubble site (Gas) was the only other sedimentary site, apart for the mussel habitat, where escape of gas into the water column was observed. Although the AOM and SR rates, as well as sulphide fluxes were in the range of values measured at the clam habitats, the gas bubble site was completely devoid of any visible megafauna. The relatively deep production of sulphide (≥9 cm sediment depth) and the low surface methane concentrations can potentially limit the dispersal of megafaunal organisms, which for their survival need more or less constant supply of energy sources i.e. sulphide and/or methane. It is possible that we had sampled a relatively fresh gas vent, which was not yet populated by the slow growing AOM communities transforming methane to sulphide and fuelling other chemosynthetic megafauna.

In this study we could show that the REGAB habitats differ in their methane efflux and porewater geochemistry, and that they are associated with different types of megafauna. However, methane efflux was relatively similar between the megafauna patches and their direct surroundings. In contrast, oxygen fluxes and oxygen penetration depths were different between the bare sediments and those populated by bacterial mats. Based on the biogeochemical analyses from this study, the sedimentary habitats at REGAB can be grouped in three categories: (1) mussel-associated sediments –
characterized by highest methane efflux, methane consumption and bacterial counts; (2) bacterial mat-associated sediments – characterized by intermediate activity and low bacterial counts (3) clam-associated sediments – characterized by no methane efflux, intermediate methane consumption and bacterial counts. Overall, the megafauna distribution reflects the underlying sediment characteristics, thus we propose that the megafauna assemblages can be used as reliable first visual indicator of the sediment geochemistry at cold seeps i.e. of the magnitude of methane and oxygen fluxes, and the depth of sulphide production within the sediments.

In comparison to other cold seeps, the REGAB pockmark is a relatively unique ecosystem due to the co-occurrence of diverse symbiont-bearing megafauna (one species of siboglin tubeworms and three bivalve species) as well as occasional thiotrophic bacterial mats, which form highly specialized non-overlapping assemblages (except for mussels and tubeworms in some areas). The only other seep sites harbouring similar chemosynthetic habitats have been also found in the Congo Fan area, in the vicinity of REGAB (Sahling et al., 2008). However, in general, the magnitude of the measured fluxes and geochemical processes at REGAB fall within the range of values previously reported from other gas hydrate- and carbonate-bearing cold seeps (Table 6). A striking exception is the total oxygen uptake, detected within the dense clam patches at REGAB, which to our knowledge represents the highest oxygen consumption ever measured in a cold seep environment (Table 6). The highest flux of dissolved methane at REGAB, detected at the mussel habitat, was in the range of maximum fluxes reported for bacterial mats at the Hydrate Ridge (Sommer et al., 2006), ampharetid polychaete habitats at the Hikurangi margin (Sommer et al., 2010) and the summit of Dvurechenskii Mud Volcano (DMV; Lichtschlag et al., 2010b; Table 6). AOM and SR rates can vary substantially among different cold seeps, but as well among different habitats within a single cold seep, as it is the case of the REGAB pockmark. The consumption of methane and sulphate at all investigated habitats at REGAB was quite low in comparison to other cold seeps i.e. Hydrate Ridge (Treude et al., 2003), DMV
(Lichtschlag et al., 2010b), with respectively ten and five times lower rates relative to maximum reported rates from other cold seeps (Table 6).

4.2 Methane flux influences bacterial community structure at REGAB

In addition to investigating the link between methane fluxes and the distribution of chemosynthetic megafauna habitats, we aimed at testing if (a) the bacterial community structure differs between the different habitats, and (b) if the underlying patterns in bacterial biodiversity could be linked to methane fluxes. In this regard, we used several independent measures of seepage activity: the magnitude of methane effluxes, AOM coupled to sulphate consumption rates, sulphide fluxes and alkalinity produced by AOM, and other associated geochemical variables (Table 5; Fig. 3; Supplement Fig. S1). In different ways, these biogeochemical processes are indicative of potential energy availability to the seep communities, with methane and sulphide representing the major sources of reduced chemical energy. Of the different indicators of bacterial diversity analysed here, the bacterial β-diversity – change in species composition and/or structure between habitats (Whittaker, 1960) – was foremost significantly positively correlated to differences in in situ methane effluxes. A positive trend existed also between the bacterial β-diversity and the difference in core geochemical processes AOM and SR, as well as alkalinity flux, which respond usually directly to variations in methane supply via AOM or via subsurface fluid advection (Bohrmann et al., 1998; Valentine, 2002; Luff and Wallmann, 2003). These results support the hypothesis that the bacterial community structure at cold seeps is influenced foremost by methane supply, as primary source of energy to anaerobic and aerobic methanotrophs (Cambon-Bonavita et al., 2009), and as a main indicator of the activity of geological processes such as gas overpressure, fluid flow and hydrate formation or dissociation.

Surprisingly, even though a much higher diversity of bacteria and animals could be biologically influenced by sulphide as energy source or as toxin, bacterial β-diversity was not significantly correlated to difference in sulphide fluxes among habitats. Sulphide is a secondary energy source, provided via microbial methane consumption with
sulphate, and it is possible that this indirect link to methane blurs relations with diversity indicators. However, generally, habitat types with similar biogeochemistry/energy availability were more similar in terms of bacterial diversity as opposed to habitats with distinct geochemistry.

A further analysis of the links between community structure and concentrations of porewater constituents across all individual depth samples also confirmed a link between β-diversity and sediment geochemistry (Fig. 5). Even when using ex situ methane concentrations in the analysis (i.e. after degassing and depressurization of the cores upon retrieval), it explained the highest proportion of the variation in the bacterial community structure. Again, sulphide concentrations could not explain any variation in the bacterial community structure, however, the confounding effect of this variable with the rest of the geochemical parameters accounted for a substantial portion of the variation in the community structure at REGAB, as did sediment depth and geographic distance between samples (Fig. 5). Geographic (spatial) distance among sampling sites and to lesser extent sediment depth of individual samples are two other variables that appeared to play a role in shaping the bacterial communities at REGAB. Of course, both are coupled strongly also to geological and geochemical processes, e.g. distance to gas seepage, or upward transport of highly reduced porewater fluids to the sulphate or oxygen penetrated sediment surface layer, and their role in structuring bacterial diversity cannot be further disentangled in this study.

Although only a small proportion of the total bacterial community can use methane directly as energy source, we propose that methane oxidation, as the primary energy producing process in the seep ecosystem, is a main driver of community structure, from bacteria to megafauna. Previous studies of β-diversity patterns of benthic bacterial communities have shown significant relationships to geobiological indicators of energy availability, such as phytodetritus sedimentation to oligotrophic continental margins (Bienhold et al., 2011), and benthic primary productivity in coastal sands (Böer et al., 2009).
4.3 Link between megafauna, geochemistry and bacterial community structure

Our data indicate that methane fluxes determine sediment geochemistry, which selects for different types of chemosynthetic megafauna at REGAB. A further aim of this study was to test if the distribution of the chemosynthetic megafauna is also influencing bacterial community structure. It has been shown previously that chemosynthetic megafauna influences its local environment by bioturbation, bioirrigation, burrowing and by exudates, altering the local seafloor biogeochemistry (Barry et al., 1997; Levin et al., 2003; Fischer et al., 2011). Clam and mussel respiration accounted for a substantial local increase in the total benthic oxygen uptake rates, as compared to the adjacent bare sediments. Furthermore, the *Bathymodiolus* mussels contain methanotrophic symbionts and consume methane efficiently (Duperron et al., 2009, 2011; Petersen and Dubilier, 2009), causing a reduction of methane efflux within the mussel patch.

Similar to sites populated by clams at other cold seeps, sulphide was absent from the surface sediments (Barry et al., 1997; Sahling et al., 2002; Levin et al., 2003; Fischer et al., 2012) and bottom water sulphate penetrated till 6 cm depth at all clam patches at REGAB – a signature ascribed to the dwelling activity of thiotrophic clams to access sulphide (Childress and Fisher, 1992), leading to a deeper sulphate penetration (Sahling et al., 2002; Cordes et al., 2005, 2010; Fischer et al., 2012). Furthermore, unusual and elevated ammonium concentrations were measured below the clam patches, where the maximum concentrations – indicating local production of ammonium, were observed within the depth range of the clam’s foot. Previous studies from the Gulf of Mexico cold seeps also showed that clam-associated sediments had very high ammonium concentrations (Joye et al., 2010). These results indicate a potential link between the clam presence and the NH$_4$ concentrations in pore water, most probably via local excretions of metabolites by the clams, but further investigations are needed to decipher this relation. However, when comparing the structure of bacterial communities of adjacent sites at REGAB with and without chemosynthetic megafauna or bacterial mats, both methane fluxes and bacterial community structure were similar (Fig. 4,
Table 2). Accordingly, no direct association of unique bacterial types with the different megafauna was detected. Even the dense clam patches had little effect on the bacterial community structure. This indicates that the abundant bacterial types in this cold seep ecosystem as detected by ARISA fingerprinting were directly affected by methane seepage and other geochemical processes, but only indirectly by the presence and absence of megafauna types. This finding may differ with other types of molecular methods, which include rare bacterial types. Furthermore, space played an important role in structuring the distribution and diversity of chemosynthetic megafauna and bacterial communities at spatial scales of meters to hundreds of meters, which needs further investigation, especially when considering the need for conservation, protection and management of cold seeps as unique deep-water ecosystems.

Supplementary material related to this article is available online at: http://www.biogeosciences-discuss.net/9/8337/2012/bgd-9-8337-2012-supplement.pdf.

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### Table 1. Overview of the sampling sites, their geographic position and type of measurements performed within this study. Samples and measurements were taken in July–August 2008. AOM = Anaerobic Oxidation of Methane, SR = Sulphate Reduction, MICP = microprofiler measurement, CHAM = Benthic Chamber Incubation.

<table>
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<tr>
<th>Location</th>
<th>Sampling site</th>
<th>Latitude</th>
<th>Longitude</th>
<th>Measurement; sample</th>
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<td>E 9°42.636′</td>
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Table 2. Biogeochemical characterization of different habitats at REGAB. Maximum H$_2$S flux in the sediment, CH$_4$ efflux, TOU, average integrated (0–10 cm sediment depth) AOM and SR rates, methane consumption efficiency calculated as the percentage of methane consumption (AOM) from the total methane flux (AOM + CH$_4$ efflux), total integrated (0–10 cm sediment depth) single cell numbers, as well as alkalinity flux and modelled values of bioirrigation and advective flow.

<table>
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<th>CH$_4$ efflux (mmol m$^{-2}$ d$^{-1}$)</th>
<th>TOU (mmol m$^{-2}$ d$^{-1}$)</th>
<th>AOM (mmol m$^{-2}$ d$^{-1}$)</th>
<th>Methane consumption efficiency (%)</th>
<th>SR (mmol m$^{-2}$ d$^{-1}$)</th>
<th>Bioirrigation ($\times 10^{-6}$) (m s$^{-1}$)</th>
<th>Single cells ($\times 10^{15}$) (cm$^{-2}$)</th>
<th>Alkalinity flux (mmol m$^{-2}$ d$^{-1}$)</th>
<th>Advective flow ($\times 10^{-8}$) (m s$^{-1}$)</th>
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<td>&lt;1</td>
<td>4</td>
<td>0.7</td>
<td>0.1</td>
<td>1.4</td>
<td>0</td>
</tr>
<tr>
<td>Clam_N_Env</td>
<td>9</td>
<td>n.d.</td>
<td>n.d.</td>
<td>5</td>
<td>n.d.</td>
<td>4</td>
<td>1.5</td>
<td>0.3</td>
<td>1.0</td>
<td>0</td>
</tr>
<tr>
<td>Clam_S</td>
<td>6</td>
<td>1</td>
<td>590</td>
<td>6</td>
<td>81</td>
<td>6</td>
<td>0.9</td>
<td>0.1</td>
<td>5.0</td>
<td>0</td>
</tr>
<tr>
<td>Clam_S_Env</td>
<td>0</td>
<td>3</td>
<td>18</td>
<td>3</td>
<td>47</td>
<td>2</td>
<td>1.0</td>
<td>0.1</td>
<td>0</td>
<td>0</td>
</tr>
<tr>
<td>Mussel_S</td>
<td>20</td>
<td>1–81</td>
<td>94</td>
<td>19</td>
<td>19</td>
<td>28</td>
<td>2.7</td>
<td>0.4</td>
<td>0</td>
<td>0</td>
</tr>
<tr>
<td>Mussel_S_Env</td>
<td>23</td>
<td>334</td>
<td>77</td>
<td>20</td>
<td>6</td>
<td>36</td>
<td>3.5</td>
<td>0.4</td>
<td>6.0</td>
<td>0</td>
</tr>
<tr>
<td>Gas</td>
<td>7</td>
<td>n.d.</td>
<td>n.d.</td>
<td>4</td>
<td>n.d.</td>
<td>8</td>
<td>1.6</td>
<td>0.2</td>
<td>0</td>
<td>0</td>
</tr>
<tr>
<td>Clam_SW</td>
<td>5</td>
<td>&lt;1</td>
<td>294</td>
<td>3</td>
<td>81</td>
<td>10</td>
<td>1.4</td>
<td>0.0</td>
<td>0</td>
<td>0</td>
</tr>
<tr>
<td>Clam_SW_Env</td>
<td>0</td>
<td>0</td>
<td>12</td>
<td>2</td>
<td>97</td>
<td>1</td>
<td>1.0</td>
<td>0.0</td>
<td>0</td>
<td>0</td>
</tr>
</tbody>
</table>

Table 3. Sediment oxygen penetration depth (OPD) and diffusive oxygen uptake (DOU) of sediments in relation to the vicinity of clam patches and bacterial mat at REGAB, as measured in situ with a microprofiler.

<table>
<thead>
<tr>
<th>Distance to clam patch/bacterial mat</th>
<th>OPD (mm)</th>
<th>DOU (mmol m(^{-2}) d(^{-1}))</th>
</tr>
</thead>
<tbody>
<tr>
<td>Bacter_N_Env Within bacterial mat</td>
<td>2</td>
<td>13</td>
</tr>
<tr>
<td>Clam_S_Env &lt;1 m from clam patch</td>
<td>2</td>
<td>6</td>
</tr>
<tr>
<td>Clam_SW_Env &lt;1 m from clam patch</td>
<td>3</td>
<td>11</td>
</tr>
<tr>
<td>Clam_N_Env &gt;1 m from clam patch</td>
<td>35</td>
<td>3</td>
</tr>
<tr>
<td>Clam_SW_Env &gt;1 m from clam patch</td>
<td>13</td>
<td>3</td>
</tr>
</tbody>
</table>
Table 4. Analysis of Similarity (ANOSIM; lower triangle), testing for significant differences in bacterial community structure between habitats and percentage of shared OTUs between habitats (upper triangle). ANOSIM R-values (lower triangle) are interpreted as follows: $R < 0.25 =$ strongly overlapping, $0.25 < R < 0.5 =$ separated but with overlap, $0.5 < R < 0.75 =$ separated with only minor overlap, $R > 0.75 =$ strongly separated. $^* p < 0.05$, $^{**} p < 0.01$, $^{***} p < 0.001$.

<table>
<thead>
<tr>
<th></th>
<th>Bacter_N</th>
<th>Clam_N</th>
<th>Mussel_S</th>
<th>Clam_S</th>
<th>Clam_SW</th>
</tr>
</thead>
<tbody>
<tr>
<td>Bacter_N</td>
<td>0.3</td>
<td></td>
<td>0.7**</td>
<td>0.8***</td>
<td>0.5**</td>
</tr>
<tr>
<td>Clam_N</td>
<td>92</td>
<td>77</td>
<td>80</td>
<td>87</td>
<td></td>
</tr>
<tr>
<td>Mussel_S</td>
<td>78</td>
<td></td>
<td>78</td>
<td>70</td>
<td>76</td>
</tr>
<tr>
<td>Clam_S</td>
<td>78</td>
<td>78</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Clam_SW</td>
<td>0.4*</td>
<td>0.8***</td>
<td></td>
<td></td>
<td>0.7**</td>
</tr>
</tbody>
</table>
Table 5. Mantel test, checking for correlation between β-diversity (calculated as Bray-Curtis dissimilarities) and difference in CH₄ efflux, Total Oxygen Uptake (TOU), integrated Anaerobic Oxidation of Methane (AOM) and Sulphate Reduction (SR) rate, alkalinity flux, advective flow, integrated sulphide flux and bioirrigation rate. The spearman rank coefficient was used for calculating correlations. * = p < 0.01; Bonferroni-correction was applied to correct for multiple testing. (**) only significant without Bonferroni correction. Mantel correlation test was performed on pooled ARISA samples (0–10 cm depth) according to habitat. For this analysis data from the following sites was used: Clam_N, Mussel_S, Mussel_S_Env, Clam_S, Clam_S_Env, Clam_SW, Clam_SW_Env.

<table>
<thead>
<tr>
<th></th>
<th>CH₄ efflux</th>
<th>TOU</th>
<th>AOM</th>
<th>Bioirrigation rate</th>
<th>Alkalinity flux</th>
<th>Advective flow</th>
<th>Integrated H₂S flux</th>
<th>SR</th>
</tr>
</thead>
<tbody>
<tr>
<td>Mantel statistics</td>
<td>0.6**</td>
<td>-0.1</td>
<td>0.5(*)</td>
<td>0.2</td>
<td>0.5(*)</td>
<td>0.2</td>
<td>0.4</td>
<td>0.5(*)</td>
</tr>
<tr>
<td>Significance p-value</td>
<td>0.003</td>
<td>0.641</td>
<td>0.048</td>
<td>0.265</td>
<td>0.021</td>
<td>0.222</td>
<td>0.076</td>
<td>0.037</td>
</tr>
<tr>
<td>Bonferroni corrected p-value</td>
<td>0.021</td>
<td>1</td>
<td>0.336</td>
<td>1</td>
<td>0.147</td>
<td>1</td>
<td>0.532</td>
<td>0.259</td>
</tr>
</tbody>
</table>
Table 6. Compilation of in situ CH$_4$ effluxes and TOU, ex situ average integrated AOM, SR rates, as well as H$_2$S fluxes from cold seep sites worldwide, including data from REGAB (this study).

<table>
<thead>
<tr>
<th>Cold seep</th>
<th>Habitat</th>
<th>CH$_4$ efflux (mmol m$^{-2}$ d$^{-1}$)</th>
<th>AOM (mmol m$^{-2}$ d$^{-1}$)</th>
<th>SR (mmol m$^{-2}$ d$^{-1}$)</th>
<th>H$_2$S flux (mmol m$^{-2}$ d$^{-1}$)</th>
<th>TOU (mmol m$^{-2}$ d$^{-1}$)</th>
<th>Reference</th>
</tr>
</thead>
<tbody>
<tr>
<td>REGAB</td>
<td>Bacterial mat</td>
<td>n.d.</td>
<td>9</td>
<td>23</td>
<td>13</td>
<td>n.d.</td>
<td>This study</td>
</tr>
<tr>
<td>(3160 m water depth)</td>
<td>Clam patch</td>
<td>1</td>
<td>3 (±2.5)</td>
<td>8 (±2.5)</td>
<td>5 (±0.6)</td>
<td>311 (±294)</td>
<td></td>
</tr>
<tr>
<td></td>
<td>Mussel patch</td>
<td>1-81</td>
<td>19</td>
<td>28</td>
<td>20</td>
<td>94</td>
<td></td>
</tr>
<tr>
<td>HMMV$^a$ (1250 m</td>
<td>Beggiaatoa mat</td>
<td>78</td>
<td>10 (±6.6)</td>
<td>14 (±6.2)</td>
<td>101; 114</td>
<td>101; 114</td>
<td>Felden et al. (2010)</td>
</tr>
<tr>
<td>water depth)</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Hydrate Ridge (600–1000 m</td>
<td>Beggiaatoa mat</td>
<td>5.7; 30; 90</td>
<td>5 (±4.4); 99 (±102)</td>
<td>32 (±34)</td>
<td>23 (±13)</td>
<td>48</td>
<td>Treude et al. (2003); Sahling et al. (2002);</td>
</tr>
<tr>
<td>water depth)</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td>Sommer et al. (2006); Torres et al. (2002)</td>
</tr>
<tr>
<td></td>
<td>Clam patch</td>
<td>0.6; 0.45</td>
<td>56 (±54)</td>
<td>65 (±58)</td>
<td>7 (±2.4)</td>
<td>4</td>
<td>Treude et al. (2003); Sahling et al. (2002);</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td>Sommer et al. (2006); Torres et al. (2002)</td>
</tr>
<tr>
<td></td>
<td>Acharax field</td>
<td>n.d.</td>
<td>2 (±1.4)</td>
<td>0.4 (±0.3)</td>
<td>0.05 (±0.05)</td>
<td>n.d.</td>
<td>Teude et al. (2003); Sahling et al. (2002)</td>
</tr>
<tr>
<td>DMV$^b$ (2060 m</td>
<td>Summit</td>
<td>458</td>
<td>0.07 (±0.1)</td>
<td>0.05 (±0.0)</td>
<td>n.d.</td>
<td>0</td>
<td>Lichtschlag et al. (2010)</td>
</tr>
<tr>
<td>water depth; anoxic</td>
<td>Geographical centre of DMV</td>
<td>n.d.</td>
<td>9 (±6)</td>
<td>20 (±5.7)</td>
<td>n.d.</td>
<td>0</td>
<td>Lichtschlag et al. (2010)</td>
</tr>
<tr>
<td>Black Sea)</td>
<td>Western edge</td>
<td>n.d.</td>
<td>11 (±9.6)</td>
<td>108 (±38)</td>
<td>n.d.</td>
<td>0</td>
<td>Lichtschlag et al. (2010)</td>
</tr>
<tr>
<td>Hikurangi Margin (1050 m water depth)</td>
<td>Ampharetids</td>
<td>max. 265</td>
<td>17 (±9.2)</td>
<td>n.d.</td>
<td>n.d.</td>
<td>max. 118</td>
<td>Sommer et al. (2010)</td>
</tr>
<tr>
<td></td>
<td>Tube worms</td>
<td>max. 5</td>
<td>n.d.</td>
<td>n.d.</td>
<td>max. 65</td>
<td></td>
<td>Sommer et al. (2010)</td>
</tr>
</tbody>
</table>

$^a$ Håkon Mosby Mud Volcano.
$^b$ Dvurechenskii Mud Volcano.
Fig. 1. Map of the REGAB pockmark (adopted from Ondréas et al., 2005), derived from ROV Victor 6000 data (Ifremer, France), with the main sampling locations. Locations of push core sediment sampling are depicted as red triangles, benthic chamber incubations as yellow circles and microprofiler measurements as squares.
Fig. 2. Typical habitats at the REGAB pockmark: the central area with extensive carbonate crust and outcropping hydrates populated by mussels and tubeworms (a); clam patches overlying blackened reduced sediment (b, c, f); mussel patch on soft sediment (d); white bacterial mat at the REGAB_N (e); Benthic chamber incubation is visible in (b) picture.
Fig. 3. Geochemical depth profiles of $\text{H}_2\text{S}$, $\text{SO}_4^-$, AOM, SR, $\text{NH}_4^+$, alkalinity and single cell counts at all investigated sites at REGAB. Closed symbols denote measurements taken within the patches/bacterial mat, and open symbols denote measurements taken at the nearby bare sediments. N.d. = not determined.
Fig. 4. Nonmetric MultiDimensional Scaling (NMDS) ordination plot (based on Bray Curtis distance matrix) of ARISA merged profiles (2–3 PCR replicates were pooled to form a consensus profile). Ordihull grouping of samples (0–10 cm) according to sampling location (samples from within the patches and the samples from the respective bare sediment were grouped together). NMDS stress 18 %. Open symbols denote samples from the bare sediment sites – outside of the respective patch; closed symbols denote samples from within the respective patch.
**Fig. 5.** Variation partitioning analysis. Effect (%) of the geochemistry (concentrations of CH$_4$, H$_2$S, Cl, Fe, Mn, PO$_4$, SO$_4$, pH), clams presence (clams directly influence the top 6 cm sediment), space (geographic distance among sampling sites) and depth (sediment depth) on the bacterial diversity at REGAB (a). Geochemistry, followed by space and depth significantly shape the diversity at REGAB. Visualization of the individual influence of CH$_4$ and H$_2$S on the bacterial community structure (b). Other = all parameters included in (a), excluding CH$_4$ and H$_2$S. Individually CH$_4$ ($p < 0.001$) significantly explains the variation in the ARISA dataset. *** $p < 0.001$, ** $p < 0.002$, * $p < 0.03$. The complete model explained 36% of the total variation in the bacterial diversity at REGAB.